

Mineralogy and geochemistry of Fe-Ti oxide ores from the Don Dieguito massif-type anorthosite suite, Sierra Nevada de Santa Marta, Colombia

Mineralogía y geoquímica de las menas de óxidos de Fe-Ti en la anortosita de Don Dieguito, Sierra Nevada de Santa Marta, Colombia

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ABSTRACT

Fe-Ti oxide ores are commonly associated with Proterozoic massif-type anorthosite bodies emplaced during the Grenville orogeny (~1.2-1.0 Ga). Some of these anorthositic bodies occur in the northernmost part of the Santa Marta Massif, Colombia. They locally contain crosscutting Fe-Ti(-V) ore bodies between the El Hierro creek and the Don Dieguito river. We have distinguished two types of Fe-Ti(-V) ores: i) oxide-apatite norite (fine grained ilmenite and magnetite disseminated in an assemblage of apatite, amphibole, chlorite, rutile and sericitized plagioclase) and ii) banded nelsonite (coarser grained ilmenite, magnetite and apatite distributed in bands, with minor baddeleyite, srilankite and högbomite). Ilmenite in the two ore types display distinct hematite exsolution features: a) needle-like in the oxide-apatite norite ore; and b) two generations of exsolutions in the banded ore. Magnetite shows similar Ti and V contents in both ore types (0.4 wt % TiO₂ and 0.5 wt % V₂O₃ on average in the oxide-apatite norite ore, 0.5 wt % TiO₂ and 0.4 wt % V₂O₃ in the banded ore), whereas ilmenite has higher hematite, pyrophanite and geikielite components in the banded ore. The ores might have formed from the combination of fractional crystallization and magma mixing, with exsolution occurring probably at 575-600°C. This first detailed mineralogical and textural study of Fe-Ti(-V) oxide ores in anorthosite massifs from Colombia suggests that the Santa Marta Massif is an interesting target for future research.

Keywords: Fe-Ti oxide ores, V, exsolution, titaniferous hermatite, ilmenite, magnetite, massif anorthosite, nelsonite.

RESUMEN

Las mineralizaciones de óxidos de Fe-Ti se asocian comúnmente a complejos anortositicos proterozoicos (massif-type) emplazados durante la orogenia Grenvilliana (~1.2-1.0 Ga). Algunos de estos cuerpos anortositicos se encuentran en la parte septentrional del Macizo de Santa Marta, en Colombia. Localmente, éstos contienen cuerpos de óxidos de Fe-Ti(-V) entre la quebrada El Hierro y el río Don Dieguito. Se han distinguido dos tipos de mineralizaciones de Fe-Ti(-V): i) norita con óxidos y apatito (ilmenita y magnetita de grano fino diseminadas en una asociación de apatito, anfíbol, clorita, rutilo y plagioclasa sericitizada); y ii) nelsonita bandeada (ilmenita, magnetita y apatito de grano grueso, distribuidos en bandas, con cantidades menores de baddeleyita, srilankita y högbomita). Las ilmenitas en los dos tipos de mineralizaciones muestran diferentes exsoluciones de hematitas: con forma de aguja en la mena diseminada; y dos generaciones de exsoluciones en la mena bandeada. La magnetita presenta contenidos en Ti y V similares en ambas mineralizaciones (un promedio de 0.4 wt % TiO₂ y 0.5 wt % V₂O₃ en la diseminada, 0.5 wt % TiO₂ y 0.4 wt % V₂O₃ en la bandeada), mientras que la ilmenita tiene mayor componente hematitas, pirofanita y geikielite en la mena bandeada. Las menas se podrían haber formado por la combinación de cristalización fraccionada y mezcla de magmas, y la exsolución tendría lugar probablemente a 575-600°C. Esta primera caracterización mineral y textural de los óxidos de Fe-Ti(-V) en macizos anortositicos en Colombia indica que la Sierra de Santa Marta es un objetivo interesante para investigaciones futuras.

Palabras clave: óxidos de Fe-Ti, V, exsolución, Ti-hermatita, ilmenita, magnetita, anortositas massif-type, nelsonita.

1. Introduction

Fe-Ti oxide-apatite rocks (nelsonites) are relatively scarce worldwide, and their occurrence is commonly associated with anorthosite massifs (e.g. Charlier *et al.*, 2015; Chen *et al.*, 2013; Duchesne, 1999; Dymek and Owens, 2001). In the broader Circum-Caribbean region they have been reported mainly in Mexico (e.g. Cisneros de León *et al.*, 2017; Solari *et al.*, 2003; Ortega-Gutiérrez, 1978) associated with Massif-type anorthosites that are part of Grenvillian granulite terrains, and that are considered important for the paleogeographic reconstructions of Rodinia (Ortega-Gutiérrez *et al.*, 2018 and references therein).

Fe-Ti oxide ores in northern Colombia were first described by Tschanz *et al.* (1969) in the Sierra Nevada de Santa Marta (SNSM), where they are associated with isotropic and gneissic anorthosites (Alfonso *et al.*, 1990; Tschanz, 1970). These Fe-Ti oxide ores have been generally poorly described, and their petrogenetic significance remains to be studied (Alfonso *et al.*, 1990; Cortes Castillo, 2013; Tschanz, 1970). On one hand, dense vegetation and weathering conditions represent an obstacle to find appropriate outcrops and, on the other hand, indigenous cultures who currently live and govern the region, grant researchers a restricted access in order to protect the natural resources of the area. The aim of the present contribution is to study the mineralogy and geochemistry of the Fe-Ti(-V) oxide mineralizations of SNSM.

2. Geological setting

The Sierra Nevada de Santa Marta (SNSM) mountain range is a complex massif isolated and completely independent from the Andean chain. It is located in northern Colombia and covers a range of altitudes from the sea level to close to 5,700 m.

The SNSM is composed of a series of geological units surrounded by Mesozoic and Cenozoic sedimentary basins and intruded by plutons of different ages (Piraquive, 2017; Cardona *et al.*, 2010a,

2010b; Tschanz *et al.*, 1969; 1974). The current configuration of the SNSM is the result of the Late Meso-Cenozoic east-northeast migration of the Caribbean plate, and the movement of the northern Andean block associated with the subduction of the Pacific plate (Montes *et al.*, 2010; Bayona *et al.*, 2010; Cardona *et al.*, 2010a; Pindell *et al.*, 1998). It is limited by three large regional fault systems: the Santa Marta-Bucaramanga Fault System trending SE-NW, the Media Luna Fault trending SW-NE, and the Oca Fault trending WE (Tschanz *et al.*, 1969). These three fault systems shape the range to the final triangle-like outline (Figure 1, Tschanz *et al.*, 1969). The basement of the SNSM comprises a series of Precambrian rocks older than ~1 Ga, subsequent complex Paleozoic metamorphic belts and several Mesozoic to Neogene intrusions (Tschanz *et al.*, 1969).

The SNSM has been divided into four different provinces in terms of their age, petrology and structures (Figure 1; Tschanz *et al.*, 1969; 1974; Piraquive, 2017). From northwest to southeast they are: 1) the Eocene Santa Marta Province, or Santa Marta Orogenic Belt, 2) the polymetamorphic Sevilla Arch Province or Sevilla Metamorphic Belt, 3) the Precambrian Granulite Province or Sierra Nevada Province, and 4) the Perijá Province, across the Cesar valley.

The Eocene Santa Marta Orogenic Belt consists in turn of two metamorphic units separated by the Eocene Santa Marta Batholith (Figure 1): 1) the Santa Marta Schists (also the Outer Santa Marta Metamorphic Belt), which lie to the west of the batholith and include greenschists, epidote-amphibolite facies rocks and metavolcanics; and 2) The Santa Marta Metamorphic Belt (also Inner Santa Marta Metamorphic Belt), which comprises the San Lorenzo and the Gaira Schists. These rocks are possible allochthonous fragments formed during accretionary events of the Mesozoic and Cenozoic (Bustamante *et al.*, 2009; Tschanz *et al.*, 1969).

The Sevilla Metamorphic Belt controls the current configuration of the SNSM. It is constrained by the Sevilla Lineament, a first order structure whose features limit the surface of a subduction

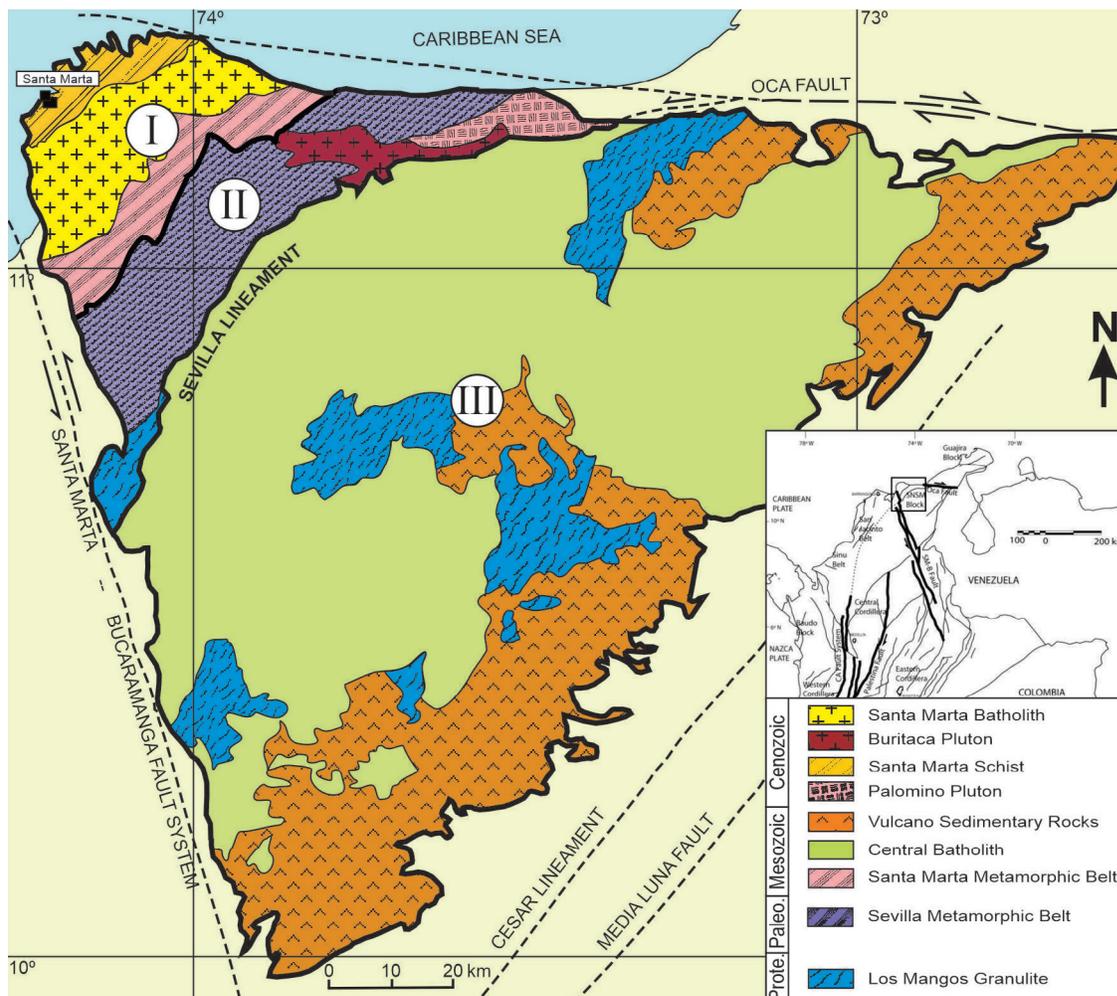


Figure 1 Geological map of the SNSM. I- Eocene Santa Marta Orogenic Belt. II- The Sevilla Metamorphic Belt. III- Precambrian granulite province (Modified after Tschanz *et al.*, 1969).

zone during the Upper Permian and Triassic (Piraquive, 2017; Cardona *et al.*, 2010a, 2010b; Tschanz *et al.*, 1969). Along this structure, mafic and anorthositic gneisses, schists and allochthonous blocks of granulites crop out, the latter making part of a Neoproterozoic arc that existed during Rodinia break-up and was later metamorphosed in Permian-Triassic times, possibly during the Laurentia-Gondwana collision (Piraquive, 2017). Different magmatic events are found to intrude the Sevilla lineament causing the formation of several parallel structures, among these, the Eocene Buritaca pluton, in contact with the Precambrian anorthosites and other Permian, Cretaceous and Paleogene magmatic bodies (Tschanz *et al.*, 1969).

Finally, the Precambrian Granulite Province (Bustamante *et al.*, 2009; Tschanz *et al.*, 1969) is the largest and oldest province in the massif. A Grenvillian oceanic crust makes up the basement of this province along with a series of gneisses, amphibolites, garnet-rich granulites and anorthosites (Tschanz *et al.*, 1969). It is intruded by Jurassic plutons and it is mostly covered by volcanosedimentary rocks of Jurassic age. The granulites are affected by a set of transcurrent faults and show wide fragmentation by the intrusion of younger igneous bodies, such as the Jurassic Central Batholith and the Eocene Buritaca Pluton. This province can be associated to the Sunsas-Putumayo province (Ibáñez-Mejía *et al.*, 2011), and represents one of the exposed Grenvillian inliers in Colombia (see review in Kroonenberg, 2018).

2.1. LOCAL GEOLOGY

The study area is located in El Hierro creek, a medium-sized tributary of the Don Dieguito river, at the northern slope of the SNSM, near the Don Diego river (Figure 2). In this location, Tschanz (1969) reported the occurrence of minor ilmenite-apatite mineralization related to the Precambrian Don Dieguito Anorthosites. The main units that outcrop along the studied area are: 1) Los Mangos granulite, 2) Don Dieguito Anorthosite, 3) Santa Marta batholith and the Buritaca pluton, covered by alluvial terraces.

2.1.1. LOS MANGOS GRANULITE

This unit forms the metamorphic basement of the major part of the SNSM with an approximate extension of 1150 km². The granulites are overlain by sedimentary and volcanic rocks and are intruded by different batholiths. The Los Mangos granulite comprises a series of parallel layers of alternating metamorphosed dark and light granulite facies gneisses, often migmatitic. They include 1) light-colored granitic granulites which include calcsilicate rocks and diopside marbles, 2) dark-colored, dense mafic and ultramafic granulites (e.g. peridotites), and 3) granulites of intermediate composition (Ordóñez-Carmona *et al.*, 2002; Alfonso *et al.*, 1990; Tschanz *et al.*, 1969).

Light grey quartz-feldspathic gneisses with a massive appearance and granoblastic texture, medium to coarse grained, make up the first of the three groups mentioned above. They are mainly composed of plagioclase (oligoclase), quartz, biotite, with sericite and chlorite as alteration minerals. The second group consists of layers of dark grey hornblende-biotite gneisses up to 50 cm in width inside the quartz-feldspathic gneisses; and are medium to coarse-grained and with a granoblastic texture. The primary minerals are hornblende, biotite, quartz and plagioclase (oligoclase), with epidote, magnetite and sericite as accessory minerals (Tschanz *et al.*, 1969). Finally, the intermediate granulites vary in composition from quartz diorite to trondjemite. Minerals of the granulite facies are

scarce and secondary minerals are abundant, such as biotite and amphibole. Plagioclase is commonly oligoclase, present with green clinopyroxene, microcline and quartz (Tschanz *et al.*, 1969). The unit is in contact with the Don Dieguito Anorthosite and the Buritaca pluton (Tschanz *et al.*, 1969).

2.1.2. DON DIEGUITO ANORTHOSITE

The Don Dieguito Anorthosite was first recognized by Tschanz *et al.* (1969) in the El Hierro creek. Together with the other anorthosites of the SNSM, it is the third occurrence of anorthosites described in South America, after the Capivarito Anorthosite in Rio Grande do Sul, Brasil (Chemale *et al.*, 2011) and the ones in Sierras Pampeanas, Argentina (Casquet *et al.*, 2005), all of them yielding Precambrian ages. The Don Dieguito anorthosite is a beige, locally light pink to brownish rock, with isotropic to foliated structure and massive appearance. Within these rocks, the ilmenite-apatite-magnetite mineralization can be observed, both layered and podiform, which gives the rock a spotted aspect. According to Tschanz *et al.* (1969), at least 80% of the rock is composed mainly of coarse-grained twinned plagioclase that ranges from labradorite to oligoclase with a proclastic texture and locally fractured; the remaining 20% is made up of augen-like mafic minerals such as hornblende and biotite displaying granoblastic texture (Alfonso *et al.*, 1990; Tschanz *et al.*, 1969). The gneissic anorthosite and accessory magnetite are considered to be plagioclase gneisses that form lenses or concordant layers within the Precambrian granulites (Alfonso *et al.*, 1990). The primary mafic minerals are completely replaced by secondary pale green actinolitic amphibole, low iron chlorite and low iron epidote. Biotite and muscovite are also common secondary minerals. Traces of quartz and calcite are locally present (Alfonso *et al.*, 1990; Tschanz *et al.*, 1969). The Don Dieguito Anorthosite is associated to a coarse-grained equigranular rock, composed mainly of oligoclase.

Tschanz *et al.* (1969) described two types of magnetite or apatite-ilmenite mineralizations in the Don Dieguito river area: 1) mafic dykes, very

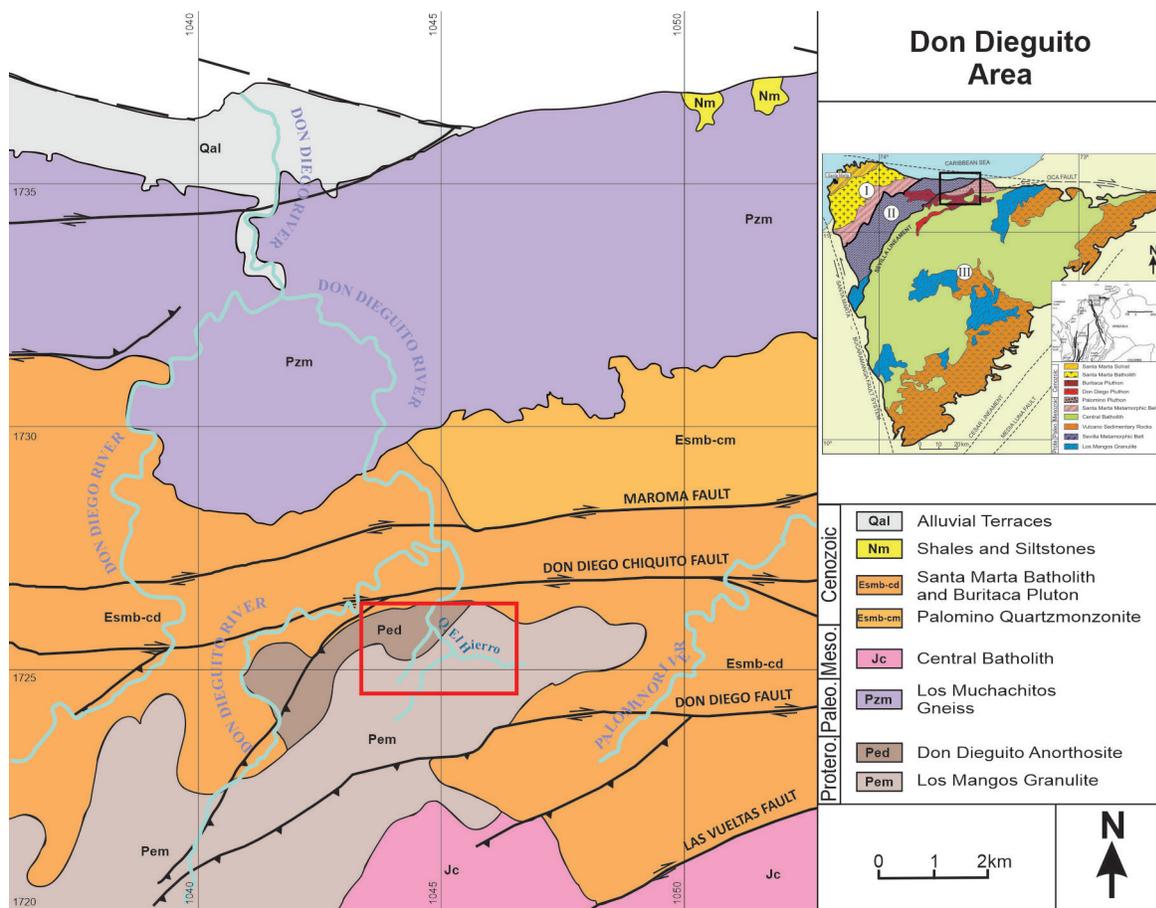


Figure 2 Geological map of the Don Dieguito river area (red area). Note the El Hierro creek in the central part of the map flowing towards the Don Dieguito river, which cuts the Los Mangos Granulite (Pem) and the Don Dieguito Anorthosite (Ped). Modified after the Colombian Geological Survey (SGC, 2019).

rich in apatite and ilmenite, and ilmenite veins with accessory magnetite; 2) a bimineralic banded apatite-ilmenite rock closely associated with the Don Diego Anorthosite. Although it is difficult to accurately locate the ilmenite-apatite mineralized zones, great amounts of boulders and cobbles of these rocks are found along the El Hierro creek.

Gabbroid rocks (mineralized norites), amphibole rich rocks, diorites and plagioclase (oligoclase) dykes locally crosscut the anorthosite. The genetic relationship between these rocks and the Don Dieguito Anorthosite is still uncertain (Alfonso *et al.*, 1990; Tschanz *et al.*, 1969). The contact between the Don Dieguito Anorthosite and the Los Mangos Granulite has not been well established yet (Tschanz *et al.*, 1969).

2.1.3. SANTA MARTA BATHOLITH AND THE BURITACA PLUTON

This unit comprises two Eocene igneous bodies (Colmenares *et al.*, 2007; Tschanz *et al.*, 1969; 1974). The Buritaca pluton intrudes the Los Mangos Granulites and the Don Dieguito Anorthosite. The composition is biotite and hornblende tonalite, but also diorite, granodiorite and granite.

3. Materials and methods

3.1. SAMPLING

The studied samples were taken along the El Hierro creek. Many black, massive and heavy boul-

ders and cobbles were observed on the riverbed indicating the existence of an ore body upstream the creek. The only accessible outcrop is located approximately 1.2 km from the El Hierro river mouth to the Don Dieguito river. Two boulder samples were selected for this study, representative of the two mineralization styles observed: NEL-002-I for the banded nelsonite and NEL-002-J for the oxide-apatite norite.

3.2. BULK ROCK GEOCHEMISTRY

Major element compositions were determined at the Activation Laboratories Ltd. (Actlabs, Ontario, Canada) by X-ray fluorescence (XRF). Precision was better than ± 1.5 wt.% for a concentration of ≥ 10 wt.% and ± 2 wt.% for a concentration ≤ 10 wt.%. The concentration of trace elements, including rare earth elements (REE, La-Lu+Y) was determined by acid dissolution using ICP-MS (WRA4B2).

3.3. POWDER X-RAY DIFFRACTION (PXRD)

The samples were carefully ground using an agate mortar and pestle and were prepared manually pressing the powders in cylindrical standard containers of 16 mm diameter and 2.5 mm height, in order to obtain the flattest possible surface. The diffractograms were obtained using a PANalytical X'Pert PRO MPD Alpha1 powder diffractometer (Centres Científics i Tecnològics of the University of Barcelona – CCiT-UB, Barcelona, Spain), in Bragg-Brentano $\theta/2\theta$ geometry of 240 mm of radius, nickel filtered Cu K α radiation ($\lambda = 1.5418$ Å), 45 kV– and 40 mA. During analysis, samples were spun at 2 revolutions per second. Axial divergence Soller slits of 0.04 radians were used. Samples were scanned from 4° to 80° (2θ) with a step size of 0.017° and measuring time of 50 s per step, using a X'Celerator detector (active length = 2.122°), and a primary Ge monochromator. The detailed structural characterization and the semiquantitative mineral determination were performed by full profile analyses of the powder diffractograms using the software TOPAS version 4.0 (Rietveld method). This method was comple-

mented with the mineral proportions obtained by examining the polished sections and polished thin sections.

3.4. SCANNING ELECTRON MICROSCOPE (SEM)

Samples NEL-002-I and NEL-002-J were prepared as polished sections and polished thin sections for the study under the optical microscope using both reflected and transmitted light. They were carbon coated and examined under the ESEM Quanta 200 FEI scanning electron microscope, couples to an Energy Dispersive Spectrometer (SEM-EDS) Inca 250 (CCiT-UB, Barcelona, Spain). The microscope was set at 15 kV and 5 nA in backscattered electron (BSE) mode.

3.5. ELECTRON MICROPROBE ANALYSIS (EMPA)

Chemical analyses of the mineral phases of samples NEL-002-I and NEL-002-J were performed using a JEOL JXA-8230 electron microprobe (EMP) at the CCiT-UB, equipped with four wavelength dispersive spectrometers (WDS) and one energy dispersive spectrometer (EDS) operated at 20 kV accelerating voltage, 15 nA beam current and with a beam diameter of 5 μm , and 1 μm for minute mineral grains. The analytical standards used were: hematite (Fe), rutile (Ti), periclase (Mg), rhodonite (Mn), corundum (Al), Cr₂O₃ (Cr), metallic vanadium (V), diopside (Si), sphalerite (Zn), NiO (Ni), wollastonite (Ca) and metallic Nb (Nb). The correction procedure PAP (Pouchou and Pichoir, 1991) was used to convert specimen intensity ratios into concentrations. The Ti K β overlap on the V K α peak was corrected by the microprobe software. The structural formulae calculations of magnetite were performed on the basis of 4 oxygens, and of ilmenite and hematite, with 3 oxygens; and using the criteria of Carmichael (1966) for the Fe³⁺ estimation. The ilmenite (Ilm), hematite (Hem), geikielite (Gk) and pyrophanite (Pph) components were calculated as follows: $\text{Ilm} = (\text{Fe}^{2+} / (\text{Fe}^{2+} + \text{Fe}^{3+} + \text{Mg} + \text{Mn})) \cdot 100$, $\text{Hem} = (\text{Fe}^{3+} / (\text{Fe}^{2+} + \text{Fe}^{3+} + \text{Mg} + \text{Mn})) \cdot 100$, $\text{Gk} = (\text{Mg} / (\text{Fe}^{2+} + \text{Fe}^{3+} + \text{Mg} + \text{Mn})) \cdot 100$, $\text{Pph} = (\text{Mn} / (\text{Fe}^{2+} + \text{Fe}^{3+} + \text{Mg} + \text{Mn})) \cdot 100$; where Fe²⁺, Fe³⁺, Mg

and Mn are in cations per formula unit. The calculation of the ulvöspinel (Usp) component in magnetite was obtained following Stormer (1983). The composition of Fe-Ti(-V) oxides was plotted in triangular diagrams using the CSpace software (Torres-Roldán *et al.*, 2000). Mineral and end member abbreviations are after Whitney and Evans (2010).

4. Results

In general terms, two distinct mineralization styles were identified, each with characteristic mineralogy and textures. On one hand, the oxide-apatite norite is characterized by having a dark greenish color, and contains ilmenite, apatite, and minor silicates and other oxides (represented by sample NEL-002J). On the other hand, the banded nelsonite presents a banded fabric, the bands consisting of alternating massive black ilmenite grains, and apatite grains embedded in ilmenite (represented by sample NEL-002I).

4.1. MINERALOGY AND TEXTURES

From PXRD, the oxide-apatite norite mineralization consists mainly of amphibole (38 %), apatite (26 %), chlorite (12 %), ilmenite (9 %) and magnetite (4 %) (Figure 3a). In contrast, the banded nelsonite is characterized by the presence of predominantly ilmenite and hematite, magnetite and apatite, and minor chlorite was detected in the apatite separate (Figure 3b and 3c). Moreover, for the banded style, separated analyses were done to obtain the mineral expression of both the ore (Figure 3b) and the rest of the minerals (Figure 3c), yielding a simple composition of ilmenite and apatite.

4.1.1. OXIDE-APATITE NORITE ORE

In general terms, the oxide-apatite norite displays an isotropic fabric, it is holocrystalline and fine-grained (most of the grains are smaller than 1 mm). The presence of amphibole and chlorite

gives it a greenish color. In descending order of abundance, amphibole retrograded to chlorite is present in the sample in contact with apatite, ilmenite, sericitized plagioclase and pyrite (Figure 4). Granoblastic microstructures can be recognized between ilmenite and apatite grains displaying straight edges and triple points, therefore suggesting evidences of textural equilibrium (Figure 4a and 4b). Minor, scattered pyrite relicts are present as inclusions within some apatite grains, and are less than ~ 20 μm in size (Figure 4a, upper right). Altered, subhedral amphibole are found across the sample with light green tones and sizes ranging from a few hundreds of micrometers to some millimeters (Figure 4c). Stubby grains of apatite smaller than ~ 300 μm appear across the sample, both as inclusions within the amphibole, and as individual grains (Figure 4d). In contrast, minor plagioclase crystals range from sub- to euhedral and display polysynthetic twinning and alteration to sericite (Figure 4d).

Large, anhedral grains of magnetite make up no more than 5 vol % of the sample and tend to have serrated edges close to the ilmenite grains (Figure 4e), and no exsolved phases are observed within the magnetite grains. Ilmenite displays different titaniferous hematite exsolutions, including both a high lamellae density (Figure 4f) and the absence or fewer lamellae (Figure 4g). The very narrow needle-like titaniferous hematite lamellae are characteristic of this mineralization style, and are systematically aligned following the {0001} ilmenite planes. These lamellae display a mildly constant width ranging between 5 and 10 μm ; however, thicker lamellae, as the ones observed in Figure 4h (center right of the image), might form when narrow lamellae weld or are in contact with each other. In some cases, external grain exsolutions can be observed (Figure 4g; NW ilmenite border).

4.1.2. BANDED NELSONITE ORE

The banded nelsonite presents an equidimensional grain size (<1 mm) and an ilmenite/apatite ratio close to 50/50 (Figure 5a). The rock has a

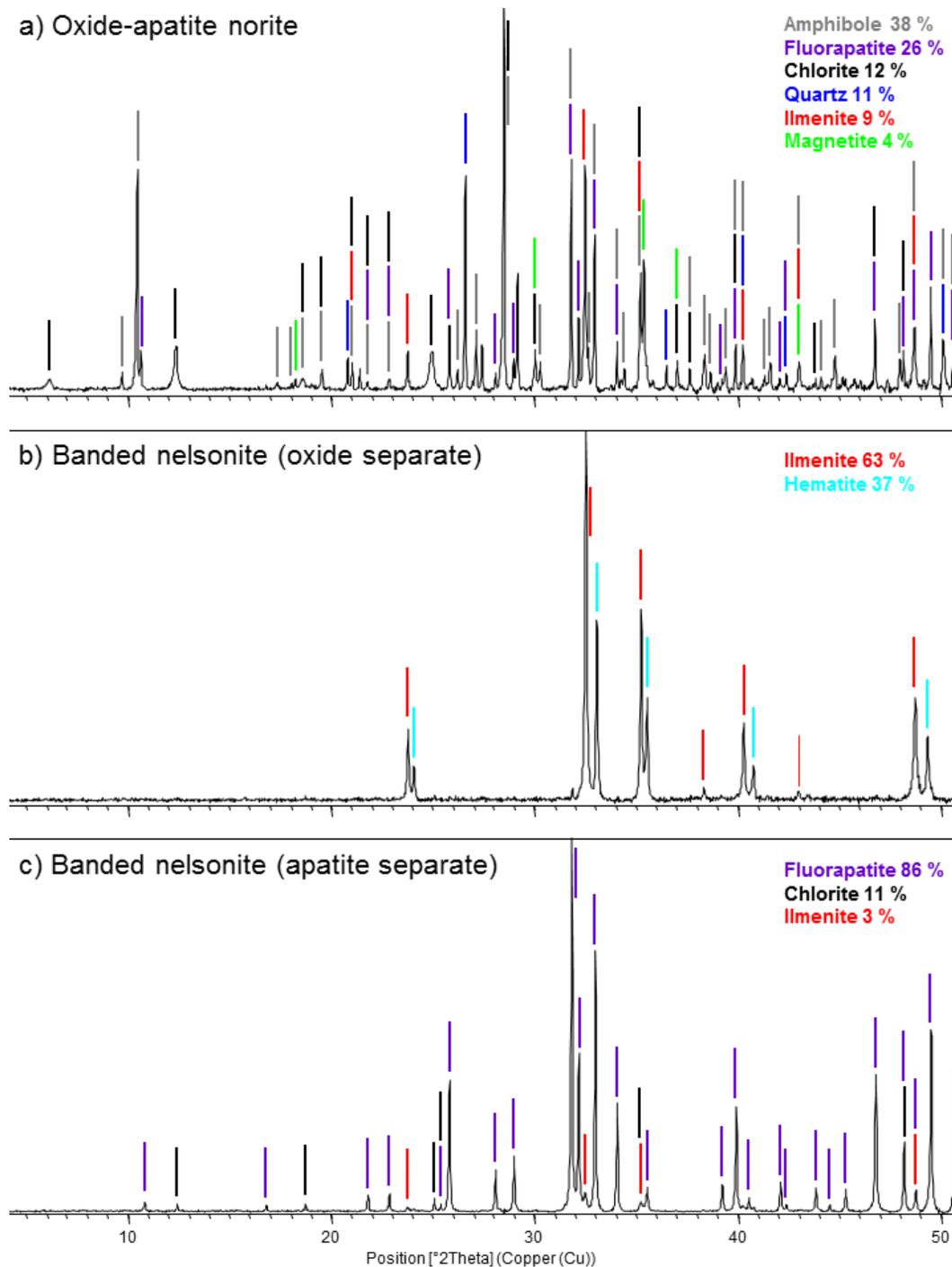


Figure 3 X-ray diffractograms (blue) with semiquantitative mineralogy of samples NEL-002-J (a: whole rock) and NEL-002- (b: Fe-Ti oxide; c: apatite grains), including the percentage of each mineral identified. Note the presence of amphibole and the abundance of chlorite in the oxide-apatite norite ore. The background is displayed in green lines.

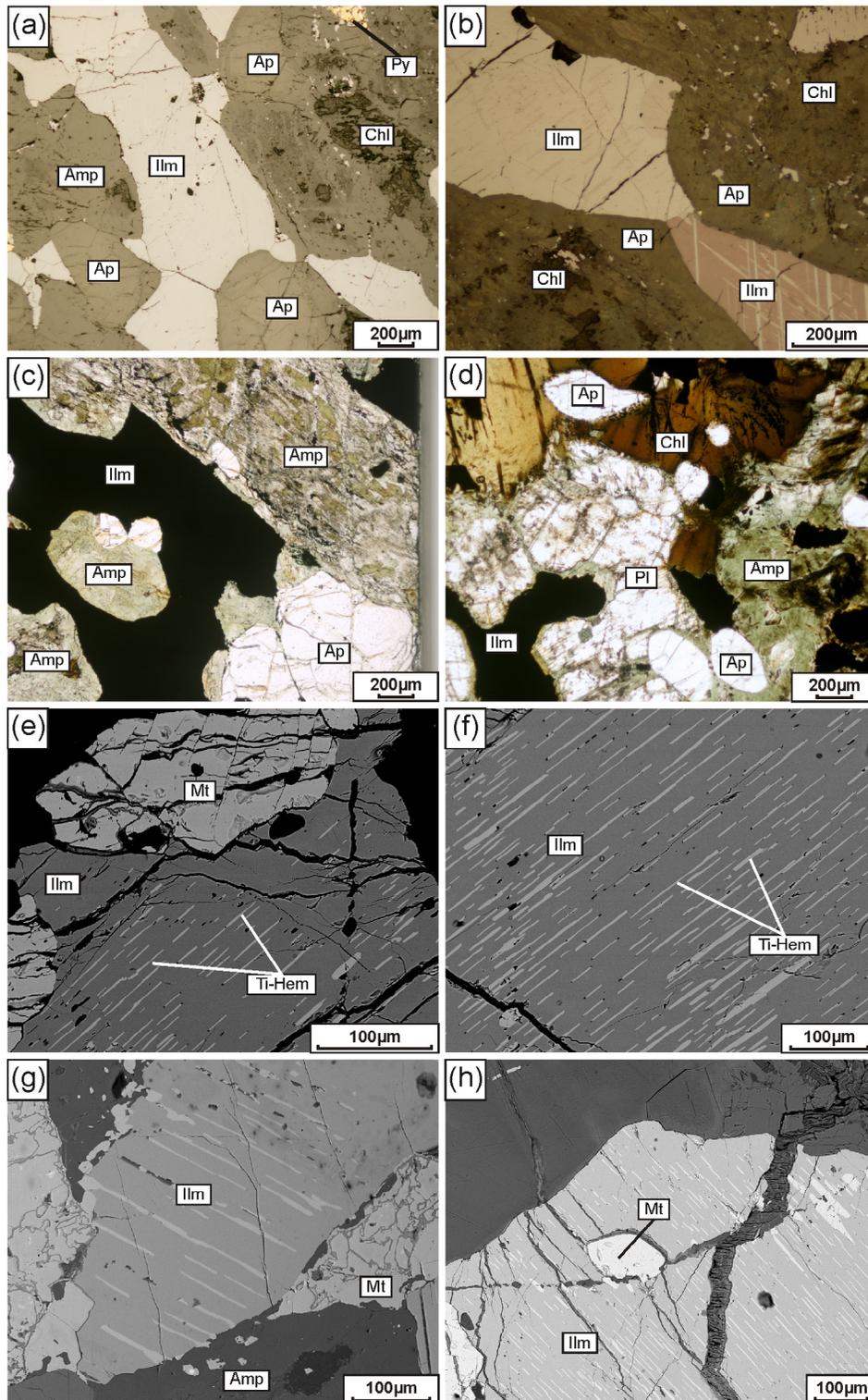


Figure 4 Optical (a, reflected, plane polarized light; b, reflected, crossed polars; c, d, transmitted, plane polarized light) and backscattered electron (BSE) micrographs (e,f,g,h,i: CCiT-UB) showing the most representative textures of the Fe-Ti(V) oxide-apatite norite ore from the Sierra Nevada de Santa Marta (SNSM), Colombia: (a) apatite grains with pyrite inclusions, in textural equilibrium with ilmenite in a granoblastic microstructure; (b) ilmenite grains surrounded by apatite and chlorite; (c) altered greenish euhedral amphibole with stubby grains of apatite; (d) sericitized plagioclase in contact with apatite, chlorite and amphibole; (e,f) anhedral grains of magnetite with serrated edges in contact with ilmenite grains with needle-like titaniferous hematite lamellae; (g) large ilmenite grains with hematite lamellae, surrounded by amphibole and magnetite (note the spinel rim in the SE ilmenite border with magnetite, and the external grain exsolutions in the NW ilmenite border with the amphibole); (h) thicker lamellae produced by agglomeration of contiguous lamellae (center, right. Mineral abbreviations are after Whitney and Evans (2010).

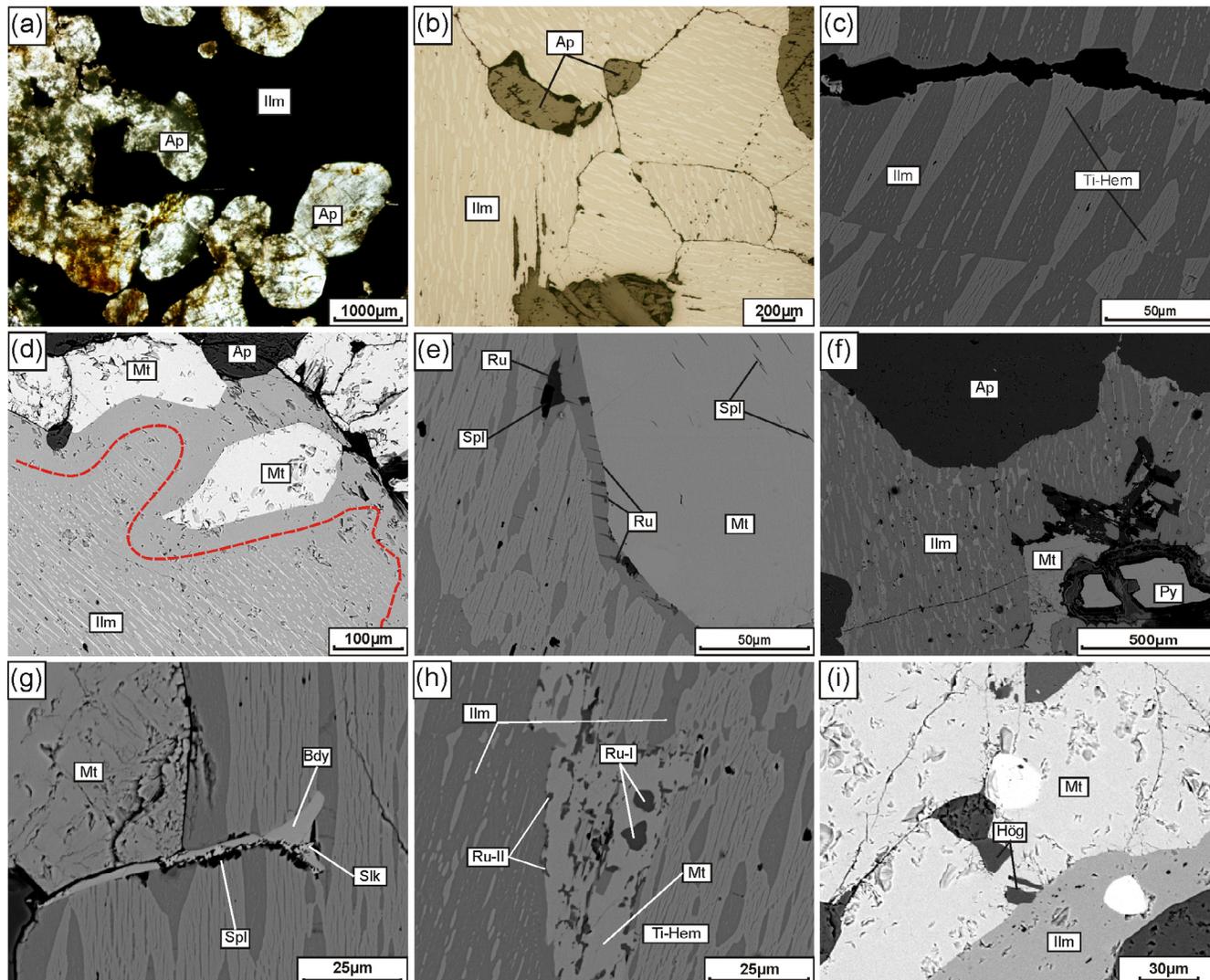


Figure 5 Optical (a, transmitted, plane-parallel light; b, reflected, plane parallel light) and backscattered electron (BSE) micrographs (c,d,e,f,g,h,i: CCiT-UB) showing the most representative textures of the banded nelsonite from the Sierra Nevada de Santa Marta (SNSM), Colombia: (a) nelsonite exhibiting rounded grains of apatite with a greenish alteration within ilmenite; (b) equilibrium textures between apatite and exsolved ilmenite grains; (c) patchy and minute exsolution lamellae exsolving from the ilmenite grains; (d) halo of a lamellae-free zone in the border of the ilmenite grain (outlined by the dashed line); (e) magnetite grains containing spindle-like grains of spinel and rutile needles in the contact with ilmenite; (f) presence of pyrite altered to iron oxides; (g) presence of srilankite (Slk) along with a baddeleyite (Bdy) lath in the border between magnetite and ilmenite; (h) symplectitic texture between magnetite and rutile; (i) possible presence of Zn-bearing spinel (dark grey) and spinel/hercynite (black) in ilmenite. Mineral abbreviations are after Whitney and Evans (2010).

banded fabric with an alternating arrangement of apatite-rich bands and ilmenite-rich bands. Large, well-developed grains ranging between 0.5 and 1 mm in diameter of both ilmenite and apatite are recognizable on polished surfaces of the hand specimen. The banded nelsonite ore presents a granoblastic microstructure, more evident than in the oxide-apatite norite, with triple points and sharp contacts between grains (Figure 5b). Ilmenite is characterized by the ubiquitous presence of two

generations of titaniferous hematite exsolutions, parallel to $\{0001\}$ ilmenite crystallization planes (Figure 5c). The first generation (hemo-ilmenite), coarser and patchy, or irregular-shaped, contains in turn later ilmenite lamellae. The second generation (ilmeno-hematite) consists of abundant minute lamellae scattered around the first generation exsolutions (Figure 5c). These exsolution features were first described by Hargraves (1959) and Carmichael (1961), and have also been reported by

McEnroe *et al.* (2002, 2007), Hebert *et al.* (2005), and Charlier *et al.* (2010), among others, in Canadian and Norwegian Fe-Ti oxide ores associated with anorthosite intrusives.

Some magnetite grains display nanometric spindle-like exsolutions of possibly spinel (too small to be analyzed) following the {100} directions of magnetite (Figure 5d, upper right). Lamellae-free halos may develop along the border of ilmenite when it is in contact with magnetite, as shown with the dashed line in Figure 5d. The decrease of hematite component towards the ilmenite rim was pointed out by Charlier *et al.* (2015) and references therein (Figure 5e). Hematite-rich areas within ilmenite grains may show ilmenite exsolutions, in some sort of inverse exsolution relationship. In this case, an exsolution-free border is locally found near the boundary with magnetite, where parallel rutile needles may grow on the magnetite-ilmenite boundary and inside the ilmenite grain (Figure 5e). Even though the banded nelsonite comprises primarily ilmenite and apatite, some isolated grains of pyrite are recognizable, generally altered to iron oxyhydroxides to the edges (Figure 5f).

In addition, a Y-shaped baddeleyite lath has been observed between two grains of magnetite and ilmenite. Around and inside this lath, small grains of possibly srilankite (Ti_2ZrO_6) are found (identified from SEM-EDS), along with minute spinel grains surrounding baddeleyite (Figure 5g).

Furthermore, a symplectitic texture is observed between rutile and magnetite grains (Figure 5h), with two different grain size groups and their location in relation with magnetite, possibly indicating different rutile formation stages: rutile I is embedded in the magnetite while rutile II is present along its margins. Finally, it must be noted that small amounts of a Zn-bearing minerals are present as small grains in contact with magnetite and ilmenite, possibly belonging to the högbomite group (Figure 5i), with 1.7-11.8 wt.% ZnO and 5.06-6.75 wt.% TiO_2 (7 analyses). Högbomite group minerals have been reported in other Fe-Ti oxide deposits (e.g., Devaraju *et al.*, 1981; Mohanty *et al.*, 2000).

4.2. BULK-ROCK GEOCHEMISTRY

The two mineralization styles identified differ from each other mainly in the TiO_2 content, which as expected, is higher in the banded nelsonite ore because of higher Fe-Ti oxide content (Table 1). The oxide-apatite norite ore contains 11 wt % TiO_2 and 33 wt % total Fe (as Fe_2O_3), whereas the banded nelsonite yields 31 wt % TiO_2 and 52 wt % Fe_2O_3 , because of the differences in ilmenite, magnetite and hematite contents. Si content is much higher in the oxide-apatite norite ore (28 wt % SiO_2) than in the banded nelsonite (0.6 wt % SiO_2), because of the presence of silicates. For the same reason, the values of CaO and MgO are higher in the oxide-apatite norite ore in comparison to the banded nelsonite, reaching 12 and 7 wt % respectively. Finally, both the oxide-apatite norite ore and the banded nelsonite contain similar P_2O_5 (6 wt % and 7 wt %, respectively). It is noteworthy that significant values of V and Cr are present in the banded mineralization style: 1857 ppm V and 280 ppm Cr.

Both the oxide-apatite norite and the banded nelsonite REE patterns depict a negative slope (Figure 6, Table 1). Additionally, both show a negative Eu anomaly. A comparison between the REE data in the studied nelsonites and other Fe-Ti(-V) oxide mineralizations around the world is displayed in Figure 6. Similarities in the spider diagrams of nelsonites found in Chiapas (Cisneros de León *et al.*, 2017), Hestnes (Duchesne and Liégeois, 2015) and Kunene (Villanova-de-Benavent *et al.*, 2017) indicate that the origin of the SNSM nelsonites could be similar. The Nelsonites in the Hestnes deposit are cumulates in an immiscible Fe-rich, Si-poor melt in oxidizing conditions (Duchesne and Liégeois, 2015).

4.3. MINERAL CHEMISTRY

4.3.1. ILMENITE

The chemical composition of ilmenite in the two mineralization types identified is very similar in terms of Ti and Fe (Figure 7, Table 2). In gen-

eral, ilmenite from the SNSM has a TiO_2 content ranging from 41.6 to 50.4 wt % (0.78-0.95 apfu Ti, 37 analyses) in the banded nelsonite, and from 49.2 to 50.1 wt % (0.93-0.94 apfu Ti, 3 analyses) in the oxide-apatite norite. No relation between the amount of Ti and the location of the measured point has been observed when measuring ilmenite points close or away from the lamellae. Note that, in the element maps of Figure 8, there is no significant compositional variation between the portion of ilmenite in contact with the lamellae and the parts away from them (Figure 8 a). V content is low, ranging from 0.3 to 0.6 wt % V_2O_3 in the banded nelsonite, and of 0.4 wt % V_2O_3 in the oxide-apatite norite.

However, the hematite component varies from 4.5 to 21.3 mol % (average of 10.4 mol %) in the banded nelsonite, and from 5.8 to 6.5 (average of 6.1 mol %) in the oxide-apatite norite.

In addition, ilmenite from the banded mineralization style contains between 0.2 and 6.6 wt % MnO which gives the mineral an important pyrophanite component (up to 14.9 mol % Pph, 5.2 mol % Pph on average), and up to 5.1 wt % MgO which translates into a remarkable geikielite component (up to 18.2 mol % Gk, 4.7 mol % Gk on average). In contrast, ilmenite from the oxide-ap-

atite norite shows much lower pyrophanite and geikielite components (3.2 and 0.2 mol %, on average, respectively), since very low values of MnO (1.5 wt % on average) and MgO (below 0.1 wt % on average) were obtained (Figure 9).

Finally, measured points within the titaniferous hematite lamellae exsolved from ilmenite give highly variable Ti (Table 3, Figure 7). TiO_2 in the banded nelsonite varies from 2 to 18.1 wt %, but possible contamination from the ilmenite host cannot be discarded; and V ranges from 0.17 to 0.88 wt % V_2O_3 .

Valid analyses of the patchy exsolutions in the banded nelsonite, and of the lamellae oxide-apatite norite mineralization style could not be obtained.

4.3.2. MAGNETITE

Magnetite grains in both mineralization styles do not differ significantly from each other in terms of chemical composition (Figure 7, Table 4). In the oxide-apatite norite, Ti content is too low to account for a titaniferous magnetite, with an average of 0.4 wt % TiO_2 (0.01 apfu Ti). In contrast, in the banded nelsonite it has an average of 0.5 wt % TiO_2 (0.01 apfu Ti), but it can reach up to

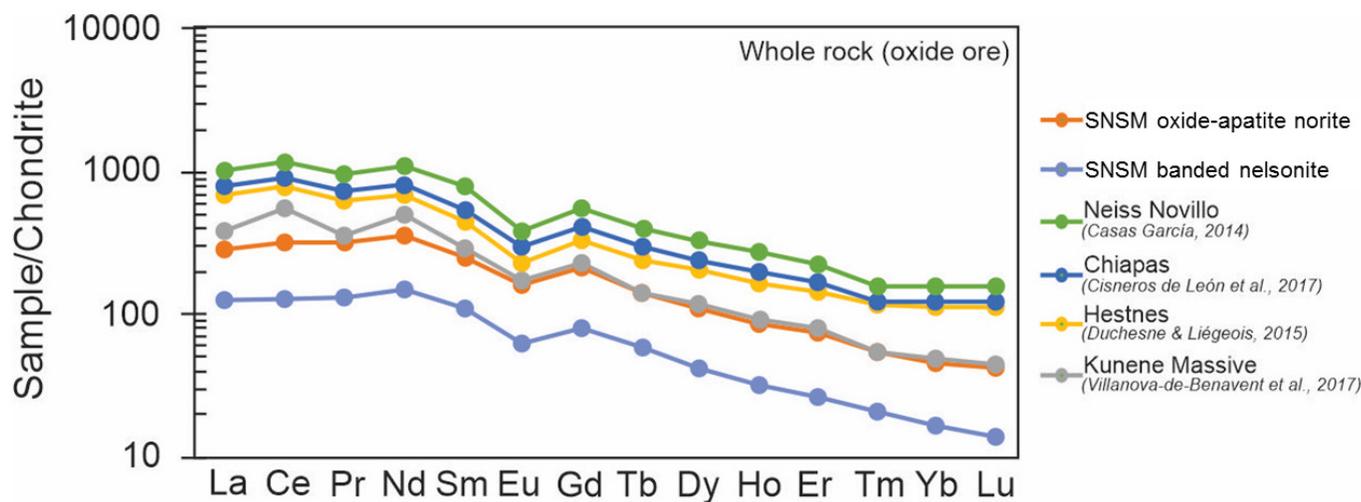


Figure 6 Spider diagram depicting the chondrite-normalized REE composition of the oxide-apatite norite and banded nelsonite oxide ores studied in the Sierra Nevada de Santa Marta, compared to values of other Fe-Ti(V) oxide ores worldwide. Chondrite values are after McDonough and Sun (1994).

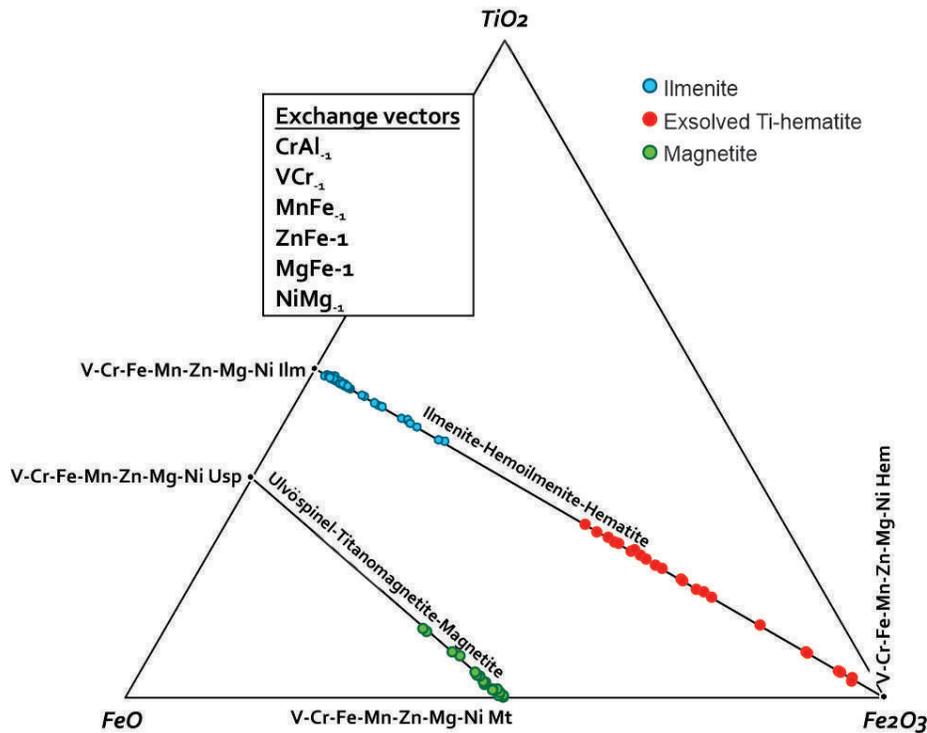


Figure 7 Ternary plot representing the composition of ilmenite, magnetite and hematite within the ilmenite-hematite and ulvöspinel-magnetite solid solutions in the SNSM. Ilm = ilmenite ($\text{Fe}^{2+}\text{TiO}_3$), Hem = hematite ($\text{Fe}^{3+}_2\text{O}_3$), Usp = ulvöspinel ($\text{Fe}^{2+}_2\text{TiO}_4$), Mt = magnetite ($\text{Fe}^{2+}\text{Fe}^{3+}_2\text{O}_4$), Rt = rutile (TiO_2). Diagram build using CSpace software from Torres-Roldán *et al.* (2000).

3.8 wt % TiO_2 . Also in the banded nelsonite, the amount of Cr is relatively high with an average of 0.25 and a maximum of 3.4 wt % Cr_2O_3 , while V content is 0.4 wt % on average, but reaching 0.78 wt % V_2O_5 . Magnetite in the oxide-apatite norite mineralization contains 0.02 wt % Cr_2O_3 and 0.5 wt % V_2O_5 on average. The presence of magnetite is outlined by the Fe predominance in the qualitative element maps in Figure 8b and 8c.

5. Discussion

The mineralogy and textures observed in the Fe-Ti(-V) oxides record the magmatic evolution, the cooling history and later modifications (*e.g.*, metamorphism) of the ores. Textures observed in both ilmenite and magnetite are the response of changes in the physical conditions and chemical composition during and after crystallization from a melt (Craig, 2001).

Firstly, according to Alfonso *et al.* (1990) and Tschanz *et al.* (1969), the mineralization in SNSM is associated with magmatic segregation processes that formed both banded nelsonite and oxide-apatite norite ores with little or no continuity. Many hypotheses explain the formation of Fe-Ti(-V) oxide deposits. One hypothesis involves fractional crystallization. Anorthosite parent magmas and their residual liquids are Ti-rich, which may enable the early saturation and crystallization of ilmenite.

However, the calculated cotectic proportion of ilmenite in a ferrobasic liquid is not enough to produce an ore deposit, so ilmenite enrichment processes are required (Charlier *et al.*, 2015; and references therein). Another hypothesis is magma immiscibility, which leads to the formation of cumulates, by segregation from a Fe-rich immiscible melt. Ilmenite and apatite could form before ferromagnesian silicates, while plagioclase, with lower density, could be separated by flotation. However, this hypothesis cannot be applied to

Table 1. Whole-rock chemical data for the Fe-Ti ore deposits in the SNSM.

Whole-rock chemical data for the Fe-Ti ore deposits in the SNSM		
Locality	Quebrada El Hierro	
Type	Banded nelsonite	Oxide-apatite norite
Coord. N	11°10'23.47"	
Coord. W	73°39'57.51"	
Sample	NEL002I	NEL002J
SiO ₂ (wt. %)	0.6	27.95
TiO ₂	30.27	11.04
Al ₂ O ₃	0.92	2.8
Fe ₂ O ₃	51.87	32.63
MgO	1.71	6.68
MnO	0.165	0.401
CaO	8.72	11.6
Na ₂ O	0.01	0.22
K ₂ O	0.01	0.07
P ₂ O ₅	6.76	6.25
LOI	-2.17	0.69
Sum	98.865	100.331
Sc (ppm)	42	44
Be	< 1	< 1
V	1857	569
Cr	280	< 20
Co	60	71
Ni	< 20	< 20
Cu	< 10	50
Zn	100	250
Ga	10	11
Ge	< 1	1
As	< 5	< 5
Rb	< 2	2
Sr	160	106
Y	48	81
Zr	79	257
Nb	23	14
Mo	< 2	< 2
Ag	< 0.5	0.7
In	< 0.2	< 0.2
Sn	25	4
Sb	< 0.5	< 0.5
Cs	< 0.5	< 0.5
Ba	12	6
Bi	< 0.4	< 0.4
La	29.8	40
Ce	80	114
Pr	12.6	18.3
Nd	67.4	98.6
Sm	16.3	24.1
Eu	3.72	5.17
Gd	16.3	24
Tb	2.1	3.1
Dy	10.6	16.6
Ho	1.8	3
Er	4.3	7.2
Tm	0.52	0.88
Yb	2.7	5.1
Lu	0.35	0.71
Hf	3.6	5.8
Ta	1.6	0.8
W	6	1
Tl	< 0.1	< 0.1
Pb	< 5	< 5
Th	0.4	0.1
U	0.2	0.1

the formation of deposits with only ilmenite, and enrichment processes are also necessary. Finally, variations of Cr in ilmenite along the ilmenite deposit of Lac Tio (Canada) suggest the possibility of various magma pulses, and alternating fractional crystallization and magma mixing stages (Charlier *et al.*, 2015; and references therein).

On one hand, the origin of nelsonites in the Grader layered intrusion (including nelsonites, massive oxides and oxide-apatite gabbro-norites), near the Tio Mine, has been interpreted as successive liquidus phases from a continuous fractional crystallization process, with ilmenite and plagioclase followed by apatite, orthopyroxene+magnetite, and finally clinopyroxene, with possible gravitational segregation of ilmenite and apatite from plagioclase (Charlier *et al.*, 2008). On the other hand, the oxide-apatite norite would develop dykes and sheets, formed by filter pressing of magmas successively depleted in Si, Al, alkali, and enriched in Fe, Ti and P. These would represent late stages of the evolution of these magmas, comagmatic with the anorthositic rocks, as suggested by McLelland *et al.* (1994). However, the limited availability of data regarding the spatial distribution and chemical variations in a larger scale of SNSM, and their relationships with the anorthosite host do not allow further constraining of the formation of the studied Fe-Ti(-V) oxide ores.

The liquid temperature at which the Fe-Ti(-V) oxide cumulates might have formed was experimentally established between 1100 to 1200°C in a jotunite parental magma for the associated anorthosites in the Rogaland ilmenite deposit (Duchesne and Liégeois, 2015). Nevertheless, according to Charlier *et al.* (2015), the emplacement of ilmenite deposits as a melt is highly questionable due to its high melting point, at least 1365°C, too high for a melt in the crust. Carmichael (1966) concluded that ilmenite, magnetite or both are the first minerals to crystallize. The SNSM oxide ores are characterized by dominant ilmenite coexisting with minor magnetite. According to McEnroe *et al.* (2004), the first oxide to crystallize from the magma is ilmenite, followed by magnetite, as a

Table 2. Representative EMP analyses of ilmenite from the SNSM.

Locality	Quebrada El Hierro											
	Banded nelsonite									Oxide-apatite norite		
	NEL002I									NEL002J		
No.												
Label	8-3	8-4	8-5	8-6	8-9	10-17	9-49	9-50	9-55	5-91	5-92	5-93
SiO ₂ (wt. %)	0.02	0.01	0.00	0.01	0.04	0.00	0.00	0.03	0.00	0.00	0.00	0.00
Al ₂ O ₃	0.01	0.00	0.02	0.05	0.03	0.02	0.04	0.00	0.03	0.00	0.00	0.02
Cr ₂ O ₃	0.08	0.00	0.00	0.05	0.01	0.06	0.02	0.07	0.02	0.00	0.00	0.02
TiO ₂	47.12	47.03	42.10	41.56	48.26	47.87	48.21	48.10	47.38	49.24	50.14	49.62
MgO	1.83	1.83	1.56	1.72	1.80	1.86	1.80	1.82	1.86	0.02	0.09	0.04
Na ₂ O	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.03	0.02	0.00	0.00
F	0.08	0.00	0.02	0.24	0.00	0.00	0.08	0.00	0.00	0.14	0.14	0.04
V ₂ O ₃	0.48	0.46	0.49	0.49	0.47	0.53	0.46	0.41	0.47	0.39	0.44	0.40
MnO	0.27	0.25	0.22	0.23	0.25	0.18	0.25	0.25	0.25	1.44	1.49	1.55
FeO *	38.86	38.72	34.80	34.06	39.94	39.50	39.86	39.72	38.82	42.63	43.40	42.89
Fe ₂ O ₃ *	11.82	12.24	21.34	22.42	9.49	10.65	10.18	10.13	11.81	6.76	6.29	6.10
NiO	0.00	0.02	0.01	0.02	0.00	0.04	0.02	0.02	0.00	0.00	0.00	0.00
ZnO	0.00	0.03	0.02	0.01	0.04	0.01	0.02	0.05	0.04	0.07	0.00	0.07
CaO	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.02	0.00
K ₂ O	0.00	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.01	0.00	0.00	0.01
Cl	0.00	0.00	0.01	0.00	0.00	0.01	0.00	0.01	0.00	0.01	0.00	0.00
Sum	100.57	100.61	100.59	100.87	100.33	100.73	100.94	100.61	100.74	100.71	102.01	100.76
Si (a.p.f.u.)	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Cr	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ti	0.88	0.88	0.79	0.78	0.90	0.89	0.90	0.90	0.89	0.93	0.94	0.94
Mg	0.07	0.07	0.06	0.06	0.07	0.07	0.07	0.07	0.07	0.00	0.00	0.00
Na	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
V	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01	0.01
Mn	0.01	0.01	0.00	0.00	0.01	0.00	0.01	0.01	0.01	0.03	0.03	0.03
Ni	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Zn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ca	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
K	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe ²⁺	0.81	0.81	0.73	0.71	0.83	0.82	0.83	0.83	0.81	0.90	0.90	0.90
Fe ³⁺	0.222	0.229	0.402	0.422	0.178	0.199	0.190	0.190	0.221	0.128	0.118	0.115
Ilm**	81.4	81.1	73.4	71.8	83.8	82.7	83.2	83.1	81.4	90.4	90.6	90.7
Hem**	11.1	11.5	20.2	21.3	9.0	10.0	9.6	9.5	11.1	6.4	5.9	5.8
Gk**	6.8	6.8	5.9	6.5	6.7	6.9	6.7	6.8	7.0	0.1	0.3	0.2
Pph**	0.6	0.5	0.5	0.5	0.5	0.4	0.5	0.5	0.5	3.1	3.2	3.3

result of a later reduction of the magma. However, this can also be explained by the observation that the presence of magnetite is controlled by $f O_2$, whereas the saturation of ilmenite is determined by TiO₂ (Toplis and Carroll, 1995). Hence, a melt in oxidizing conditions but a high TiO₂ can reach ilmenite saturation before magnetite (Charlier *et al.*, 2015; and references therein). Secondly, both ilmenite and magnetite in SNSM display subsolvus

features and symplectites, which reflect isochemical processes during cooling (*e.g.*, Craig, 2001; Sauerzapf *et al.*, 2008; Duchesne *et al.*, 2015). It is worth mentioning that size and spatial arrangement of the exsolution features are a function of the cooling history (McEnroe *et al.*, 2002). Once crystallized and upon cooling, ilmenite developed hematite exsolutions along the {0001} planes of the host (*e.g.*, Duchesne, 1970), and the process is

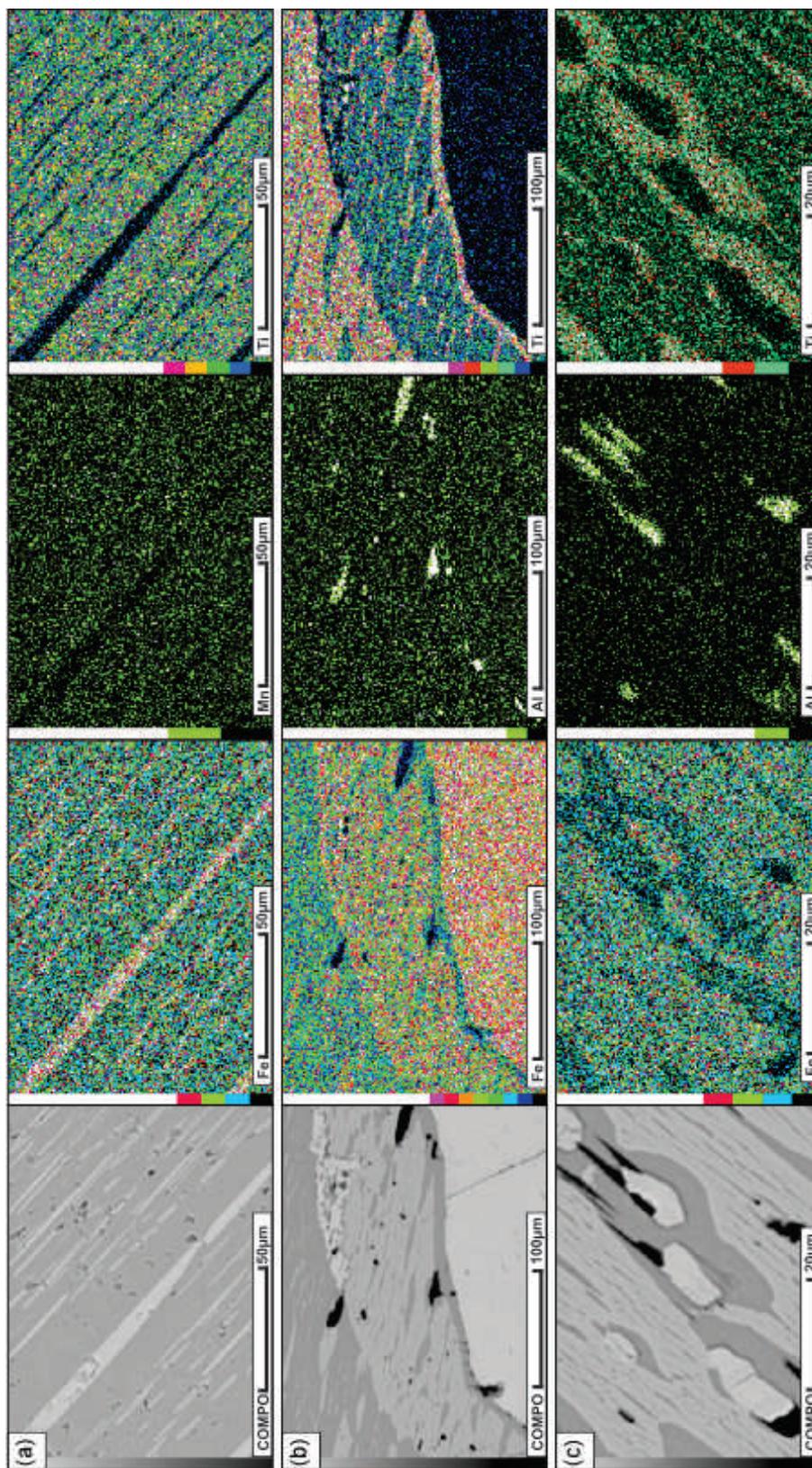


Figure 8 BSE images (left) and Fe, Al, Ti and Mn X-ray maps displaying hematite exsolution in ilmenite, and magnetite in the banded nelsonite. (a) ilmenite with hematite lamellae; (b) halo of inverse exsolution around a magnetite grain (light gray) where titaniferous hematite contains ilmenite lamellae. Note the intermediate colors in the Fe and Ti maps for the titaniferous hematite exsolved area around magnetite. (c) Micrometer sized, rounded magnetite grains (light gray) within ilmenite-hematite, exsolution.

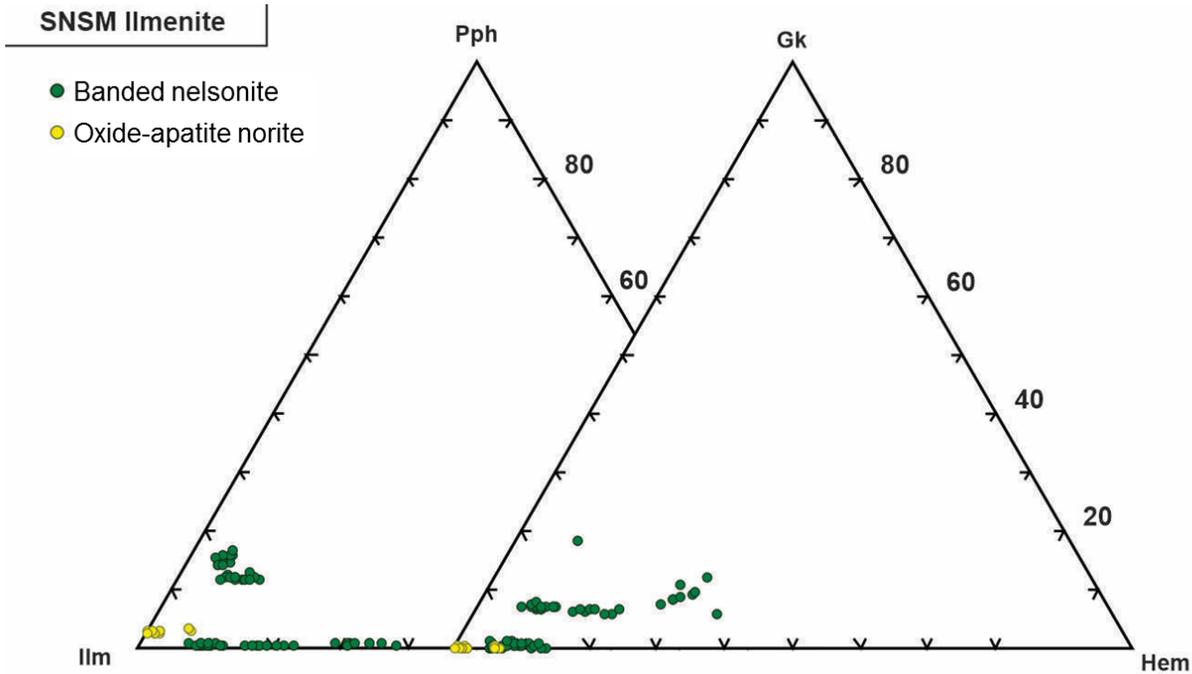


Figure 9 Triangular plots showing the ilmenite (Ilm)-pyrophanite (Pph)-hematite (Hem) (left) and ilmenite (Ilm)-geikielite (Gk)-hematite (Hem) (right) components of ilmenite in the SNSM.

isochemical (Arguin *et al.*, 2018; and references therein). McEnroe *et al.* (2002) pointed out that the equilibrium recrystallization occurred prior to the low temperature exsolution process, since every ilmenite grain has its own set of {0001} planes in different directions, and is interrupted by grain boundaries, as the textures observed in the SNSM banded nelsonite (Figures 5a and 5b). Carmichael (1961) established that the temperature of the hemo-ilmenite Precambrian deposits of Allard Lake (Québec) was between 575-600°C when the hematite lamellae exsolved, and that these conditions were maintained during a long period of time. Alternatively, in studies combining mineral chemistry and magnetism, thermodynamic models indicate that there is a miscibility gap below 700-800°C between a paramagnetic hematite-rich phase and a paramagnetic ilmenite-rich phase (McEnroe *et al.*, 2007). As temperature decreases, canted ferromagnetism in hematite appears at lower temperatures (McEnroe *et al.*, 2007).

Balsley and Buddington (1958) already stated that the number and size of hematite lamellae are related to the hematite component in ilmenite: exsolutions are not visible for Hem below 7 mol %;

exsolutions appear when Hem exceeds 7–9 mol %; and when Hem is above 18 mol%, two generations may appear. This is coherent with the textures and mineral chemistry obtained in the banded nelsonite of the SNSM (4.4-22.4 wt% Fe₂O₃, 4.5-21.3 mol % Hem in the banded nelsonite, with two generations of exsolutions; and 6.1-6.8 wt% Fe₂O₃, 5.8-6.5 mol % Hem in the oxide-apatite norite, with smaller and fewer exsolutions). According to X-ray diffraction and Curie point studies in Allard Lake by Carmichael (1961), the small hematite lamellae exsolved at a lower temperature and later than the larger lamellae. In particular, Carmichael (1961) suggested that the smaller lamellae might have formed during erosion that enabled the Allard Lake deposits to outcrop. McEnroe *et al.* (2002) were able to analyze both sets of lamellae by EMP and transmission electron microscopy, and determined that the original ilmenite, Ti-poor (Ilm84) would exsolve at high temperature to form Ilm94 host and micrometer-sized Ilm23 lamellae. Further cooling would produce nanometer-sized Ilm15 surrounded by Ilm98 hosts. The fact that the smaller lamellae would be richer in hematite, if exsolution was so quick that the larger lamellae

Table 3. Representative EMP analyses of Ti-hematite lamellae from the SNSM.

Representative EMP analyses of Ti-hematite lamellae from the SNSM							
Locality	Quebrada El Hierro						
Type	Banded nelsonite						
No.	NEL002I						
Label	1-38	10-15	10-20	10-21	9-44	9-47	9-48
SiO ₂ (wt. %)	0.00	0.06	0.06	0.02	0.00	0.00	0.04
Al ₂ O ₃	0.13	0.05	0.55	0.09	0.03	0.04	0.08
Cr ₂ O ₃	0.01	0.23	0.11	0.21	0.12	0.19	0.17
TiO ₂	9.69	13.70	15.58	15.16	12.76	15.59	17.24
MgO	0.00	0.36	0.54	0.35	0.22	0.35	0.51
Na ₂ O	0.03	0.00	0.00	0.00	0.00	0.01	0.01
F	0.07	0.00	0.00	0.00	0.03	0.00	0.00
V ₂ O ₃	0.25	0.97	0.81	0.88	0.80	0.80	0.80
MnO	1.09	0.04	0.04	0.06	0.06	0.06	0.05
FeO *	38.14	42.65	44.28	44.14	41.91	44.33	45.72
Fe ₂ O ₃ *	48.55	40.02	36.47	37.65	42.24	36.87	33.65
NiO	0.01	0.01	0.08	0.01	0.00	0.00	0.00
ZnO	0.03	0.05	0.06	0.00	0.05	0.01	0.04
CaO	0.01	0.01	0.00	0.00	0.00	0.00	0.01
K ₂ O	0.01	0.00	0.00	0.00	0.00	0.00	0.00
Cl	0.00	0.00	0.00	0.00	0.00	0.00	0.01
Sum	98.01	98.15	98.59	98.57	98.22	98.27	98.34
Si (a.p.f.u.)	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	0.00	0.00	0.02	0.00	0.00	0.00	0.00
Cr	0.00	0.01	0.00	0.00	0.00	0.00	0.00
Ti	0.21	0.30	0.34	0.33	0.28	0.34	0.37
Mg	0.00	0.02	0.02	0.02	0.01	0.02	0.02
Na	0.00	0.00	0.00	0.00	0.00	0.00	0.00
V	0.01	0.02	0.02	0.02	0.02	0.02	0.02
Mn	0.03	0.00	0.00	0.00	0.00	0.00	0.00
Ni	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Zn	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ca	0.00	0.00	0.00	0.00	0.00	0.00	0.00
K	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe ²⁺	0.93	1.03	1.06	1.06	1.02	1.07	1.10
Fe ³⁺	1.07	0.87	0.79	0.81	0.92	0.80	0.73

a.p.f.u. = atoms per formula unit

* Calculated

would not have reached equilibrium, was also predicted by Carmichael (1961).

Extensive diffusion of Fe²⁺, Fe³⁺ and Ti+Fe³⁺ from the Ilm-Hemss not only yields the formation of the exsolved lamellae but also generates relatively large areas of exsolution-free haloes (Arguin *et al.*,

2018), as the ones seen in Figure 5d. In any case, it should be noted that most thermodynamic and experimental approaches to the Fe-Ti oxide system do not take into account the effect of components other than Ti, Fe²⁺ and Fe³⁺ (*e.g.*, Mg and Mn) and only apply to certain pressure ranges (Robinson *et*

Table 4. Representative EMP analyses of magnetite from the SNMS.

Locality	Quebrada El Hierro										
	Banded nelsonite								Oxide-apatite norite		
Type	NEL002I								NEL002J		
Sample											
Label	1-1	1-5	1-8	1-45	1-46	10-38	10-40	9-42	5-88	5-89	5-90
SiO ₂ (wt. %)	0.03	0.01	0.05	0.01	0.04	0.01	0.03	0.00	0.03	0.03	0.00
Al ₂ O ₃	0.26	0.21	0.19	0.28	0.33	0.10	0.95	0.07	0.45	0.28	0.18
Cr ₂ O ₃	0.04	0.03	0.07	0.01	0.00	0.37	0.31	3.39	0.00	0.04	0.02
TiO ₂	0.06	0.06	0.15	0.13	0.11	1.58	3.76	0.09	0.63	0.14	0.29
MgO	0.00	0.02	0.04	0.03	0.00	0.00	0.35	0.02	0.00	0.05	0.00
Na ₂ O	0.00	0.05	0.00	0.00	0.06	0.04	0.00	0.00	0.00	0.00	0.06
F	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
V ₂ O ₃	0.23	0.22	0.20	0.21	0.22	0.63	0.63	0.78	0.50	0.55	0.45
MnO	0.04	0.09	0.08	0.06	0.05	0.00	0.03	0.01	0.04	0.00	0.02
FeO *	30.80	30.47	30.75	31.25	31.08	32.42	34.17	31.10	32.04	31.56	31.52
Fe ₂ O ₃ *	67.61	67.89	67.45	68.79	68.78	65.40	60.30	64.59	67.56	68.86	69.18
NiO	0.0341	0.01	0.00	0.00	0.04	0.12	0.12	0.12	0.03	0.04	0.01
ZnO	0.0115	0.00	0.07	0.07	0.00	0.04	0.14	0.00	0.05	0.09	0.06
CaO	0.0216	0.00	0.02	0.02	0.05	0.00	0.00	0.00	0.00	0.01	0.02
K ₂ O	0.0000	0.00	0.00	0.01	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Cl	0.0000	0.00	0.01	0.00	0.01	0.01	0.00	0.00	0.00	0.00	0.01
Sum	99.13	99.07	99.07	100.88	100.78	100.73	100.80	100.17	101.34	101.66	101.82
Si (a.p.f.u.)	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Al	0.01	0.01	0.01	0.01	0.01	0.00	0.04	0.00	0.02	0.01	0.01
Cr	0.00	0.00	0.00	0.00	0.00	0.01	0.01	0.10	0.00	0.00	0.00
Ti	0.00	0.00	0.00	0.00	0.00	0.05	0.11	0.00	0.02	0.00	0.01
Mg	0.00	0.00	0.00	0.00	0.00	0.00	0.02	0.00	0.00	0.00	0.00
Na	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
V	0.01	0.01	0.01	0.01	0.01	0.02	0.02	0.02	0.02	0.02	0.01
Mn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ni	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Zn	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Ca	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
K	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00	0.00
Fe ²⁺	1.00	0.99	1.00	1.00	0.99	1.03	1.08	1.00	1.02	1.00	1.00
Fe ³⁺	1.97	1.98	1.97	1.97	1.97	1.88	1.71	1.86	1.93	1.96	1.97
Usp ^{**}	0.58	0.52	1.30	1.14	1.00	12.94	27.99	0.95	5.43	1.19	2.51

a.p.f.u. = atoms per formula unit

* Calculated

**Percent, calculated following Stormer (1983)

al., 2004). In parallel, a similar process takes place within the magnetite grains where spinel exsolves in the form of minute needles along the {100} planes of magnetite, like the ones observed in Figure 5e. Turnock and Eugster (1962) found that the magnetite-hercynite solvus lies at 860 ± 15 °C. In any case, there are very few aluminous spinel exsolutions (observed only in the banded nelsonite, Figure 5d), indicating that magnetite originally contained a very

small amount of Al₂O₃. These aluminous spinel exsolutions may also occur in ilmenite, as seen in Figure 5i and in Charlier *et al.* (2006), but are less frequent. The fact that magnetite in the SNSM does not display ulvöspinel-magnetite exsolutions, and that it contains very low Ti, implies that Ti was mostly spent on the crystallization of ilmenite.

Apart from that, symplectitic textures involving rutile and magnetite, similar to those observed in

Figure 5h, were also described by Tan *et al.* (2015) in the Xinjie mafic-ultramafic Fe-Ti oxide deposit (China). These textures could be explained by a local low temperature alteration of hemo-ilmenite to magnetite + rutile under oxidizing conditions. The subsolidus oxidation of the original Ilm-Hemss would lead to a more Fe-rich Ilm-Hemss, in turn producing a heterogeneous nucleation of rutile. The rutile growth would consume Ti^{4+} , generating an excess of Fe^{2+} along the Ilm-Hemss boundary. This process enhances the diffusion of the Fe^{3+} from the Ilm-Hemss towards the boundary, and together with Fe^{2+} , leads to the nucleation of magnetite between ilmenite and rutile. This process ends below $\sim 650^{\circ}C$, and hematite exsolutions would start forming after.

Finally, there are mineralogical evidences that the oxide-apatite norite and the nelsonite underwent metamorphism to some extent, as exemplified for example, by the presence of amphibole and chlorite instead of pyroxene in the oxide-apatite norite (Figure 4d). Likewise, högbomite group minerals have been reported in high grade metamorphosed peraluminous or mafic rock, together with other Al (*e.g.*, spinel) and Ti (*e.g.*, ilmenite) minerals (Petersen *et al.*, 1989). Furthermore, the assemblage of baddeleyite, zircon and srilankite (Ti_2ZrO_6), associated with ilmenite, has also been reported in a Proterozoic anorthosite-mangerite-charnokite-granite suite which underwent granulite facies metamorphism (Bingen *et al.*, 2001). According to these authors, ilmenite (with tens of ppm Zr) can exsolve baddeleyite during subsolidus cooling of the intrusion; and srilankite could be produced after a reaction between ilmenite and baddeleyite. Ultimately, during granulite facies metamorphism, zircon could form after the reaction of baddeleyite with silica (Bingen *et al.*, 2001). However, in SNSM baddeleyite, with inclusions of srilankite and a spinel rim, is included in ilmenite, whereas in Bingen *et al.* (2001) baddeleyite and srilankite are mostly in the contacts with, and rarely included in, ilmenite (also coexisting with spinels, of gahnitic composition). Since the zircon coronas described in Bingen

et al. (2001) and Charlier *et al.* (2015) do not have any systematic spatial relationship with hematite lamellae in ilmenite, these authors determined that Zr exsolution happened before hematite exsolution. In contrast, baddeleyite in SNSM seems to crosscut hematite, although more data are needed to obtain a solid conclusion.

6. Conclusions

The Fe-Ti(-V) oxide ores in the Proterozoic massif type anorthosite in SNSM display slightly differentiated REE patterns and negative Eu anomalies. According to the mineralogy, textures and mineral chemistry, two ore types were distinguished: i) oxide-apatite norite (fine grained ilmenite and magnetite disseminated in an assemblage of apatite, amphibole, chlorite, rutile and plagioclase); and ii) massive, banded nelsonite, with alternating layers of coarse grained ilmenite and magnetite, and apatite, with minor baddeleyite, srilankite and högbomite. In both, ilmenite displays different hematite exsolution features: needle-like in the oxide-apatite norite ore, and two generations of exsolutions in the banded ore (an earlier, coarser grained; and a later, finer grained). The composition of magnetite in terms of Ti and V is quite homogeneous in both ore types, whereas ilmenite has higher Fe^{3+} , Mn and Mg contents in the banded ore.

The mineral assemblages and textures observed are similar to those of other major ilmenite deposits in the world, such as Lake Allard in Canada, and Tellnes in Norway. However, more information is needed concerning the spatial location of the different ore types in order to extend the knowledge on the formation of these ores.

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