

Vertebrate fossils from the San José de Gracia quarry, a new Late Cretaceous marine fossil site in Puebla, Mexico

Fósiles de vertebrados de la cantera San José de Gracia, un nuevo sitio fósil marino del Cretácico Superior en Puebla, México

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ABSTRACT

The San José de Gracia Quarry, located within the Municipality of Molcaxac, southern Puebla, Mexico, is a new paleontological site discovered about a decade ago. This paper is the first formal scientific contribution of this site. The quarry occupies an area no larger than the two hectares that is exploited for commercial purposes, where slabs are extracted from a marine sequence of cream-brown strata of poorly carbonated clays, with centimetric thickness. Although the upper and lower limits of this sequence are not known; the characteristic lithology and fossil content do not correspond with those of the regional geological units previously reported. Although the fossil association recovered from the San José de Gracia Quarry is composed mostly of fish remains; it also contains ammonites, belemnites, inoceramids, indeterminate ostreids, reptiles and few plant remains. In this work, the first fossil fish remains from this locality are described, including representatives of *Enchodus* and clupeid both previously known from Mexico; as well as the first record of *Dercetis* and *Xenyllion* from this country. In addition, a yaguarasaurine mosasaur is also reported. Based on its fossil content and its lithology, the fossil-bearing strata of the San José de Gracia quarry can be possibly referred to the Mexcala Formation (Morelos-Guerrero Basin) of Turonian age.

Keywords: San José de Gracia quarry, Turonian, Puebla, Mexcala Formation, vertebrates.

RESUMEN

La Cantera San José de Gracia, ubicada dentro del Municipio de Molcaxac, al sur del Estado de Puebla, México, es un sitio paleontológico descubierto hace cerca de una década. Este trabajo representa la primera contribución científica formal. Esta cantera ocupa una superficie no mayor a las dos hectáreas explotadas con fines comerciales, donde se extraen lajas de una secuencia de origen marino, compuesta de margas arcillosas pobremente carbonatadas depositadas en estratos paralelos, de color marrón-crema, cuyo grosor alcanza varios centímetros. A pesar de que los límites, superior e inferior, de esta unidad rocosa no son conocidos; sus características litológicas y contenido fósil no corresponden con aquellos de las unidades geológicas previamente reportadas para esta región. Aunque la asociación fósil recuperada de la Cantera San José de Gracia se compone en su mayor parte de restos de peces; ésta también contiene amonites, belemnites, inoceramidos, ostreidos indeterminados, reptiles y muy escasos restos de plantas. En este trabajo se describen los primeros restos de peces fósiles de esta localidad, entre los que se destacan *Enchodus*, *Laminospondylus* y un clupeido, grupos ya previamente conocidos en México; así como el primer registro mexicano de *Dercetis* y *Xenyllion*. Aquí, también se reporta el hallazgo de un mosasaurio yaguarasaurino en México. Con base en su contenido fósil y su litología, se sugiere que los estratos fosilíferos de la Cantera San José de Gracia son referibles posiblemente a la Formación Mexcala (Cuenca Morelos-Guerrero) y son de edad Turoniana.

Palabras clave: Cantera San José de Gracia, edad Turoniana, Formación Mexcala, Puebla, vertebrados.

1. Introduction

In 2006, people of the San José de Gracia town, Municipality of Molcaxac, State of Puebla, central Mexico, began to exploit a deposit of sedimentary rocks within the limits of their communal property (Figure 1). In this site, herein formally named as “San José de Gracia quarry”, cream-brown marly flagstones with reddish veins are extracted. Given its good appearance and resistance, these rocks are used to cover floors and walls. The artisanal procedures used during their extraction have allowed us

to discover a peculiar fossiliferous association that we partially report in this contribution (see also Alvarado-Ortega *et al.*, 2009, Pulido-González, 2011; Romero-García, 2013; Zavaleta-Villareal, 2015).

Fossils from the San José de Gracia quarry are scarce and poorly preserved if they are compared with those from the Tlayúa quarry, a Mexican emblematic Lagerstätte paleontological site located about 10 km south from Molcaxac town (Applegate *et al.*, 2006). For this reason, the workers and traders of flagstones extracted in this new quarry have scorned the scientific and aesthetic

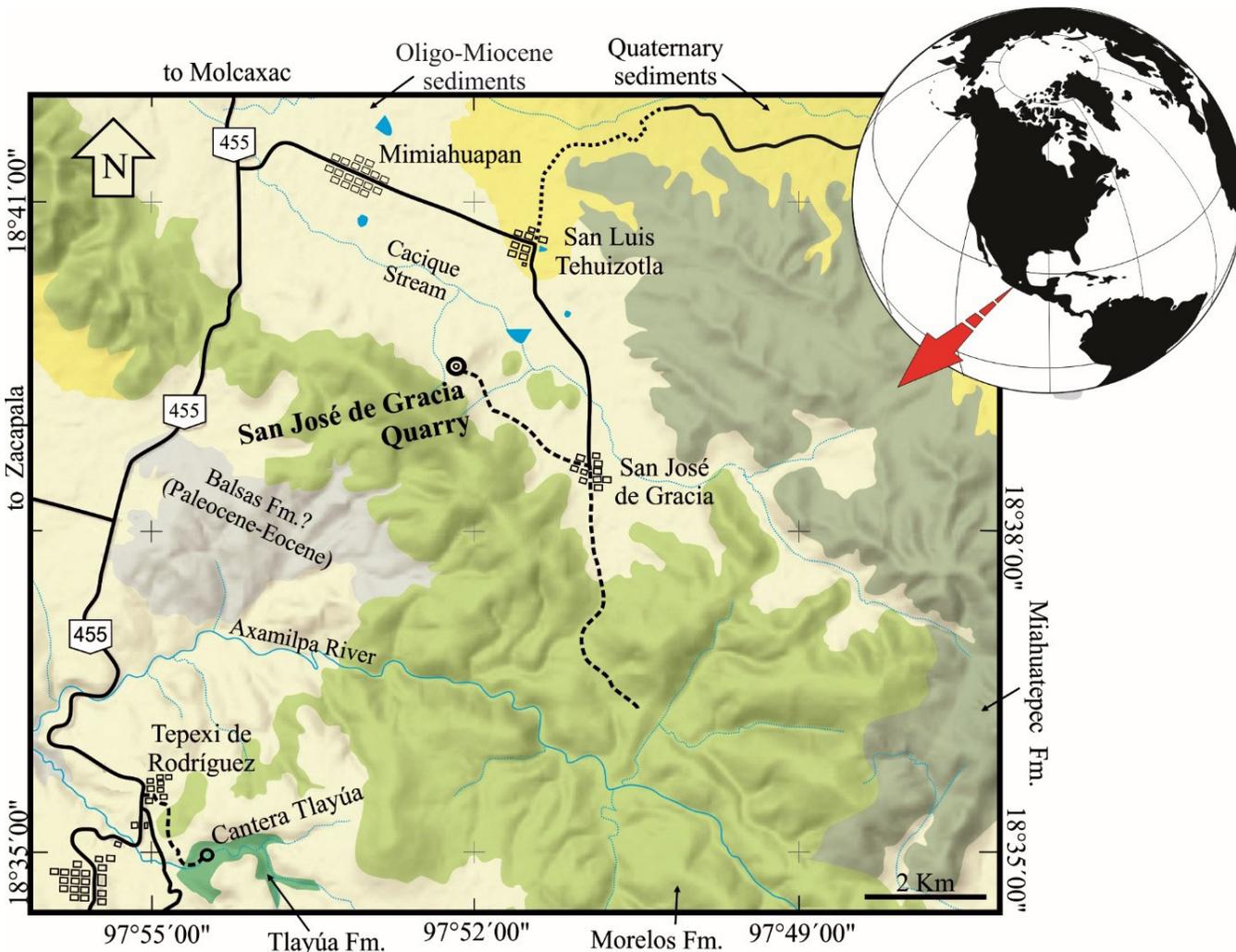


Figure 1 Map of San José de Gracia quarry showing the geological formations present in this region of Puebla, Mexico (based on map of Servicio Geológico Mexicano, 2006).

value of their fossils; this has caused the loss or destruction of a great quantity of San José de Gracia specimens during the extraction and sale of these rocks.

In 2007, Messrs. Félix and Sebastián Arangu thy, traders of flagstones from different sources, inhabitants of Tepexi Rodríguez town (neighboring town of Molcaxac), and collaborators of the Tlayúa Paleontological Project, supported by the Instituto de Geología, Universidad Nacional Autónoma de México (UNAM), gave some fossils from the San José de Gracia to the first author of this work. Thus, in 2008, researchers of the UNAM and the Benemérita Universidad Autónoma de Puebla (BUAP), as well as workers of the San José de Gracia have worked together in order to recover fossils from this new locality, with the final goal of preserving and studying them formally. Currently, about 200 macrofossils from this quarry have been deposited in the scientific collections of UNAM and BUAP, including invertebrates (ammonites, belemnites, gastropods, bivalves, and crabs), vertebrates (fishes and reptiles), and scarce plant remains.

The San José de Gracia quarry is located 2.5 km west from the center of homonymous town, within the Municipality of Molcaxac, Puebla, central Mexico. This site is in the southeastern side of the Cacique Stream. Currently, the center of the quarry is at the altitude of 1930 m.a.s.l. and between the coordinates 18° 39' 27.25" N and 97° 52' 13.94" W. The exploited surface in San José de Gracia quarry extends about 150 m from East to West and 40 m from North to South (Figure 1).

Up to now, a sequence of about 4 m thick of decimetric marly strata has been exploited in the San José de Gracia quarry (Figure 2); which includes marl layers of clays poorly carbonated, showing normal gradation, and deposited in finely laminated parallel planes (Figure 2). Although the soil covers the geological features around the San José de Gracia Quarry making it impossible to recognize the presence of possible faults and contacts with other lithological units; the strata exposed in this quarry have a strike and dip that suggest the presence of an anticline. At the eastern end of the quarry the strike is variable between N50-88°W, while in the west this geomorphological feature is



Figure 2 The San José de Gracia quarry; a) general view of one of the exploited sections in this quarry; b) Details of the San José de Gracia quarry showing the centimetric marl strata.

close to N90°W. At the same time, the dip in these layers is variable between 23-35° S and 23-40° S respectively (Figure 2). As it is noted in the results of this work, the hitherto known fossil assemblage of San José de Gracia quarry includes Late Cretaceous marine vertebrates and invertebrates.

Since our first glance at the fossils from the San José de Gracia quarry, we recognized that they represent a peculiar assemblage preserved into a distinctive lithology, never observed in this region of Puebla. Therefore, we are convinced that this fossil assemblage deserves to be recognized in greater detail. Thus, the aim of this contribution is to report the San José de Gracia fossils and describe the geological details of this paleontological site.

2. Materials and methods

Preparation methods: Fossils were mechanically prepared with needles and odontological scavengers under stereoscopic microscope. When necessary, the fossils were glued with cyanoacrylate; they were hardened with weak solution of plexigym diluted in methacrylate acetate applied with a fine brush.

Institutional abbreviations: Fossils from the San José de Gracia are deposited and catalogued under the acronym IGM that corresponds to the Colección Nacional de Paleontología, housed into the Instituto de Geología, UNAM, as well as BUA-PALZ that is the acronym of the Colección de Paleontología of BUAP. The San José de Gracia quarry is recorded in the catalog of paleontological sites of IGM as IGM-loc 3876.

Anatomical nomenclature and abbreviations: The taxonomical identification of fossil taxa described here, as well as the anatomical nomenclature and abbreviations used here (Figure 3 to 8), are based on the publications respectively cited.

3. Results

3.1. FOSSIL VERTEBRATE ASSEMBLAGE OF SAN JOSÉ DE GRACIA

3.1.1. SYSTEMATIC PALEONTOLOGY

Order Clupeiformes sensu Grande, 1985
Suborder Clupeoidei sensu Grande, 1985

Family Clupeidae Cuvier, 1817
Genus and species indeterminate

Referred material: IGM 11190 A and B, part and counterpart, fairly complete specimen without the jaws and fins (Figure 3).

Description: This almost complete specimen preserved in part and counterpart shows the trunk strongly curved, reaches 40 mm of standard length (SL), with none of its fins is preserved. In life this should be a chubby fish, with a triangular head and a shallow caudal peduncle, in which the maximum height of the trunk was about 13 mm.

Although the skull is poorly preserved, it is possible to observe that the skull roof bones are smooth. In addition, the parietals have the supra-temporal commissure and are laterally displaced, separated by the supraoccipital bone. In the ethmoid region, both the mesethmoid and vomer are well preserved, the first is located just above of the latter bone. The bones of both jaws are not preserved in this specimen; hence, the hyoid bones are partially exposed showing that the anterior ceratohyal is an elongated, rectangular, and not pierced bone.

There are thirty-four centra in the vertebral column, including 18 abdominals, 14 caudals, and two urals. The abdominal and caudal centra are somewhat distorted in IGM 11190; however, they seem to be rectangular, slightly higher than long. Probably, this fish has 16 preural ribs. Since the trunk is entirely scaled, it is not possible to see more

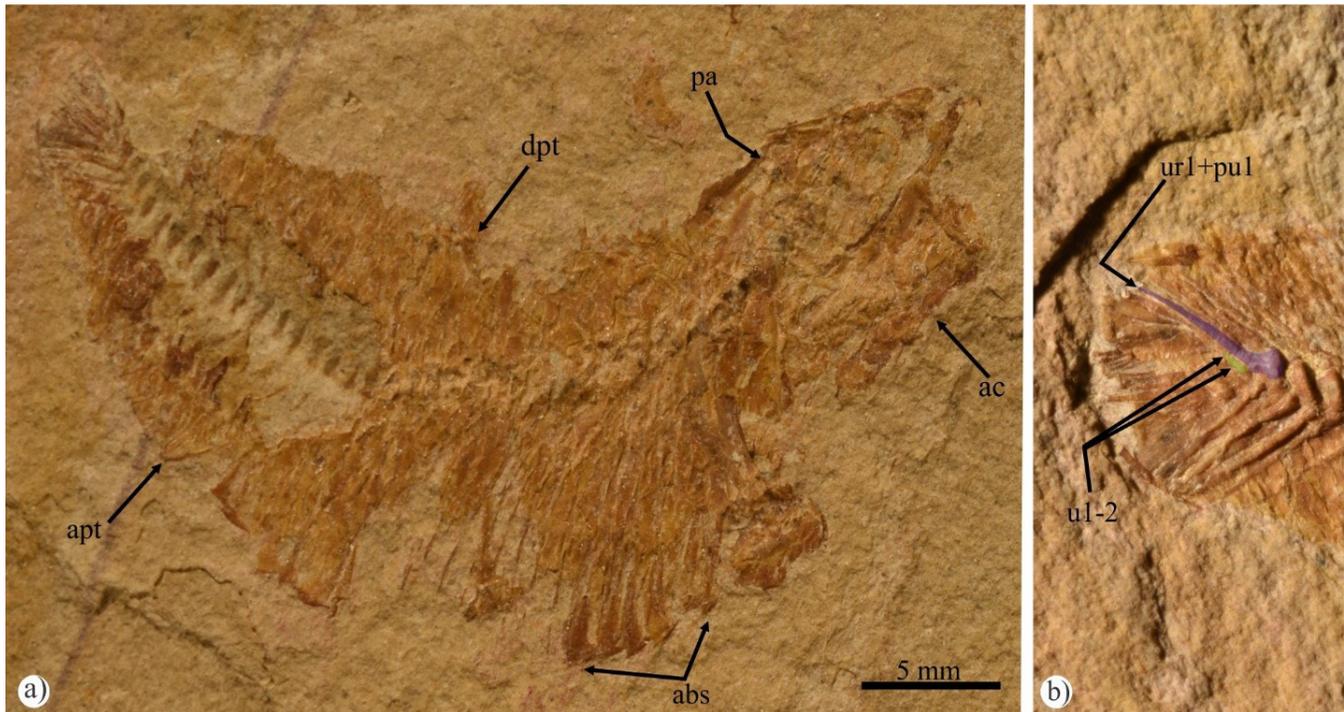


Figure 3 Specimen IGM 11190, representative of an indetermined genus and species belonging to of the family Clupeidae from the San José de Gracia quarry; a) general view of the specimen; b) closeup of the caudal fin. Abbreviations: abs, abdominal scutes; ac, anterior ceratohyal (no pierced); apt, anal pterygiophore; dpt, anal pterygiophore; pa, parietal; ur1+pu1, ural 1 and pleural centrum 1 fused; u 1-2, ural centra 1 and 2.

details of the vertebral column and associated bones.

The pectoral fin is not preserved; however, in the pectoral girdle the cleithrum is a C shaped bone that is slightly widened at its ventral end. Behind this bone, there are two postclithra, the dorsal one is ovoid and the ventral one is rod-like. The single visible dorsal pterygiophore shows that the dorsal fin rises near to the middle of the SL. There are 10 visible anal pterygiophores showing that the anal fin is located between 75 and 82.5% of the SL (30 and 35 mm). The pelvic and caudal fins are not preserved.

The trunk entirely covered for ovoid cicloid scales, taller than long, ornamented with numerous concentric circulli. No dorsal scutes are preserved in this fish. A ventral row of at least 14 abdominal scutes extends between the isthmus and the first anal pterygiophores. The abdominal scutes are triangular; they have a thickened ventral edge that projects back and forth from small spines, as well

as a laminar triangular wing slightly curved forward, directed upward, and covering less than one third of the abdominal area below the vertebral column. These scutes barely overlap each other, leaving space to observe the extension of the ribs. In the caudal skeleton, the preural 1 and the first uroneural bone are fused and there is a caudal diastema present between the hypurals 2 and 3.

Remarks: IGM 11190 is recognized as part of the Clupeoidei (Clupeomorpha, Clupeiformes) because it shows the diagnostic features of this clade, described by Grande, 1985). This has the abdominal scute series as all clupeomorphs; its parietals are entirely separated by the supraoccipital bone and its anterior ceratohyal is not medially pierced due the loss of the beryciform foramen, as occur in all clupeiforms (see Grande, 1985); additionally, as in all clupeoideis, this specimen shows the fusion of the first uroneural with the first preural centrum, the ural centra are relatively small in relation with the most caudal preural centrum,

confirms that this fish belongs to the mentioned Suborder.

In addition, based on data from Grande (1985), it is possible to consider IGM 11190 as part of the superfamily Clupeoidea. The number of pleural ribs to pleural vertebrae ratio of IGM 11190 (15/32) is 0.46 and matches with that known for members of this clade (0.47 to 0.71).

The family Clupeidea is an ancient living group with a fossil record dating back to the Barremian-Aptian (Early Cretaceous) in marine deposits of the Morro de Chaves Formation, northeastern Brazil (Figuereido, 2009; Lavoué *et al.*, 2013; Malabarba and Di Dario, 2017). Although, IGM 11190 deserves an accurate study; additional fossil materials are required to reach an accurate taxonomical determination of this clupeid fish.

Order Aulopiformes Rosen, 1973
Family Dercetidae Woodward, 1901
Genus *Dercetis* Agassiz, 1833

Dercetis sp.

Type species: According Taverne (2005a) the type species of this genus is *Dercetis elongatus* Agassiz, 1833 (p. 258). The first reference of this genus was published by Agassiz (1833, p. XLIII) but its description was published ten years later by Munster and Agassiz, 1843 (in Agassiz, 1843, p. 258), from the Turonian and Senonian marine deposits of England [Goody (1969) recognized *Dercetis scutatus* Agassiz, 1834, from the Turonian and Campanian deposits of Baumberg, in Westphalia, Northwestern, Germany, as the type species of *Dercetis*; however, this can be considered a *nomen nudum* because the single specimen of this species was never illustrated and now is lost].

Referred material: IGM 11191, an elongated almost complete fish without caudal fin exposing its left side (Figure 4).

Description: IGM 11191 is an elongated and longirostrine fish of at least 140 mm length that exposes its left lateral surface, in which the caudal fin is not preserved (Figure 4). The head is 10 mm

high and 38 mm length. The skull length is about 30 mm; this is divided into the length of orbital-postorbital and preorbital (= snout) regions, which lengths are 12 and 18 mm (or 40 and 60% of the skull length) respectively. The elongated trunk is slightly lower and uniformly high all along.

Although bad preserved, in the head both jaws are projected forward to the same level. The dentary shows two rows of slightly curved, conic teeth, of similar size. The premaxilla has a discernible wide and high ascending process. The right frontal bone is almost entirely lost; however, above the orbit the surface of this bone has some small tubercles.

This fossil preserves seventy vertebrae showing the shape of thread reels, in which the intervertebral articulation surfaces are wide and massive while their middle part is strongly constricted. Although, the vertebrae are about 1.5 time longer than high in the abdominal region; they tend to be less long. The last two vertebrae preserved in this specimen, probably the first preural centra, are slightly higher than long. A single transverse process laterally rises from the anterior half of each abdominal centra. At least 40 vertebrae are abdominal and attach with thin preural ribs. The neural and hemal spines are short and curved backward. Numerous epineurals and epipleurals are present all along the trunk.

The pectoral fin is behind the middle of the pectoral girdle, between the vertebral column and the abdominal border of the trunk and consists of 11 rays. The pre-pelvic distance is 55 mm; this fin is located 25 mm behind the pectoral girdle, is opposed to the middle part of the dorsal fin consisting of at least 5 rays. The dorsal fin is 52 mm length and is located between the 54 mm (= predorsal length) and 106 mm of the fish length consisting of about 40 rays. The anal fin is extremely short; its length is just about 8 mm and it is located between the 110 mm (= preanal length) and 118 mm of the fish length consisting of eight rays.

Although these are poorly preserved, it is possible to see that IGM 11191 has three rows of smooth and triradiate scutes along the body; one

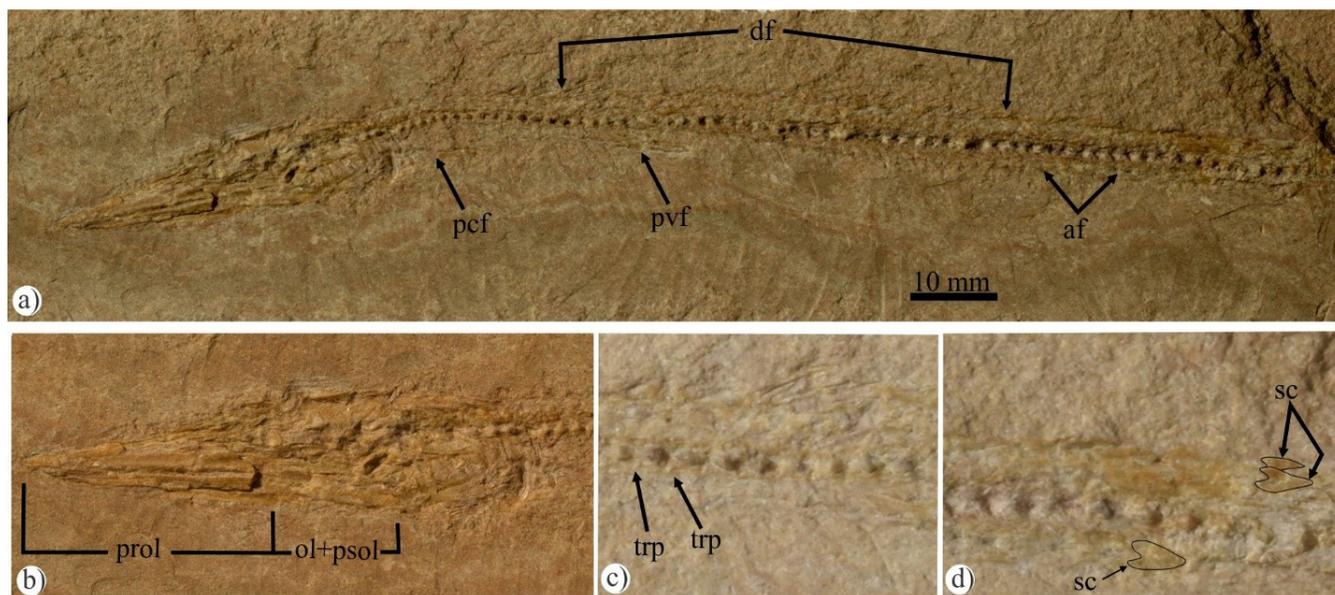


Figure 4 Specimen IGM 11191, *Dercetis* sp. from the San José de Gracia quarry; a) general view of the specimen; b) Details of the head; c) close up of the abdominal region; d) close up of the posterior part of the trunk; Abbreviations: af, anal fin; df, dorsal fin; ol+psol, orbital length + postorbital length; pcf, pectoral fin; prol, preorbital length; pvf, pelvic fin; sc, scutes; trp, transverse process.

along the dorsal body edge; another along the ventral body edge; and the last covers the flank of the trunk just above the vertebral column.

Remarks: Recently Vernygora *et al.* (2017) reviewed the phylogenetic relationship of 14 genera included in the clade Dercetidae; they point out that this family is supported on a single non-ambiguous synapomorphy: the presence of reduced neural spines. Other homoplastic features are the presence of shallow head, long snout, and arrow shaped flank scutes. IGM 11191 shows all these characters supporting its inclusion within this family.

Vernygora *et al.* (2017, suppl. 1) also revealed two monophyletic groups within Dercetidae. One of these groups includes *Dercetis* as sister group of *Cyranichthys* Taverne, 1987, and *Ophiodercetis* Taverne, 2005b; these three genera share the presence of ascending process in the premaxilla, a single pair of transverse processes on abdominal vertebrae, a long dorsal fin (with more than 20 rays) and more than one tooth row in the mandible. IGM 11191 is included within this *Dercetis-Ophiodercetis* group because it shows the first three of these characteristics.

Taverne (2006, fig. 13, p. 41–42) pointed out that *Dercetis* differs from other dercetid genera in at least four characteristics. Although in all dercetids the preorbital region or snout is elongated; in *Dercetis* the snout is relatively short and does not exceed 60% of the skull length. Contrary, in other dercetids the snout is considerably longer (at least two times than the orbito-postorbital part of the neurocranium), including the recently published *Candelarhynchus* Vernygora, Murray, Luque, Parra Ruge, and Páramo Fonseca, 2017. In *Dercetis*, the ascending process of the premaxilla is well developed whereas in other dercetids this disappear or is reduced to a tiny nodule. Here again, IGM 11191 show both features and its identification as a representative of *Dercetis* is undeniable.

Although once the genus *Dercetis* included a large number of species (see Goody, 1969; Silva and Gallo, 2011); today, two nominal species are included within the genus, the type species *D. elongatus* from Turonian-Senonian deposits of England and *D. triqueter* Pictet, 1850, from Santonian sediments of Sahel Alma, Lebanon. The temporal distribution of this genus could reach the limit of the Cretaceous if the finding of

Dercetis in Maastrichtian sediments of Muwaqqar Chalk Marl Formation, near Harrana, Lebanon, reported by Kadummi (2006) is confirmed.

Both nominal species of *Dercetis* differ in the proportions of the skull, of the preorbital region of the cranium, the lengths of the preorbital-postorbital regions of the skull are 60% and 40% in *D. triqueter* while in *D. elongatus* they are 50% (Woodward, 1901, p. 184–185; Goody, 1969, p. 62). In the present context, IGM 11191 resembles *D. triqueter*; however, it is desirable to have more specimens from San José de Gracia before reaching a conclusion. At this time, IGM 11191 is referred to as an indeterminate species of the genus *Dercetis*; however, it is noticeable that this specimen represents the first finding of this genus in America.

Family Enchodontidae sensu Fielitz, 2004

Genus *Enchodus* Agassiz, 1833 (p. XLIV)

Enchodus sp.

Type species: *Enchodus lewesiensis* (Mantell, 1822), Turonian and Senonian rocks of the Middle Chalk in Lewes, Sussex, southeastern England (Goody, 1969).

Referred material: IGM 11192, a complete palatine bone bearing an incomplete tooth (Figure 5).

Description: IGM 11192 includes an almost complete palatine bone that is exposed on its labial surface. This stout bone and its terminal tooth form a squad-like structure. The palatine is 42.97 long and 7.43 mm high; its ventral edge is almost straight and the dorsal is slightly curved (Figure 5). The medial surface of this massive bone is rounded and practically smooth, being wider along and near its ventral edge. A shallow groove for the palato-premaxillary ligament is projected backward from the middle of the dorsal edge of this palatine; unfortunately, neither the end of this groove nor the groove for the palate-maxillary ligament are preserved. The palatine has a remarkable wide rounded cavity with rough surface

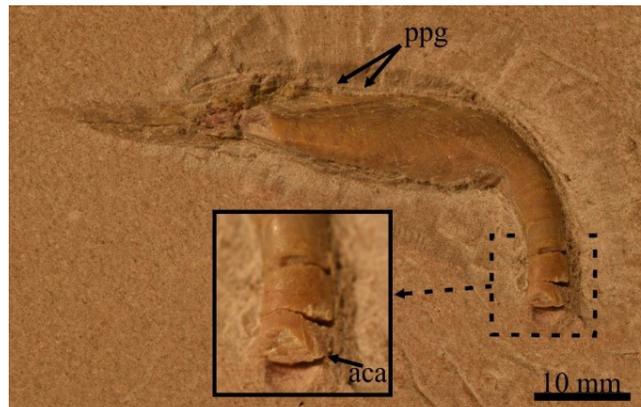


Figure 5 Specimen IGM 11192, *Enchodus* sp. from the San José de Gracia quarry, showing a close up of the palatine tooth in cross section. Abbreviations: aca, anterior carina; ppg, groove for the palato-premaxillary ligament.

opening on the lateroventral surface of the bone just behind the palatine tooth.

Although, the tip of the single stout palatine tooth is not preserved; this ankylosed structure should be straight and smooth. The basal and preserved part of this tooth is 11 mm measured from the ventral surface of the palatine; however, its total length could reach 20 mm or about 50% of the palatine length. This tooth is slightly tilted forward forming an obtuse angle, of about 100°, with the ventral edge of the palatine. The position of this tooth also is clearly terminal, the palatine does not project in front of this tooth. The broken surface of the palatine tooth is ovoid, compressed laterally, and has an anterior cutting edge or carina (Figure 5).

Remarks: The genus *Enchodus*, commonly named as saber-tooth fish, consists of about 70 nominal species, being the most diverse fossil member of the family Enchodontidae (Order Aulopiformes). The species of *Enchodus* were active Upper Albian-Eocene marine predators that inhabited practically all the seas of the planet (Fielitz and González-Rodríguez, 2010). In Mexico, fossil of *Enchodus* have been collected in numerous Albian-Maastrichtian deposits throughout Coahuila, Tamaulipas, Nuevo Leon, San Luis Potosi, Hidalgo, Guerrero, and Chiapas (Maldonado-Koerdell, 1956; Blanco-Piñón 1998;

González-Barba and Espinosa-Chávez, 2005; Alvarado-Ortega *et al.*, 2006a, 2006b; Giersch *et al.*, 2008; Díaz-Cruz *et al.*, 2016; among others); however, only five species have been identified in Mexico at the species level.

These comprise complete specimens of the species *E. zimapanensis* Fielitz and González-Rodríguez, 2010, from Albian-Cenomanian deposits of the Muhi quarry, Hidalgo, central Mexico; isolated palatine teeth referred to *E. petrosus* Cope, 1874, *E. gladiolus* (Cope, 1875), and *E. ferox* Leidy, 1855, from Maastrichtian deposits of the Ocozocuatla Formation, Chiapas, southern Mexico, (Carbot-Chanona and Than-Marchese, 2013); as well as possible representatives of *E. venator* Arambourg, 1954 from the Turonian deposits of Vallecillo quarry, Nuevo León, northern México, which were reported by Giersch (2014) as *E. cf. venator*. Among these, *E. zimapanensis* is a remarkable because it represents the oldest species known of the genus (Albian-Cenomanian) and the only species restricted to Mexico.

Despite its taxonomic richness, in the most recent systematic studies on *Enchodus*, only 18 species have been considered (Goody, 1976, Chalifa, 1996, Bogan and Agnolin, 2010, among others) and just 14 of them were included into the last phylogenetic approaches (Fielitz, 2004; Fielitz and González-Rodríguez, 2010; Cavin *et al.*, 2012; Díaz-Cruz *et al.*, 2016; Holloway *et al.*, 2017; among others). Considering the results of Díaz-Cruz *et al.* (2016); the presence of a single palatine tooth is a homoplastic condition shared only for *Rharbichthys ferox* Arambourg, 1954 and all members of the family Enchodontidae minus *Unicachichthys multidentatus* Díaz-Cruz, Alvarado-Ortega and Carbot-Chanona, 2016. *Unicachichthys* is a peculiar enchodontid fish, in which the palatine is multidentated.

Among enchodontids with a single palatine tooth, the cross section of this tooth is symmetrical, ovoid, and shows a single anterior carina (uni-carinated) only in five species, including *E. shumardi* Leidy, 1856; *E. dirus*, Leidy, 1857; as well as *E. gladiolus*, *E. zimapanensis*, and the unnamed

Enchodus species from Gavdos (Parris *et al.*, 2007; Fielitz and González-Rodríguez, 2010; Cavin *et al.*, 2012; Holloway *et al.*, 2017; among others). In contrasts, in other enchodontids in which the palatine tooth is known, the cross section of this tooth is rounded without carinae or laterally compressed and bi-carinated. In *Rharbichthys ferox* this tooth is also ovoid in cross section and uni-carinated; however, this species clearly differs from IGM 11192 and the nominal *Enchodus* spp. because its single palatine tooth is strongly tilted forward and so extremely developed that it is noticeably longer than the palate bone (Fielitz, 2004, fig. 4E). Moreover, among the group of five species of *Enchodus* noted above, in *E. zimapanensis* and the unnamed *Enchodus* species from Gavdos the palatine tooth shows longitudinal striations (Fielitz and González-Rodríguez, 2010; Cavin *et al.*, 2012) whereas in *E. shumardi*, *E. gladiolus*, and *E. dirus* this tooth is entirely smooth (Parris *et al.*, 2017).

Although the palatine tooth of IGM 11192 shares the characteristics discussed above with *Enchodus shumardi*, *E. gladiolus*, and *E. dirus*, including a smooth its palatine that in cross section is symmetrical, ovoid, and is uni-carinated (Figure 5); it is not possible to advance further in the attempt to determine the taxonomic identity of this fossil from San José de Gracia quarry. Another feature shared by IGM 11192 and these same species is the anterior terminal position of the palatine tooth. Although this feature deserves to be analyzed in greater detail and beyond the scope of this work; it is noteworthy that in other species of *Enchodus* where the palatine is well known, as well as in *Rharbichthys ferox*, the palatine tooth is not terminal because this bone has a projection in front of the tooth base (Goody, 1969, fig. 36; Fielitz, 2004, fig. 4E; Chalifa, 1996; figs. 3, 5, 6; among others).

Nonetheless, the taxonomical approximation of IGM 11192 with the three nominal *Enchodus* species noted above is geographically congruent since all of them are present only in North America, within a range between Late Cenomanian

and Early Maastrichtian (Parris *et al.*, 2007). The distribution of *Enchodus gladiolus* includes Maastrichtian deposits of the southern part of México (Carbot-Chanona and Than-Marchese, 2013) and probably also reaches Maastrichtian deposits of South America (Bogan and Angolin, 2010). Although based on the present context, IGM 0001 is considered as an uncertain species of *Enchodus*; its morphological features suggest that it is related with or belongs to some one of these three North American *Enchodus* species.

Order Sphenocephaliformes Rosen and Patterson,
1969

Family Sphenocephalidae Patterson, 1964

Genus *Xenyllion* Wilson and Murray, 1996

Xenyllion sp.

Type species: *Xenyllion zonensis* Wilson and Murray, 1996, from the Early Cenomanian deposits of the Fish Scale Formation, in Alberta, Canada.

Referred material: IGM 11193, complete and articulated specimen (Figure 6).

Description: High diamond shaped fish, with the skeleton completely preserved but with the trunk somewhat curved upward. In this fish the standard length is about 18 mm, the head length is about 7 mm, and the body height is 8 mm. The short dorsal fin consists of 4 thin spines plus six or seven rays; it is located just at the end of the first half of the SL (predorsal length = 8mm). The anal fin is short, is composed of at least four thin spines plus eight rays, and located in the last quarter of the SL (preanal length = 14 mm). The pectoral fin consists of about 10 rays and is located just above the ventral border of the trunk. The pelvic fin consists of about seven rays plus an anterior rudimentary (short) spine; this fin is located in the middle of the distance between the pectoral and the anal fin base, opposed to the middle part of the dorsal fin (Figure 6).

Although bones of the head are strongly cracked, the skull is triangular, with a high supraoccipital crest, and mainly roofed by the frontal

bones. The lower jaw is projected upward from the posterior edge of the orbit. The premaxilla is about triangular with ascending process as short as the articular process, which is separated by a shallow notch. The maxilla is a bar-like structure projected up to the anterior end of the premaxilla. No teeth are exposed.

The preopercle is almost entirely preserved as an impression; however, it shows a short horizontal limb, whose length is contained four times in the height of the vertical limb of this bone. Both preopercular limbs form an angle of about 90°. There is an impression of a lateral flange along the way of the preopercular sensory canal, which seems to be open along a groove. The posterior edge of this preopercular flange is sinuous. Also, there are impressions of two long spines in the posteroventral preopercular edge.

Vertebral column consists of 26 centra, including about 13 abdominals, 11 preurals, and two urals. About 12 thin and short preural ribs are present just covering the dorsal half of the abdominal cavity. Neural and hemal arches are small and their respective spines are thin. Some epineurals and epipleurals are preserved; however, their position, number, and sizes are unclear. There is a single supraneural bone located above the neural spines of the anterior abdominal centra. In the caudal skeleton, the preural 2 bears a complete, long, and thin neural spine. The preural 1 has a triangular short neural arch and above of this, there are two epineurals. The trunk is covered with cycloid scales with crenate posterior edge and strong circuli on the anterior half.

Remarks: Sphenocephalidae is a family of extinct acantomorph marine fishes erected by Patterson (1964): the species *Sphenocephalus fissicaudus* Agassiz, 1839 (p. 4), from the Campanian deposits of Baumberg, Westphalia, Germany. Later, another German species from Sendenhorst, Westphalia, was described as *Sphenocephalus brachypterygius* Rosen and Patterson, 1969. Recently, *Xenyllion zonensis* Wilson and Murray, 1996, from the Early Cenomanian deposits of the Fish Scale Formation, in Alberta, Canada, was described

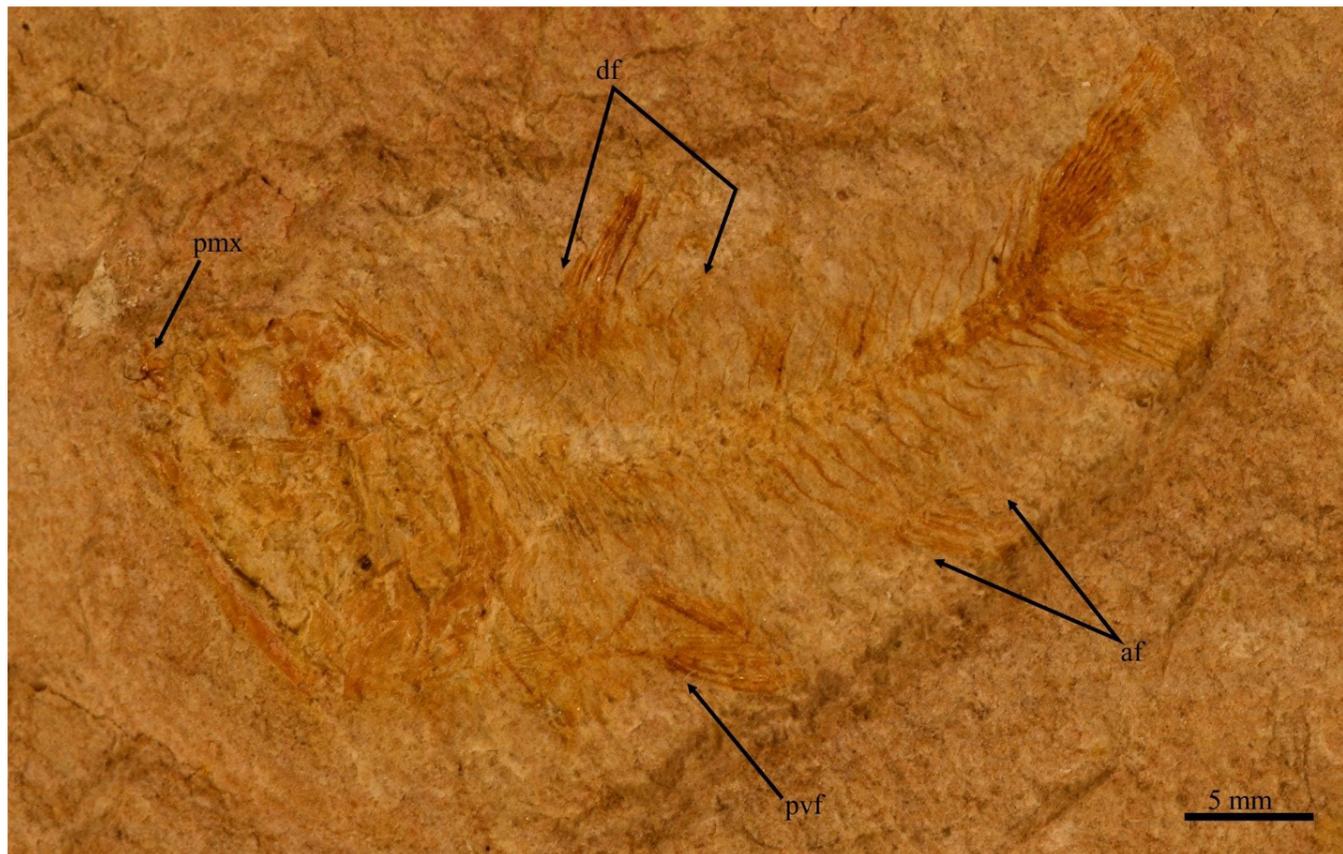


Figure 6 Specimen IGM 11193, *Xenyllion* sp. from the San José de Gracia quarry. Abbreviations: af, anal fin; df, dorsal fin, pcf, pectoral fin; pmx, premaxilla; pvf, pelvic fin.

as a new sphenoccephalid member. Finally, *X. stewarti*, Newbrey, Murray, Wilson, Brinkman and Neuman, 2013, from the Early Cenomanian deposits of Mowry Formation in Vernal, Utah, USA was described and added to *Xenyllion*. In the following paragraphs, it is demonstrated that IGM 0005 is a member of the family Sphenoccephalidae and the genus *Xenyllion*.

The presence of dorsal and anal spines in this fish support its inclusion within Acanthomorpha, a group well characterized for the presence of spines in the dorsal and anal fins (Nelson *et al.*, 2016). The Paracanthopterygii is a problematic group of living fishes, with the oldest record dating back to the Cenomanian (Newbrey *et al.*, 2013). This group was named by Greenwood *et al.* (1966) as a Superorder and since then its monophyly and composition have been discussed (Rosen and Patterson, 1969; Wiley and Johnson,

2010; Borden *et al.*, 2013; among others). The presence of a complete neural spine on second preural, two epipleurals in the caudal support, and a single supraneural present are features of Paracanthopterygii (see Greenwood *et al.*, 1966, p. 352, 353) and Murray and Wilson (1999) that are seen in the specimen IGM 11193.

One of two diagnostic characters of the order Sphenoccephaliformes described by Rosen and Patterson (1969) are present in IGM 11193, its preopercle has large posteroventral spines. The other of these features, the opercle with posteroventral edge curved, is unknown in IGM 11193. Additionally, this Mexican specimen shows some characters included in the diagnosis of the single family of this order. According to Patterson (1964) and Newbrey *et al.* (2013), the diagnostic characters of Sphenoccephalidae include the presence of a pelvic splint (= rudimentary anterior spine),

seven pelvic fin rays, as well as anal and dorsal fin with four to five spines each.

Newbrey *et al.* (2013) provided a comparative exercise between the four sphenoccephalid species. Based on their results, IGM 11193 is identified here as a new representative of the genus *Xenyllion* because they share two characteristics: the trunk is covered with crenate cycloid scales, and the preopercle shows similar proportions (both limbs form an angle of 90° and the horizontal limb is short and contained four times in the height of the preopercular vertical limb). In contrast in both *Sphenocephalus* species, the preopercles have a relatively long horizontal limb (its length is the half of the height of the vertical preopercular limb) and both limbs form a posteroventral obtuse angle of about 120°; *S. fissicaudus* has oval spinoid cycloid scales whereas *S. brachypterygius* are rather rounded, apparently do not show spines, and the circuli form a triangular wave-form pattern.

The poor preservation of the skull in IGM 11193 prevents carrying out a deeper comparative analysis to determine its taxonomic identity with better precision. Despite this, the finding of IGM 11193 represents the first occurrence of a sphenoccephalid fish in Mexico and the expansion of the geographical distribution of the genus *Xenyllion* in America.

Orden Squamata Opperl, 1811

Superfamily Mosasauoidea Camp, 1923

Family Mosasauridae Gervais, 1853

Subfamily Yaguarasaurinae Palci, Caldwell, and Papazzoni, 2013

Referred material: IGM 11194, isolated premaxilla and right maxilla (Figure 7).

Description: A preliminary description of this specimen is available in Zavaleta-Villareal (2015). The premaxilla is an unpaired compound bone that is drop-like shaped in dorsal view; its anterior edge is rounded, the posterior tapers forming an acute angle where the internarinal bar arises. Although this premaxillary process is almost entirely

preserved as an impression; the thin internarinal bar is at least as long as the rest of the premaxilla. The dorsal-lateral surface of this bone is curved and exposes numerous pores of the ophthalmic nerve.

Ventrally, this bone is rather flat and is divided into four regions framed by two crests, one transversal that runs from side to side and other longitudinal that rises behind the anterior edge of this bone and it projects backwards. In each of the two previous regions two teeth are presented forming a line; while the two posterior regions have no teeth (Figure 7).

The right maxilla is an elongated bone tapered anterior and posteriorly, with the ventral edge straight and the dorsal one somewhat sinuous. The premaxilla-maxilla suture border is straight and extends over two maxillary teeth. The posterior edge of this bone is inclined downward from the top, but its bottom develops a thin and long postero-ventral process where there are no more teeth. The labial surface is curved and has a longitudinal axis protruding near the ventral edge of the bone. Although the surface of the maxilla shows numerous pores of the maxillary nerve; nine of these form a row below the medial longitudinal edge of this bone and are arranged parallel to the ventral edge of this bone (Figure 7).

Eight conical teeth and three empty alveoli are placed on the maxillary alveolar border; these are ordered following the dental formula 1, 2, 3, 4, x, 6, x, 8, 9, x, 11. The teeth have ankylosed crowns; they have broad bases protruding from the alveolar maxillary edge. All teeth are entirely smooth and strongly curved backward. In cross section, teeth are rounded at the base, oval in the middle part, and form a weak anterior cutting edge or carina near to the tip (Figure 7).

Remarks: The presence of an unpaired premaxilla (resulting of the fusion of a couple of premaxillae) with four teeth in IGM 11194 supports its identification as a member of the Family Mosasauridae (see Leblanc *et al.*, 2012, among others). According to Palci *et al.* (2013), the occurrence of tooth roots largely exposed in

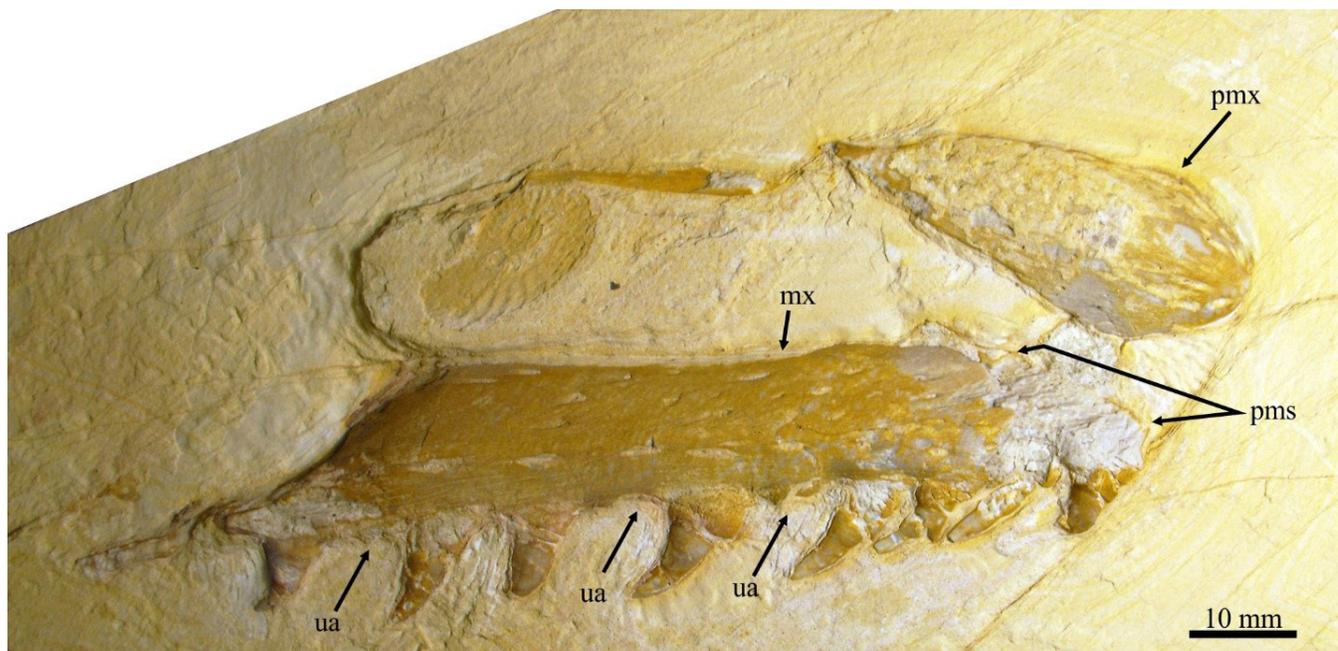


Figure 7 Specimen IGM 11194, indeterminated *Yaguarasaurinae* from the San José de Gracia quarry. Abbreviations: mx, maxilla; pmx, premaxilla; pms, premaxilla-maxilla suture; ua, unoccupied or empty alveolus.

medial view of jaw bones represents one of the synapomorphies supporting the monophyly of the subfamily *Yaguarasaurinae*, which gathers three genera. There include two monospecific genera, *Yaguarasaurus columbianus* Páramo, 1994, collected in Turonian sediments belonging to Villeta Formation along the Magdalena Valley, Colombia (Páramo, 1994; Páramo-Fonseca, 2000); *Russellosaurus coheni* Polcyn and Bell, 2005, from the middle Turonian deposits of the Arcadia Park Shale, approximately 15 cm above Kamp Ranch Limestone, at Cedar Hill, Dallas County, Texas, USA; as well as to species of the genus *Romeosaurus* Palci, Caldwell, and Papazzoni, 2013, collected in Late Cretaceous localities of Verona, northern Italy, which recently were named as *R. fumanensis* (from the middle Turonian to early Santonian deposits near of the S. Anna di Alfaedo and Fumane villages) and *R. sorbinii* (from the middle Turonian deposits near of the S. Anna di Alfaedo Villages. These characteristics are present in IGM 11194 supporting its recognition as part of this subfamily (Figure 7).

Among other characteristics, the nominal members of *Yaguarasaurinae* differ from each other by the size of the premaxillary-maxillary suture, as well as the number and shape of maxillary teeth. The suture maxillary-premaxillary is over three maxillary teeth in *Romeosaurus* while it is over only two maxillary teeth on *Yaguarasaurus* and *Russellosaurus* (Palci *et al.*, 2013, p. 600). The number of premaxillary teeth is 14–15 in *Yaguarasaurus* (Páramo-Fonseca, 2000, p.127), 13–15 in *Romeosaurus* (Palci *et al.*, 2013, p. 600), and 16 in *Russellosaurus* (Polcyn and Bell, 2005, p. 323). The maxillary teeth are moderately curved backward, probably smooth and with no carenae in *Yaguarasaurus* (Páramo-Fonseca, 2000, fig. 3); and contrary, teeth are strongly curved, covered with fine medial longitudinal striations, and probably have no carenae in *Russellosaurus* (Polcyn and Bell, 2005, fig. 2). Finally, teeth are subcircular in cross section and strongly curved backward in both species of *Romeosaurus*; however, in *R. fumanensis* there are 15 maxillary teeth, entirely smooth, and bearing weak anterior and posterolateral carinae whereas *R. sorbinii* there

are 13 maxillary teeth that are ornamented with longitudinal ridges and apparently with no carinae (Palci *et al.*, 2013, p. 601, figs. 2, 4, and 7).

IGM 11194 shows a peculiar combination of features that avoid determining its belonging to any previously known yaguarasaurine genus. In this Mexican fossil the maxillary-premaxillary suture is short as in *Yaguarasaurus* and *Russellosaurus* (extended over two maxillary teeth). The number of maxillary teeth (11) in IGM 11194 is lower than in the nominal yaguarasaurins (13–16). Contrary to *R. sorbinii*, teeth of IGM 11194 are superficially smooth. Peculiarly, in cross section the teeth of IGM 11194 are rounded in the base, ovoid in the middle and clearly compressed laterally in the tip in which there is a small anterior carina. Although the accurate taxonomic identification of IGM 11194 is desirable; this represents a major effort that goes beyond this work. The discovery of IGM 11194 represents the first record of the subfamily Yaguarasaurinae in Mexico.

3.2. OTHER ELEMENTS IN THE FOSSIL ASSEMBLAGE FROM SAN JOSÉ DE GRACIA

For now, about a hundred fossil specimens from San José de Gracia have been catalogued into the Colección Nacional de Paleontología (IGM), UNAM, and a similar amount is within the paleontological collection of the Benemérita Universidad Autónoma de Puebla (BUAP). These fossils show different degrees of conservation, from isolated and fractured elements to complete specimens. Until now, they represent a fossil assemblage that include complete ammonites, belemnites, inoceramids and other indeterminate ostreids, scarce impressions of plants, and vertebrates.

Centimetric coprolites are also preserved in San José de Gracia quarry; these are preserved as phosphatized masses deposited on the bedding planes. In some small fishes, the gastrointestinal contents also preserve similar phosphatic mass. Bony cavities are filled with calcium carbonate crystals. Practically all the fossils from this site are distorted, compressed in a dorsoventral plane, but

only some of them show a horizontal elongation that distorts their shape in this plane.

Among the fossils from San José de Gracia, the heteromorph ammonites are remarkable because some of them represent the families Hamitidae and Anisoceratidae, which ranges from Albian to Turonian (Wright *et al.*, 1996) and Albian to Coniacian (Kennedy and Gale, 2015), respectively. Some inoceramids specimens deposited in the BUAP were identified by Gloria Alencáster (JAO, pers. observ.) as potential representatives of the species *Inoceramus (Mityloides) labiatus* Schlotheim (1813) that has been widely recognized as a fossil species index of the Turonian (Riccardi, 1981; Damian and Lazăr, 2005).

If the temporal distributions of Hamitidae and Anisoceratidae ammonites, the pelecypod *Inoceramus labiatus*, and the vertebrates described above are considered, it is possible to attribute a Turonian age to the fossiliferous strata exploited in the San José de Gracia quarry (Figure 8). For the moment, the formal studies of all the fossil invertebrates and the possible microfossils of San José de Gracia are pending; once these are

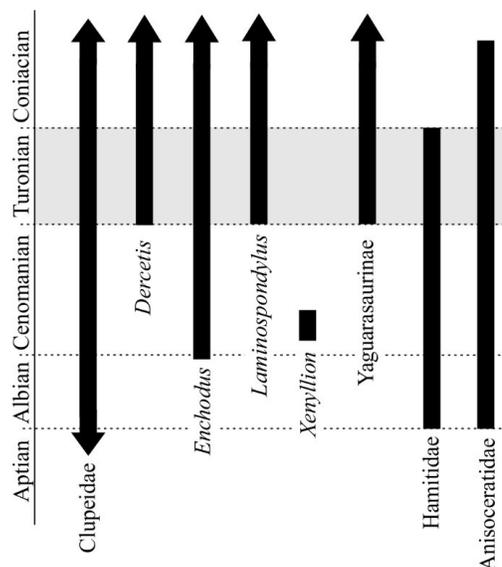


Figure 8 Specimen IGM 11194, indetermined Yaguarasaurinae from the San José de Gracia quarry. Abbreviations: mx, maxilla; pmx, premaxilla; pms, premaxilla-maxilla suture; ua, unoccupied or empty alveolus.

developed, it will be possible to reach a more accurate age for this deposit.

4. Discussion

Maps already published by Servicio Geológico Mexicano (2006) show that the Oligocene-Miocene continental deposits of the Coatzingo Formation are exposed around the town of San José de Gracia (Figure 1). The base of this continental formation is composed of limestones with some intercalated flint horizons that are covered with a sequence of volcanic ash and sand layers deposited in a lacustrine environment of low energy within the Zacapala Basin (Beraldi-Campesi, 2003). The lithology of the San José de Gracia quarry does not match with that of the Coatzingo Formation; hence, there is no possibility that the rocks of this new locality belong to that formation.

Besides, the local geological framework of San José de Gracia quarry is somewhat complex. In the south, Cambrian-Devonian shales and phyllites belonging to the metamorphic complex of the Acatlán Complex are the oldest rocks in the region (Ortega-Gutiérrez, 1978). Above these Paleozoic strata, there are terrigenous red layers and calcareous lenses from the Upper Jurassic-Early Cretaceous (Neocomian-Aptian). Near the town of Tepexi de Rodríguez, a few kilometers south of San José de Gracia, Pantoja Alor (1992) described the Tlayúa Formation as an incomplete sequence, which is recognized as a sequence of at least 300 meters thick composed of marine strata of Aptian-Cenomanian lower age. The low topographic parts near San José de Gracia and Tepexi de Rodríguez show continental Tertiary clastic sediments belonging to different units; among these, the Paleocene-Eocene conglomerates of the Balsas Group (Fries, 1960) lie on the eroded surface of the Tlayúa Formation; above, a Eocene-Oligocene sequence of fluvial, lacustrine and alluvial deposits of the Pie de Vaca Formations as well as a Pleistocene sequence of lacustrine travertines belonging to the Agua de

Luna Formation are present (Pantoja-Alor, 1992; Beraldi-Campesi, 2003; Applegate *et al.*, 2006; among others).

The hills south to San José de Gracia are formed by the Cenomanian massive gray marine limestones of the Morelos Formation, which was originally described by Fries (1960) as a shallow platform marine deposit. Otherwise, the hills and highlands to the East and Northeast of José de Gracia consists of the Aptian marine limestones belonging to Miahuatepec Formation, whose fossil-lacking strata were deposited in shallow conditions (Buitrón and Barceló-Duarte, 1980) (Figure 1).

Regarding its geological context, it is possible to conclude that the marine strata exploited in the San José de Gracia quarry do not belong to any of the Tertiary continental geological units previously documented in this region (Coatzingo, Pie de Vaca, and Agua de Luna). The fossiliferous marly strata of San José de Gracia quarry cannot be identified as part of any of the Early-middle Cretaceous marine limestone formations (Tlayúa, Miahuatepec, and Morelos) previously mapped near San José de Gracia (Figure 1). Therefore, the San José de Gracia fossiliferous strata represent a unit for the first time observed in this region of Puebla.

The sedimentary rocks of Turonian age closest exposed to the San José de Gracia quarry are located 20 km northwest from this site, along the Atoyac River, and belong to the Mexcala Formation, previously described by Fries (1960). This formation is a sequence of Cenomanian-Maastrichtian marine sediments deposited within a marginal depocenter, the Morelos-Guerrero Basin, occurred under moderate currents, low turbidity, and pelagic and hemipelagic environments, which today is exposed along Guerrero, Morelos, Puebla, and Oaxaca (Alencáster, 1980; López-Ramos, 1983; Hernández-Romano, 1999; Aguilera-Franco, 2000; among others). It is tempting, based on this observation, to suggest that the San José de Gracia quarry is an outcrop of the Mexcala Formation; however, comparison of the fossil assemblage

from San José de Gracia quarry with those from Turonian sites (the Arroyo Las Bocas quarry and Taxo El Viejo) of the Mexcala Formation in Guerrero already reported by Alvarado-Ortega *et al.* (2006b) shows scarce common elements (the occurrence of the genus *Enchodus* and the family Dercetidae). Today, the recovery of more geological data in the San José de Gracia quarry is problematic, given the small extension and depth of the sequence that is being exploited; hence, the final recognition of the belonging of these rocks to a formal geological unit remains as a pending task.

5. Conclusions

The marls of San José de Gracia Quarry, Puebla, Mexico, constitute an outcrop of marine fossiliferous Turonian rocks observed for the first time in this region of Puebla. Although it is possible to suggest these rocks as part of the Mexcala Formation; additional geological data are required to confirm this idea. The fossil vertebrate assemblage recovered from this new site represents an unexpected diversity. This includes taxa previously known in Mexico (such as *Enchodus*), taxa for the first time recorded in Mexico (such as *Dercetis*, *Xenyllion* and *Yaguarasaurinae*), as well as a specimen that represents the oldest record of the Family Clupeidae in the country (Figure 8).

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