Systematic paleontology and taphonomic studies of Ypresian mollusks at the Kopet-Dagh Basin, NE Iran

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ABSTRACT

Marine invertebrate fossils have long been considered important tools for age dating and stratigraphic interpretation of the Paleogene deposits of Central Asia, however information has not yet been provided from the Kopet-Dagh Basin (NE Iran). In this research, fossiliferous horizons of the Chehelkaman Formation at the Sheikh and Ghaleh-Zou sections (which have never been recognized previously), both in the Sheikh Syncline area, are discussed. These sedimentary beds overlay the terrestrial siliciclastic Pesteligh Formation and are the first evidence of marine flooding of the Paleogene transgression in the central Kopet-Dagh. Systematic paleontological studies of molluscan fossils in this succession led to the identification of ten genera, twelve species and two subspecies of bivalves and five genera and five species of gastropods. The assemblage belongs to four different fossiliferous horizons: (1) Turritellidae Dominated Assemblage (TDA), (2) Pycnodonte - Turkostrea horizon, (3) Cardiopoda - Cardium horizon and (4) Globularia shell beds in the Sheikh section and only TDA horizon in the Ghaleh-Zou section. The age of the whole succession, based on Cardiopoda subathooensis - Turritella subathooensis Zone co-occurrence with nannofossil zones NP12 and NP13, is Late Ypresian (early middle Cusian). The TDA in both sections confirms the abundance of nutrients and their bimodal orientation demonstrates effects of oscillatory waves rather than unidirectional paleo-currents. High rates of bioerosion, encrustation, fragmentation, disarticulation and corrosion of the shells in the Pycnodonte - Turkostrea horizon reveals high volumes of nutrients and energy-rich conditions. This interval, overlain by the Cardiopoda - Cardium horizon in a pavement arrangement, represents the stable conditions after a storm. The Globularia shell beds interbedded with marly beds indicate stability and a deepening trend of the environment, terminating to deeper marine shales and marls of the Khangiran Formation.

Keywords: Taphonomy, paleoenvironment, Early Eocene, Kopet-Dagh Basin, Chehelkaman Formation.

RESUMEN

Los fósiles de invertebrados marinos se han considerado herramientas importantes para la resolución de datos por estudios y la interpretación estratigráfica de los depósitos de paleógenos de Asia Central, sin embargo, aún no se ha proporcionado información de la cuenca de Kopet-Dagh (NE Irán). En esta investigación, se discuten los horizontes fosíferos de la Formación Chehelkaman en las secciones Sheikh y Ghaleh-Zou (que nunca antes se habían reconocido), ambos en el área Sheikh Syncline. Estos lechos sedimentarios se superponen a la Formación Pesteligh siliciclástica terrestre y son la primera evidencia de inundación marina de la transgresión del Paleógeno en el centro de Kopet-Dagh. Los estudios paleontológicos sistemáticos de fósiles de moluscos en esta sucesión condujeron a la identificación de diez géneros, doce especies, dos subespecies de bivalvos, cinco géneros y cinco especies de gastéropodos. El conjunto pertenece a cuatro horizontes fosíferos diferentes: (1) Conjunto Dominado por Turritellidae (TDA), (2) Pycnodonte - Turkostrea horizonte, (3) Cardiopoda - Cardium horizonte, (4) Camas de concha Globularia en la sección Sheikh y solo horizonte TDA en la sección de Ghaleh-Zou. La edad de toda la sucesión es Ypresiano tardío (Cusiano medio temprano). Esta edad está basada en Cardiopoda subathooensis - Turritella subathooensis co-ocurrencia de la zona con las zonas nannofósiles NP12 y NP13. El TDA en ambas secciones confirma la abundancia de nutrientes y su orientación bimodal demuestra los efectos de las ondas oscilatorias en lugar de las paleocorrientes unidireccionales. Las altas tasas de bioerosión, incrustación, fragmentación, desarticulación y corrosión de las conchas en Pycnodonte – Turkostrea horizonte revela altos volúmenes de nutrientes y condiciones ricas en energía. Este intervalo, cubierto por el Cardiopoda - Cardium horizonte en una disposición de pavimento, representa las condiciones estables después de una tormenta. Las camas de conchas de Globularia intercalados con lechos de mugoso, indican estabilidad y una tendencia cada vez más profunda del medio ambiente terminando en lutitas y margas más profundas de la Formación Khangiran.

Palabras clave: Tafonomía, paleo-ambiente, Eoceno temprano, cuenca de Kopet-Dagh, Formación Chehelkaman.
1. Introduction

The Kopet-Dagh Mountains of north-eastern Iran had long been recognized as a petroliferous basin (Afshar-Harb, 1969). These Mountains extend from the eastern margin of the Caspian Sea to Afghanistan, limited by the Ashgabat fault at the northern border and the Alborz Range and Central Iran from the southern part (Afshar-Harb, 1979; Buryakovsky et al., 2001). The basin was formed as a result of the southeastern extension of the South Caspian Basin by Neotethyan back-arc rifting, followed by the closure of the Palaeotethys in the Middle Triassic by the Early Cimmerian Orogeny (Robert et al., 2014). In the basin, continuous sedimentation took place from the Jurassic to the Neogene (Berberian and King, 1981; Afshar-Harb, 1994; Garzanti and Gaetani, 2002; Wilmans et al., 2009; Robert et al., 2014).

The Paleogene rocks of this sedimentary basin are Early to early Middle Paleocene Pesteligh Formation (Robert et al., 2014), the Middle Paleocene - Early Eocene Chehelkaman Formation (Rivandi et al., 2012), and the Eocene - Early Oligocene Khangiran Formation (Hadavi and Senemari, 2002). These successions are overlain by unnamed continental Neogene red beds and Quaternary alluvial deposits (Saadat-Mehri et al., 2012).

The Chehelkaman Formation is a fossil rich rock unit with nannoplankton, foraminifers, dinoflagellates, ostracods, fish remains and a molluscan assemblage (Hadavi and Senemari, 2002; Salahi, 2011; Moshirfar et al., 2013; Rivandi et al., 2013). The formation is useful and informative for stratigraphic correlation within the basin and also with the Central Asia region. Nonetheless, the macrofossil content of this formation has previously never been investigated. The only published report on the molluscan fossils of the Paleogene succession of the Kopet-Dagh Basin is from Salahi and Vahidinia (2011), who focused on the Lutetian- Bartonian oysters of the upper part of the Khangiran Formation. In the current research, we carry out the systematic paleontological, palaeoecological and paleoenvironmental interpretation of the molluscan fossils of the Chehelkaman Formation in the Sheikh Syncline area (SSA), which is located at the northeast of Bojnourd City in the central Kopet-Dagh (Figure 1).

2. Geological setting

During the Early Paleocene, the Laramide orogeny caused the retreat of the sea, followed by continental deposits of the Pesteligh Formation above the completely marine Kalat Formation (Afshar-Harb, 1994; Robert et al., 2014). The red beds of the Pesteligh Formation in the type area of the eastern Kopet-Dagh gradually change to alternating siliciclastic and carbonate beds of the Chehelkaman Formation, which indicates the first phase of the Paleogene sea transgression. Toward the upper part, gradual change in the lithology from the limestone into silty marl and marl are characteristic for the Khangiran Formation.

Along the central part of the basin, particularly in the SSA, it is difficult to follow the typical succession of the Paleogene and Neogene rock units laterally because of its lateral facies changes than the type locality area. The shallowing trend toward the central Kopet-Dagh Basin and replacement of the carbonate sequence by siliciclastic-evaporitic succession is the most important feature of the depositional system. These facies changes cause some difficulties for stratigraphic interpretations; therefore the Chehelkaman Formation has never been depicted in the published reports of the area (e.g., Afshar-Harb, 1979, 1994). It was also not mentioned in the geological maps of the area, such as 1/250,000 quadrangle map of Bojnourd (Bolourchi and Afshar-Harb, 1987) and 1/100,000 quadrangle map of Kakli (Ghaemian, 2004). It means the previous geologists depicted the Khangiran Formation directly on the Pesteligh Formation through a main stratigraphic gap, which was considered temporally equivalent to the depositional interval of the Chehelkaman Formation in the type locality area in the eastern Kopet-Dagh Basin.
Figure 1  Geological map of the studied area in central part of the Kopet-Dagh Basin; red stars point to the sections discussed (modified after Afshar-Harb, 1982). Left photograph: Outcrop at Sheikh Synclinal region with sequence of lithostratigraphic units from Upper Cretaceous to Neogene. Right upper photograph: stratigraphic section at the Sheikh area. Right lower photograph: stratigraphic section at the Ghlaeh-Zou area.
The fossiliferous beds in the SSA have been investigated in two stratigraphic sections, 22 km apart (Figures 1 and 2). The Sheikh section in the western part of the SSA is located between the coordinates N37°34′42″ and E57°32′48″ and the Ghaleh-Zou section in the southeastern part of the SSA, between the coordinates N37°31′40.7″ and E57°47′13.2″. Our fieldwork in both sections confirms the presence of the Chehelkaman Formation shell beds, but in much less thickness than the type section area (Figure 1). Like typical shell beds (Kidwell et al., 1986), they are laterally well-extended and show a high abundance of macrofossils.

3. Material and Methods

Amongst the collected bivalve and gastropod specimens, the best preserved were selected for systematic description. They were biometrically measured and their shape, outline, shell thickness, ornamentations, convexity, ribs and growth lines were determined and compared with the published descriptions of the known fauna from the other regions, especially Central Asia (Berizzi Quarto di Palo, 1970; Lan, 1997; Mathur and Juyal, 2000; Halder and Sinha, 2014). In addition, depositional and ecological conditions of the formation of

Figure 2 Stratigraphic logs at the Sheikh and Ghaleh-Zou sections.
shell beds (Dorsey and Kidwell, 1999; Dorsey and Umhoefer, 2000) and taphonomic attributes of macrofossils such as transportation, fragmentation, disarticulation, corrosion, encrustation, boring, and predation were recorded (Gordillo et al., 2014). The shells of turritelline were classified following Allmon et al. (1995) in two classes as complete and incomplete specimens. The complete shells have three to five whorls preserved and variable diameter of the twisting. Photographs were taken using a Cannon 60D camera and macro lens L100.

4. Results

4.1 SHEIKH SECTION

In the Sheikh section, the sedimentary sequence includes Cenozoic sedimentary units, from the Pesteligh Formation to Neogene red beds, with a total thickness of 774 m. This study was carried out on only 20 m of the upper part of the Pesteligh Formation to the lower part of the Khangiran Formation (Figure 2).

4.1.1 PESTELIGH FORMATION

The formation is characterized by 250 m of medium to thick planar-parallel bedsets of siltstone, sandstone, claystone, and interbedded thin gypsum layers. Some charophyte remains were recorded from the top claystone horizons (Figure 3). This flora was deposited in a freshwater environment.

4.1.2 CHEHELKAMAN FORMATION

The Chehelkaman Formation is mainly composed of fossiliferous and porous limestone, siltstone, calcareous sandstone, interbedded marl and evaporites. The top of the formation is composed of calcareous shell beds, siltstone and calcareous sandstone, with many mollus-

Figure 3 a: Sandy beds of Chehelkaman Formation in the Ghaleh-Zou section with turritelines at the lowermost and uppermost of the bed; b: Silty marls and gypsum beds in the boundary between Chehelkaman and Khangiran formations in the Ghaleh-Zou section; c: Lateral outcrops of Chehelkaman Formation in SSA region; as it is clear in the figure, the formation totally turns to evaporitic facies. This facies change was probably the main reason of neglected Chehelkaman Formation definition in the previous studies.
can fossils. The thickness of this formation is up to 300 m, while in the type section area at the Eastern Kopet-Dagh it is 234 m (Afshar-Harb, 1994).

In the Sheikh section, the formation is represented by 14 m of medium to thick, planar-parallel, silty to sandy white limestone, shell beds, sandstone and silty marl beds, with abundant molluscan fossils (Figure 2). This unit overlays directly the Pesteligh Formation and underlies the silty and marly Khangiran Formation. Its lower boundary with the Pesteligh Formation is sharply erosive, but its upper boundary with the Khangiran Formation is in a fining upward trend to the marine silty marls beds of the Khangiran Formation (Figure 2). The fossil rich beds, especially those with ampullinid gastropods, are associated to the uppermost part of the Chehelkaman Formation in both the type section and the Sheikh section.

4.1.3 KHANGIRAN FORMATION

The Khangiran Formation is composed of 270 m of gray to green, medium to thick planar-parallel beds of siltstone, silty marl, and marl beds. The geological study focused on the first 4 to 6 m above the lower boundary.

4.2 GHALEH-ZOU SECTION

The Paleogene sedimentary sequence in the Ghaleh-Zou section comprises the Pesteligh Formation to the Neogene red beds with a total thickness of 774 m. In this research, only 18 m of the Pesteligh Formation upper parts to the lower beds of the Khangiran Formation have been studied (Figure 2).

4.2.1 PESTELIGH FORMATION

This unit is characterized by beds of siltstone, sandstone, and claystone generally of red color (Figure 3).

4.2.2 CHEHELKAMAN FORMATION

The Chehelkaman Formation is 12 m thick; it is composed of white sandstone and siltstone with interbedded white to gray thin-bedded sandstone, and lamina sets of gypsum varying from 2 to 8 cm thick. These gypsum sets become laterally thicker to evaporitic beds (Figure 4). This unit shows the medium to thick beds of loose and non-loose sandstone, both types of beds with curved, non-parallel lamination (Figure 4). Turritelline gastropods are only found in these sandy beds. The boundary of the Chehelkaman and the Khangiran formations is characterized by a gradual fining upward, comprehending terrigenous, evaporitic and marly rocks.

4.2.3 KHANGIRAN FORMATION

The formation is 270 m thick, consisting of medium to thick beds of siltstone, silty marl, and marl which show continuous, planar and parallel lamination. The lower part of the formation consists of evaporitic beds interbedded with silty marl and marly beds (Figure 4), ranging from thin laminae to thick beds in white to gray color. From the basal part of this formation, abundant nannofossils were obtained.

Comments.— Most of the sedimentary rocks in the eastern SSA were deposited in a marine environment, as it is comprehensible from the lithology and fossil evidence (e.g., shell beds). The Chehelkaman Formation likely formed in subaqueous conditions, from very shallow subtidal to sabkha environments. The calcareous algae remains from the Khangiran Formation and the presence of a considerable amount of glauconite suggest deeper conditions, likely subtidal, implying a sea level rising. The nannofossils determined an age interval from NP12 to NP13 (Early Eocene), for the Chehelkaman Formation and the lower part of the Khangiran Formation. This age interval will be discussed below.

5. Systematic Paleontology

Phylum Mollusca Linnaeus, 1758
Class Bivalvia Linnaeus, 1758
Order Ostreoidea Rafinesque, 1815
Family Gryphaeidae Vialov, 1936
Genus Pycnodonte Fischer von Waldheim, 1835
Pycnodonte sp.
Material: five left valves, MASSH 0001 - MASSH 0005.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base, horizon 2 at 1.1 m from the base, horizon 3 at 9 m from the base.

Description.— Relatively thick shell of medium size, subcircular to ovate in outline to dorsoventrally elongate, higher than long and highly convex. Specimens show a broad but broken posterior flange. Convex cup-like and generally smooth left valve with very fine concentric growth squamae and traces of a few discontinuous radial ribs in the umbonal region. Triangular ligament area (gryphaeoid type), mostly longer than high with flat bourelets that are shorter in length to the resilifer. Catachomata is absent, but vermiculate-chomata are present. Vermiculate chomata short, rather shallow, hardly visible; muscle scar is not visible due to covering by chalky sediments. Attachment area very small to absent. Vesicular layers not recognized. Umbo relatively small, subcircular and slightly opisthogyral.

Dimensions.— valve length (L) 48 mm, valve height (H) 74 mm and maximum shell thickness of left valve 27 mm.

Taxonomic comments.— Based upon shell structure, smoother cuplike shell, shell height, triangular gryphaeoid type ligament with flat bourelets, presence of flange, types of chomata and the poor development of the ribs and attachment area (Stenzel, 1971; Machalski, 1988; Callapez et al., 2015) the specimens are representatives of the gryphaeid genus Pycnodonte.

Family Ostreidae Rafinesque, 1815
Subfamily Ostreinae Rafinesque, 1815
Genus Ostrea Linnaeus, 1758
Subgenus Turkostrea Vialov, 1936
Ostrea (Turkostrea) turkestanensis turkestanensis (Romanovskiy, 1878)
Figures 5.2a and 5.2b
1948 *Ostrea turkestanensis*, Vialov, Pl. XXII, figs. 2-4
1962 *Ostrea turkestanensis*, Gekker, Osipova, Belskaya, Pl. V, figs. 1-3
1970 *Ostrea (Turkostrea) turkestanensis turkestanensis*, Berizzi Quarto di Palo, Pl. 33, fig. 2; Pl. 34, fig. 3; Pl. 35, figs. 2-3

Material: two left valves, MASSH 0031 - MASSH 0032.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

Description.— Thick shell with small size, inequivalve, oval to subtriangular, moderately to strongly convex, with very strong convexity in median region. Massive with narrow umbonal region and inclined apex, sculptured left valve with regular, almost equal and symmetric radial ribs with equal interspace in dorsal view, wider toward the valve margin, rather deep and narrow curved posteriorly. The number of radial ribs is 22 to 28. Concentric growth lamellae, more noticeable at ventral margin. Higher than long, with small to moderately attachment area. Internal shell smooth with broad crescentic kidney-shaped muscle scar, triangular ligament area (turkostreoid type) that strongly curves posteriorly, longer than high with narrow and diagonal bourelets. Catachomata and shallow-short vermiculate-chomata are present. Vesicular layers not seen.

Dimensions.— L: 21 mm, H: 22 mm and maximum shell thickness of left valve 7 mm.

Taxonomic comments. — Small size, turkostreoid type of ligament area with kidney shaped muscle scar that has a semi flange in the posterior side.
Ostrea (Turkostrea) turkestanensis baisunensis Böhm 1899

Figures 5.3, 5.4, 5.5a and 5.5b

1948 Ostrea (Turkostrea) turkestanensis baisunensis, Vialov; Pl. XXIII, figs. 1, 4

1970 Ostrea (Turkostrea) turkestanensis baisunensis, Berizzi Quarto di Palo, Pl. 36, figs. 4-7

Material: twelve left valves, MASSH 0014 - MASSH 0026.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

Description.— Relatively thin to thick shell of small to medium size, moderately inequivalve and inequilateral, nearly moderately to strongly convex, subtriangular to suboval in outline. Right valve concave marked by concentric growth lines. Sculptured somewhat irregular, strong and large radial ribs on the left valve (26 numbers) with rounded tops and bifurcated with equal and narrow interspace that are wider toward the ventral margin. The internal view is not visible in all studied specimens because of sediment casing. Umbo inflated, attachment area is clearly visible, mostly extended as a relatively large area of the umbral region. Triangular ligament area (turkostreoid type) slightly posteriorly curved, mostly longer than high. Catachomata can be seen and shallow-short vermiculate-chomata are present. Muscle scar is unknown due to coating by sediment. Vesicular layers not observed. Umbo rather big and non-coiled. Thin concentric fairly compressed growth lamellae slightly visible. Riels particularly in the umbral region become denser toward anterior and posterior sides. Anterior margin weakly convex, becoming a broad arc at the ventral margin. Posterior margin slightly hollowed beneath part of the umbo.

Dimensions.— Figure 12.5, L: 62 mm, H: 66 mm and maximum shell thickness of left valve 31 mm. Figure 12.6, L: 43 mm, H: 52 mm and maximum shell thickness of left valve 22 mm. Figure 12.7 (broken form), L: 34 mm, H: 47 mm and maximum shell thickness of left valve 21 mm.

Taxonomic comments.— The specimens are larger than Ostrea afghanica and Ostrea (Turkostrea) turkestanensis turkestanensis. Ostrea turkestanensis baisunensis has more rounded and rough ribs in comparison with Ostrea afghanica and Ostrea turkestanensis turkestanensis. Convexity of left valve is generalized unlike Ostrea afghanica and Ostrea (Turkostrea) turkestanensis turkestanensis and mostly focused on the central part of the left valve. Also, a broader umbo can be seen in this subspecies. In many sections in Central Asia, the subspecies was reported from the Middle Eocene (Berizzi Quarto di Palo, 1970), but in this study we report the subspecies from the Early Eocene.

Ostrea (Turkostrea) afghanica Vialov 1938

Figures 5.6 and 5.7

1948 Ostrea (Turkostrea) afghanica, Vialov, Pl. XXV, figs. 1-3

1970 Ostrea (Turkostrea) afghanica, Berizzi Quarto di Palo, Pl. 33, figs. 3-4

Material: eight left valves, MASSH 0006 - MASSH 0013.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

Description.— Thin shell of small size, moderately to strongly convex, rhomboid to quadrangular with symmetric outline. Sculptured left valve with blunt, rounded and high-bifurcated radial ribs with equal interspace in dorsal view and wider toward the valve margin, rather deep and narrow; 29 radial ribs, the secondary ribs positioned on tops of primary ribs. Some specimens have a developed small to large attachment area. Triangular ligament area (turkostreoid type) posteriorly curved, mostly longer than high with narrow and diagonal bourelets. Catachomata and shallow-short vermiculate-chomata are present. Muscle scar unknown because of coating by sediment. Attachment area small to large. Vesicular layers not seen. Umbo small, relatively acute, inflated and slightly orthogyrate.
Figure 6 1: Cubitostrea sp., #MASSH 0029; 2: Cubitostrea plicata, #MASSH 0027; 3: Argopecten sp., #MASSH 0036; 4: Cardium sp., #MASSH 0038; 5a-5c: Pterolucina sp., #MASSH 0044; 6a-6b, 7a-7b: Pelecyora (Cordiopsis) subathooensis, #MASSH 0047-0048; 8a-8b: Meretrix sp., #MASSH 0052; all specimens from Chehelkaman Formation, Sheikh section.

Dimensions.— Figure 11.6, L: 28 mm, H: 29 mm and maximum shell thickness of left valve 11 mm. Figure 11.7, L: 21 mm, H: 26 mm and maximum shell thickness of left valve 7 mm.

Taxonomic comments.— Vialov (1936) and Stenzel (1971) considered Ostrea (Turkostrea) afghanica as a subgenus characterized by their oval to triangular-ovoid, strongly convex left valves. Gryphaea is homeomorphous with strong chomata, continuous radial ribs with rounded crests, and slight to strong posteriorly curved ligament area (turkostreoid type). The difference between Turkostrea and Sokolovia is the less strongly gryphaeate shape of Turkostrea (Berizzi Quarto di Palo, 1970; Griffin et al., 2005). The specimens of this subgenus in the present study are smaller, with a shell more convex, and a sharp and pointed umbo.

Subfamily Crassostreinae Torigoe, 1981
Genus Cubitostrea Sacco, 1897

Cubitostrea plicata Solander 1766

Figure 6.2
1958 Ostrea (Cubitostrea) plicata, Kljušnikov, Pl. 24, figs. 2-3
1964 Ostrea (Cubitostrea) plicata aralensis, Karajuleva, Pl. 10, fig. 1; Pl. 12, figs. 1a-1b
1972 Ostrea cf. plicata, Kecskemétiné-Körmendy, Pl. 30, figs. 3-5

Material: two left valves, Massh 0027 - Massh 0028.
Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

Description.— Small size, short, high valves ornamented with many thick ribs, curved and crescent in outline; thick left valve slightly to moderately convex, increase in height of radial ribs on left valve and progressively stronger curvature of
shell. About 16 to 21 rough and coarse ribs that are long and radial in the irregular lines, rising of two and rarely more secondary costae that comprehensively similar to primaries can be seen. Interspace of ribs is deep and narrower than ribs, shell wider than high. The inner part of the specimens is covered by sediment, so it is impossible to check the ligament area, chomata and muscle scar. Spirogyrate umbo is slightly twisted.

**Dimensions.**— L: 34 mm, H: 31 mm and maximum shell thickness of left valve 7 mm.

**Taxonomic comments.**— Almost all specimens of *Cubitostrea* are small to medium in size and have a curved outline (Stenzel, 1971), but all of our specimens are small and show a curved and crescent shape. This genus can be separated from *Sokolovia* and *Turkostrea* by sharply triangular to sickle-shaped shell (Griffin *et al.*, 2005). *Cubitostrea* contains strong radial ribs.

**Cubitostrea sp.**

Figure 6.1

Material: two left valves, MASSH 0029 - MASSH 0030.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

**Description.**— Small size, short, vertically ovate, slightly convex, sculptured with irregular rounded-thick ribs almost toward ventral side, slightly crest arched toward the margins and concentric lamellae inconspicuous. The height and rounding of the radial ribs increase on the left valve, almost 10 to 12 long and radial ribs that are in irregular lines. Rising of two or more secondary costae toward ventral side can be seen. Interspace of ribs is deep and narrower than ribs. The width of the shell is greater than the height. Inner part of specimens was coated by sediment, impossible to see the ligament area, chomata and muscle scar. An opisthogyrate umbo can be seen in this material.

**Dimensions.**— L: 19 mm, H: 28 mm and maximum shell thickness of left valve 4 mm.

**Taxonomic comments.**— Small sized *Cubitostrea* with strong radial ribs that are not curved and crescent shaped as in *Cubitostrea plicata*. The opisthogyrate form of umbo in this species differs from *Cubitostrea plicata*.

Order Pterioidea Gray, 1847

Family Pectinidae Rafinesque, 1815

Genus *Argopecten* Monterosato, 1889

**Argopecten sp. A**

Figure 6.3

Material: three left valves, MASSH 0035 - MASSH 0037.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

**Description.**— Small convex shell, subcircular, height is slightly larger than its length. Beak rounded, protruding over the dorsal margin, auricles with similar extension and anterior auricle is broken. The posterior auricle has a straight margin and 5 to 6 fine riblets on the surface. Byssal sinus poorly preserved, but apparently shallow and wide. Commarginal lamellae poorly preserved, barely exhibited as straight outline on ribs and interspaces. Radial ribs narrow with rounded tops, wider toward the ventral margin, with a maximum of 38 ribs. The ribs of this species in cross-section are square even the interspaces are more or less square.

**Dimensions.**— L: 38 mm, H: 37 mm.

**Taxonomic comments.**— The shell convexity, orbicular outline, oblique growth, the symmetric size of the auricles, absence of interspaces riblets and the sculpture characteristics are the main features that distinguish *Argopecten* from the close related genus *Aequipecten* and other members of the subfamilies Pedinae and Pectininae (Stenzel, 1971; Hertlein and Jordan, 1927; Waller, 1969, 2011; Bouchet, 2012).

Differences in the numbers of ribs (38) in comparison with the other species such as *Argopecten levicostatus* are the most specific property of this species. The normal rib number in different species of *Argopecten* ranges from 17 to 25 (Moore, 1984; Morales-Ortega *et al.*, 2016).
Stratigraphic distribution of *Argopecten* is Oligocene to recent (Moore, 1984), except the only report from the Eocene of Bateque Formation, Baja California Sur, Mexico (Morales-Ortega et al., 2016). Our new finding from the Kopet-Dagh Basin is the second report of this genus from the Early Eocene and the first occurrence from the Tethyan realm in the Asiatic provinces; it extends its geographic distribution eastward.

Subclass Heterodonta Neumayr, 1884
Superfamily Carditoidea Fleming, 1828
Family Cardiidae Lamarck, 1809
Genus *Cardium* Linnaeus, 1758

**Cardium sp.**

Figure 6.4

Material: four cast of right valves, MASSH 0038 - MASSH 0041.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base, horizon 2 at 9 m from the base.

Description.— Small size, oval-sub-triangular outline and rather inflated internal casts. Shell height is slightly greater than its length. Umbo very prominent, incurved, rounded and orthogyrated. The sculpture consists of at least 20 rounded ribs, denser in umbonal region and the interspaces widely spaced toward ventral side. Surface slightly arched between the umbos and ventral margin.

Dimensions.— L: 34 mm, H: 37 mm.

Taxonomic comments.— Although specimens are imperfect, they show some features such as outline and radial ribs, which are separated by a wide and conspicuous channeled interspace. Ribs on the middle part are flatter with interspace, wider and more channeled toward ventral side. Less compact ribs and incurved umbo are index features for *Cardium halaense* and *Cardium kanleanum* that have already been reported from Eocene sedimentary succession of Afghanistan (Berizzi Quarto di Palo, 1970).

Superfamily Cardioidea Lamarck, 1809
Superfamily Lucinacea Fleming, 1828
Family Lucinidae Fleming, 1828
Genus *Pterolucina* Chavan, 1942

**Pterolucina sp.**

Figures 6.5a – 6.5c

Material: four left valves, MASSH 0042 - MASSH 0045.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base, horizon 2 at 1.1 m from the base, horizon 3 at 9 m from the base.

Description.— Internal bivalve cast, small in size, shell equivalve, equilateral and slightly oblique toward posterior, slightly spherical and circular in outline shape, height slightly more than width, regularly biconvex and compressed, very hollowed beneath the umbo. A type of marginal keel can be seen running from the dorsal side to ventral. Central part of the valves more convex in comparison with laterals. Ornamentation not well preserved but in some parts concentric convex can be conspicuous. Umbos slightly prominent.

Dimensions.— L: 21 mm, H: 26 mm.

Taxonomic comments.— Maximum convexity of this species is in the mid- umbonal region and slowly decreasing toward the posterior and anterior sides. Umbo slightly recurved forwards. As a result of the poor preservation in our specimens, ornamentations are not visible. Cosel (2006) discussed that the new genus *Joellina* (Cosel, 2006) most closely resembles *Pterolucina*, but *Pterolucina* has more separated areas and a stronger hinge dentition.

Superfamilia Veneroidea Rafinesque, 1815
Family Veneridae Rafinesque, 1815
Genus *Pelecyora* Dall, 1902
Subgenus *Cordiopsis* Cossmann, 1910

**Pelecyora (Cordiopsis) subathooensis**

d’Archaic and Haime, 1854

Figures 6.6a, 6.6b, 6.7a and 6.7b

1975 *Cordiopsis subathooensis*, Mathur, figs. 4A-F
2000 *Cordiopsis subathoensis*, Mathur and Juyal, Pl. 10, figs. 5-6

Material: six internal molds, MASSH 0046 - MASSH 0051.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base, horizon 2 at 9 m from the base.

*Description.* — Convex specimens with an oval outline and anteroposteriorly elongate. Umbo prominent and prosogyrate, lunule faintly marked. Commarginal growth lines weakly impressed in the surface. A narrow keel running along the margin indicates a partial opening of the valves.

*Dimensions.* — MASSH 0054; L: 49 mm, H: 42 mm, thickness: 25 mm. MASSH 0055; L: 43 mm, H: 38 mm, thickness 19 mm. Range: L: 27-49 mm, H: 21-42 mm, thickness: 9-25 mm. Average: L: 35 mm, H: 29 mm, thickness: 15 mm.

*Taxonomic comments.* — The species shows a considerable variation in shape from ovate to bluntly trigonal, subtrapezoidal to subtrigonal (Mathur and Juyal, 2000). *Pelecyora* (*Cordiopsis*) *subathoensis* has a subcentral umbo that shifted anteriorly and somewhat rolled inward. Lunule not impressed but large and cordiform. The surface possesses concentric striations. The species was only recorded for the Ypresian rocks of the Indus Basin (Mathur and Juyal, 2000).

Genus *Meretrix* Lamarck, 1799

*Meretrix sp.*

Material: five internal molds, MASSH 0052 - MASSH 0056.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base, horizon 2 at 9 m from the base.

*Description.* — Small size, convex and ovate, elongated anteroposteriorly. Shell equivalve but right valve slightly smaller than the left one. Umbo prominent, rounded and prosogyrate. Umbonal apex of right and left valves not in contact. Concentric growth line faintly impressed.

*Dimensions.* — L: 29 mm, H: 23 mm, convexity of left valve: 7 mm.

*Taxonomic comments.* — The form of the shell, convexity and features of the umbo best fits with the genus *Meretrix*.

Order *Myoida* Goldfuss, 1820

Superfamily *Hiatelloidea* Gray, 1824

Family *Hiatellidae* Gray, 1824

Genus *Panopea* Ménard de la Groye, 1807

*Panopea sp.*

Material: two external molds, MASSH 0057 - MASSH 0059.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 1.1 m from the base.

*Description.* — Shell is elongate, rather thick, large, left valve moderately convex, right valve slightly convex. Beak small, acute and wholly recurved, umbonal part is triangular. External surface is covered with fine concentric ornament. Antero-posteriorly elongated. Width of the shell greater than its height.

*Dimensions.* — L: 39 mm, H: 27 mm.

*Taxonomic comments.* — Okan and Hoşgör (2009) discussed in the Tethys provinces that most of Early Paleogene deep burrowing bivalves have been assigned to different *Panopea* species. The main distinguishing characters of this genus is its elongated shape and its rather short anterior margin (Okan and Hoşgör, 2009).

Subclass *Anomalodesmata* Dall, 1889

Order *Pholadomyoida* Newell, 1965

Family *Pholadomyidae* Gray, 1847

Genus *Pholadomya* Sowerby, 1823

*Pholadomya sp.*

Material: one articulated specimen of internal mold, MASSH 0060.
Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base.

Description.—The external shell is missing however the internal mold exhibits sufficient morphological features to identify at generic level. Large dimensions, inequilateral, ovate globose, oblong in outline. Umbo prominent and ornamentations consist of many thin ribs. Anterior portion reduced and posterior portion elongated. Strongly inflated.

Dimensions.—L: 55 mm, H: 47 mm, inflation: 37 mm.

Taxonomic comments.—Prominent features for identification of this specimen are the outline, inequilateral, strongly inflated and reduced anterior portion that is in contrast to the elongated posterior portion (Harzhauser and Mandic, 2001).

Class Gastropoda Cuvier, 1795
Superfamily Cerithioidea Ferrusac, 1819
Family Turritellidae Lovén, 1847
Genus Turritella Lamarck, 1799
Turritella subathooensis d’Archaic and Haime, 1854

2000 Turritella subathooensis, Mathur and Juyal, Pl. 11, figs. 4-12

Material: thirty one specimens, MASSH 0061 - MASSH 0091.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base, horizon 2 at 1.1 m from the base, horizon 3 at 9 m from the base. Ghaleh-Zou section, horizon 1 at 4.3 m from the base, horizon 2 at 4.59 m from the base.
Description. — Medium to large, turriculate to conical, and high-spired gastropods. Spire elongated, consisting of about 4 non-overlapping whors (incomplete specimen), separated by slightly deep sutures. Whorl surfaces gently convex, each whorl ornamented with about 15–17 tuberculate thin spiral cords. Protoconch and fist whors not preserved. Diameter of whors increasing gradually from apex to aperture.

Dimensions. — H: 35 mm (incomplete), W: 15 mm.

Taxonomic comments. — This species is only recorded in the Ypresian sediments of Indus Basin, characterized by turreted, gently convex whors, rather loosely coiled and body whorl ornamented by concentric oblique growth lines (Mathur and Juyal, 2000). Preservation of the specimens described here is better than the Indus region forms.

Genus Haustator Montfort, 1810

Haustator gilberti Bowles 1939

1939 Turritella gilberti, Bowles, Pl. 32, fig. 16
1996 Turritella gilberti, Allmon, Pl. 7, figs. 4-5
1969 Turritella gilberti, Toulmin, Pl. 2, fig. 10
1980 Turritella gilberti, Dockery, Pl. 2, fig. 2

Material: four specimens, MASSH 0092 - MASSH 0095.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base. Ghaleh-Zou section, horizon 1 at 4.3 m from the base, horizon 2 at 4.59 m from the base.

Description. — Medium to large, turriculate to conical and high-spired shell. Spire elongated consists of about 2 to 4 non-overlapping tube like whors (incomplete specimens), with loose coiling separated by strongly and spaced deep sutures. Whorl surfaces strongly convex, Protoconch and first whors not preserved. Diameter of whors increasing gradually from apex to aperture.

Dimensions. — H: 41 mm (incomplete), W: 15 mm.

Taxonomic comments. — Strongly and spaced deep sutures, loose whors and absence of carina separated this specimen with Turritella. This species is similar to the specimens from the Kutch district of India reported for first time by Halder and Sinha (2014).

Family Siliquariidae Anton, 1838

Genus Tenagodus Guettard, 1770

Tenagodus sowerbyi Halder and Sinha, 2014

Figures 7.5 and 7.6

2014 Tenagodus sowerbyi, Halder and Sinha, figs. 3s-3u

Material: three specimens, MASSH 0096 - MASSH 0098.

Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base.

Description. — Medium to large, turriculate to conical and high-spired shell. Spire elongated consists of about 4 non-overlapping whors (incomplete specimen), separated by slightly deep sutures. Whorl surfaces gently convex, each whorl ornamented with about 15–17 tuberculate thin spiral cords. Diameter of whors increasing gradually from apex to aperture.

Dimensions. — H: 53 mm (incomplete), W: 15 mm.

Taxonomic comments. — Shape of adult whorl and occurrence of basal carinate in Haustator is a prominent feature in the Paleocene and lower Eocene turritellid assemblages (Allmon, 1996; Gale et al., 1999).

Family Siliquariidae Anton, 1838

Genus Globularia Swainson, 1840

Globularia sp.

Figures 7.7 – 7.9

Material: twelve specimens, MASSH 0099 - MASSH 0110.
Stratigraphic occurrence in the Kopet-Dagh Basin: Chehelkaman Formation, Sheikh section, horizon 1 at 0.3 m from the base, horizon 2 at 1.1 m from the base, horizon 3 at 10.6 m from the base.

**Description.**— Shell globose, small to large, low-spired with overlapping whorls that are slightly convex and separated with deep sutures; 4 to 6 whorls that rapidly decrease toward the apex. Greatly inflated body whorl and large aperture.

**Dimensions.**— H (range): 25 - 35 mm, W (range): 17 - 23 mm.

**Taxonomic comments.**— *Globularia* has a large last whorl and low spire. A lot of species for this genus have been reported from Paleocene and Eocene deposits of Tethys realm. Their stratigraphic occurrence in the Kopet-Dagh Basin is similar to the same occurrence in Turkey (see: Okan and Högör, 2008).

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6. **Faunal features of the shell beds**

A *sensu lato* shell bed is considered here as any horizontal and continuous bed with high concentrations of mollusk fossils including shells, molds or casts (Dorsey and Kidwell, 1999). By this definition, four shell beds are recognized in the Sheikh section, and only one in the Ghaleh-Zou section. They are described as follow and named based on the most abundant macrofossil contents.

6.1 **SHEIKH SECTION**

6.1.1 **TURRITELLIDAE SHELL BED**

This shell bed is located at +0.3 m from the base of the section (Figure 2), with 0.3 to 0.4 m thick. The bed consists of high-spired shells in a silty, moderately sorted, fine-grained and poorly cemented sandstone matrix. In this bed, fossils belong to two classes of mollusks. Bivalves are represented by shells of *Pycnodonte* sp., *Pelecyora* (*Cordiopsis*) *subtahooensis*, *Cardium* sp., and *Pholadomya* sp. representing about twenty-five percent frequency of collected fossils from this bed. Gastropods are signified by shells of *Globularia* sp. and Turritellidae with the latter the most abundant mollusk taxa (Table 1) in the beds, which includes the species *Turritella subathooensis*, *Hausator gilberti*, and *Tenuodus sowerbyi* in a very high relative percentage.

Rare gastropods (almost 5%) and some bivalves (nearly 30%) preserve the shell. The remainder fossils consist of sandy cemented shells of both internal and external cast. Based on some biostratigraphic data, such as degree of disarticulation and corrosion, fragmentation of fragile and robust fossils and bioerosion of skeletal elements (shown in Table 1, Relative abundance of macrofossil groups from the Sheikh section in the Turritelline horizon.)

<table>
<thead>
<tr>
<th>Groups</th>
<th>Percentages</th>
</tr>
</thead>
<tbody>
<tr>
<td>Turritelline gastropod fauna</td>
<td>68.86</td>
</tr>
<tr>
<td>Non Turritelline gastropod fauna</td>
<td>5.13</td>
</tr>
<tr>
<td>Oyster fauna</td>
<td>15.03</td>
</tr>
<tr>
<td>Other bivalves</td>
<td>10.98</td>
</tr>
</tbody>
</table>

Table 1. Relative abundance of macrofossil groups from the Sheikh section in the Turritelline horizon.
the Table 2), perspective of fossil preservation was provided for the studied sections. More than half of the bivalve and gastropod shells and their casts represent fragmentation and almost half of them show fracturing. However, almost all of them are disarticulated. The external surface of the shell is moderately to strongly abraded, but there are no signs of encrustation, bioerosion, and predation.

The shell size of turritellids ranges between 2 and 6 cm, which represent the different stages of their growth. The presence of different shell sizes indicates a weak sorting. The gastropods have a bimodal orientation (Figure 8) as turritelline shells are commonly oriented along their axis, which indicates effects of oscillatory waves (e.g., Seilacher, 1959, 1973; Fürsich and Oschmann, 1993) rather than unidirectional currents.

6.1.2 PYCNODONTE – TURKOSTREA SHELL BED

The lower basement of this shell bed is located at +1.1 m from the base of the section. The bed has a thickness of 2 m and the interval between two consecutive shell beds is unfossiliferous. The matrix is composed of fine-grained muddy sandstone. Two subtypes are recognized, including moderately sorted with edgewise-oriented bioclasts (Figure 6b). Toward the top bed, the matrix changes to a well sorted with random to convex up oriented bioclasts (Figures 9b and 9c). The bivalves consist of *Pycnodonte* sp., *Ostrea* (*Turkostrea*) *afghanica*, *Ostrea* (*Turkostrea*) *turkestanensis*, *Argopecten* sp., *Cardium* sp., and *Pholadomya* sp. The gastropods contain a small amount of turritelline and *Globularia* sp. The most abundant group is the oysters, showing a relative 80% frequency. Toward the upper part of the bed, fossil content strongly increases. The best-preserved fossils are oysters and *Argopecten* sp., while *Cardium* sp., and *Pholadomya* sp. consist of internal or external molds.

All shells are moderately fragmented (Figures 9 and 10), and shell fragments were observed floating in the calcareous matrix toward the bed top, with a sparse packing (Figure 9). Almost all oyster shells are disarticulated, predominating left valves

<table>
<thead>
<tr>
<th>Skeletal type</th>
<th>Intact transport</th>
<th>Fragmentation</th>
<th>Disarticulation</th>
<th>Corrosion</th>
<th>Encrustation</th>
<th>Boring predation</th>
</tr>
</thead>
<tbody>
<tr>
<td>Gastropod fauna</td>
<td>+</td>
<td>+</td>
<td>-</td>
<td>+</td>
<td>+</td>
<td>Absent</td>
</tr>
<tr>
<td>Oyster fauna</td>
<td>+</td>
<td>+</td>
<td>±</td>
<td>+</td>
<td>+</td>
<td>+</td>
</tr>
<tr>
<td>Other bivalves</td>
<td>+</td>
<td>+</td>
<td>±</td>
<td>+</td>
<td>Absent</td>
<td>Absent</td>
</tr>
</tbody>
</table>
(90%). Their elements almost situated in inverse life position (convex-upward) and their left valves are strongly encrusted and bored (Figures 9c, 9d and 10). Attachment areas of the most presence oysters are small.

6.1.3 PELECYORA (CORDIOPSIS) SUBATHOOENSIS – CARDIUM SHELL BED

This shell bed with the thickness of 0.4-meter thickness is located at +9 m from the base of the section and the barren distance between the second and third shell bed. The matrix is composed of calcareous siltstone. Collected taxa are represented by the bivalves Pycnodonte sp., Cardium sp., Pelecyora (Cordiopsis) subathooensis, and the gastropod Turritella subathooensis. The dominant fossils belong to Pelecyora (Cordiopsis) subathooensis and Cardium sp. that is relatively high in abundance. Most bivalves are articulated. The remaining specimens are casts derived from articulated valves. The only recognized taphonomic feature is a moderate fragmentation in the shells and casts. The specimens do not exhibit a particular trend in arrangement or orientation. All shells show a random convex up or down arrangement of the shells. Except by the Pelecyora (Cordiopsis) subathooensis, being all in situ life position (Figures 11a - 11d). The fossil density in the shell bed is especially higher in Cordiopsis - Cardium horizons, which represent a pavement structure in the Cordiopsis - Cardium shell bed (Figure 11).

6.1.4 GLOBULARIA SP. SHELL BED

The shell bed with 0.5-meter thickness is located at +10.6 m distance from the base of the section, with a matrix composed of silty marl. Most of fossils are internal molds of gastropods. The
fossils of Ampullinidae shells, Globularia sp., are the only element with moderate frequency with preserved shell and remnants of internal casts. The shells are rarely broken. The only recognized taphonomic event is moderate encrustation marks on the shells and casts aperture area with no signs of fragmentation.

6.2 GHALEH-ZOU SECTION

6.2.1 TURRITELLIDAE SHELL BEDS

In the Ghaleh-Zou section, there is a couple of shell beds in the distance of 4.3 and 4.6 m from the base of the section (Figure 3a), separated from each other by a cross-stratified sandstone. The thickness of the lower bed is 0.15 m and the upper is 0.12 m. The matrix in both beds is composed of white loose sand in non-weathered surfaces and buff in weathered surfaces. The only macrofossil contents of both shell beds are turrellids, such as Turritella sp. and Haustator sp. They do not have preserved shells, but rather consist of almost entirely fragmented, incomplete specimens with a random and irregular arrangement (Figures 12b - 12c).

7. Age dating

The assemblage of Pelecyora (Cordiopsis) subathooensis and Turritella subathooensis in two shell beds of the Sheikh section, and the Turritellidae shell bed and Pelecyora (Cordiopsis) subathooensis – Cardium shell bed, allows us to assign them to the Cordiopsis subathooensis - Turritella subathooensis Zone, established by Mathur and Juyal (2000) from the Subathu area of India. This biozone is equivalent to the middle parts of the planktonic foraminifera P8 Zone (Berggren et al., 1995). Therefore the association at the Sheikh section indicates an early Eocene age. Sahni and Kumar (1983) in the Lesser Himalaya, Northwestern India, reported several bivalve and gastropod genera such as Venericardia, Trachycardium, Seila, Venus, Pinna, and some species of turritellids as associate fauna with the Pelecyora (Cordiopsis) subathooensis - Turritella subathooensis Zone. In the Sheikh section, associate macrofossils in this zone are oysters and Cardium sp.

To verify of an early Eocene age for the Chehelkaman Formation at the Sheikh section, 10 samples were prepared for nannofossils investigations.
These samples were taken from thin muddy lamina and calcareous parts of the formation. Based on the extracted nannofossils (Figure 14) the biozones NP12 and lower part of NP13 were recognized in this section, which concur with the ages of the macrofossil assemblages. The documented nannofossils zones are as follow:

### 7.1 TRIBRACHIATUS ORHOSTYLUS ZONE (NP12)

According to Perch-Nielson (1985) this biozone comprises the interval from the first occurrence of *Discoaster lodoensis* to the last occurrence of *Tribrachiatus orthostylus* or the first occurrence of *Toveius crassus*. This biozone, which is equivalent to CP10 of Okada & Bukry (1980), corresponds to the lower 8 m rocks of the Chehelkaman section. The most important associate nannofossils of this biozone are *Chiasmolithus consuetus, Chiphragmalithus armatus, Coccolithus pelagicus, Discoaster barbadiensis, Discoaster diastypus, Discoaster lodoensis, Discoaster kuepperi, Ellipsolithus macellus, Ericsonia formosa, Helicosphaera lophota, Markaliu simversus, Neococcolithes dubius, Neococcolithes protonus, Reticulofenestra dictyoda, Sphenolithus conspicus, Sphenolithus editus, Sphenolithus moriformis, Sphenolithus radians, Tribrachiatus orthostylus, and Zygophyllolithus bijugatus.

### 7.2 DISCOASTER LODOENSIS CONCURRENT RANGE ZONE (NP13)

The lower boundary, defined above, coincides with the upper boundary markers of NP12. Its upper boundary characterized by the first occurrence of *Discoaster sublodoensis* (Perch-Nielson, 1985), is equivalent to CP11 of Okada & Bukry (1980). This interval zone includes the succession from +8 m from the base of the Chehelkaman Formation into 2 m within the Khangiran Formation. Due to the absence of *Discoaster sublodoensis* in the studied sections last occurrence of *Toveius crassus* is used as the upper zonal marker (Perch-Nielson, 1985). Other nannofossil associations within this biozone are *Chiasmolithus consuetus, Coccolithus pelagicus, Discoaster deflandrei, Discoaster kuepperi, Discoaster lodoensis, Ericsonia formosa, Helicosphaera lophota, Reticulofenestra dictyoda, Sphenolithus moriformis, and Toveius crassus*.

### 8. Discussion

#### 8.1 PALEOECOLOGICAL CONDITIONS

Shell beds are skeletal concentrations taphocenosis, common features of shelf environments in the
past and present (Kidwell et al., 1986). They are important tools to retrieve fossil taxa, and have been extensively considered for paleoecological analysis (e.g., Muller, 1951; Schafer, 1972; Brett and Baird, 1986; Fürsich and Oschmann, 1993; Fürsich and Pandey, 1999). The shell beds work as environmental indicators and useful records for both ecological interpretations and recognition of the processes related to the formation of layers (Kidwell et al., 1986; Gordillo et al., 2014).

The turritellid shell bed at the SSA is characterized by a high abundance of elements, but with a low diversity, less than 10 genera or 10 species. This shell bed matches well with the Turritelline Dominated Assemblage (TDA) described by Allmon and Knight (1993); however they reported them from the Cretaceous succession. TDAs are dense macrofossil assemblages, dominated by turritellids in a low total diversity (<20 species). The TDAs are associated to sandstone and limestone rocks, being a widespread geological phenomenon from Cretaceous through Cenozoic (Allmon and Knight, 1993). According to Allmon (1988), communities with high abundance of turritellids required a nutrient enriched environment. Some other studies on the turritellid populations have revealed a correlation among high levels of nutrients, high productivity, and high abundance of these gastropods (Allmon et al., 1992, 1994, 2011). Hence, an important factor in the formation of TDAs is a high biological productivity due to nutrient accretion, which is associated with runoff, volcanism or upwelling (Allmon, 1988, 2007).

The TDA of the Sheikh section therefore may indicate a nutrient enrich environment in the basin during an early moment just after the Cenozoic transgression (Figure 2). Volcanic activities have never been reported from the Kopet-Dagh Basin (Afshar-Harb, 1979) and coastal upwellings are ignored due to the absence of evidences, but a significant input of nutrients from continent might have enriched the marine basin. During the Early Eocene, the high levels of carbon dioxide in the atmosphere increased the weathering of rocks (Leckie, 1989; Ravizza et al., 2001); and hot and humid conditions, accelerated the global hydrologic cycle (Schroeder, 1992; Berner, 1994; Dickens et al., 1995, 1997; Ravizza et al., 2001). Both conditions favored the entrance of continental runoff into the basin (Leckie, 1989; Ravizza et al., 2001).

The enriched basin by runoff is also suggested by the Pycnodonte – Turrostre shell beds. The habitat, life position and feeding mechanisms of both oyster groups (gryphaeid and ostreids) found in the basins connected to a well supply of sediments and
DISCUSSION

Systematic paleontology and taphonomic studies of Ypresian mollusks at the Kopet-Dagh Basin, NE Iran

nutrients. Most of the Cenozoic oysters live in tropical and subtropical shallow marine basins, being epifaunal organisms, living reclined or cemented, and being efficient suspension feeders (Stanley, 1970; Stenzel, 1971; Seilacher, 1984). The high boring and encrusting evidences on the fossils from the *Pycnodonte* – *Turkostrea* shell beds display the nutrient enriched environment. High levels of nutrients favored the growth of different mollusk populations and consequently increased the bioerosion of shells yielded by bivalve (*e.g.*, encrusting) and gastropods (*e.g.*, drilling) (Wood, 1993; Lescinsky *et al.*, 2002).

Recent oysters occupy the intertidal zone to a water depth of 30 m. Occurrence of *Turkostrea* shell beds belong to intertidal zone. *Pycnodonte* spp. inhabit at the subtidal zone to the depth of 30 m. The bivalves *Cardium* sp., *Meretrix* sp., *Pholadomya* sp., and *Panopea* sp. occupy both inter and subtidal environments (Lan, 1997). Furthermore, the stout gastropod *Turbinella* sp. were reported from the depth range of 30 to 50 m (Kachhara *et al.*, 2011), and turritellids are known from less than 2 to more than 150 m depth (Gekker *et al.*, 1962; Stanley, 1970, 1971; Houbrick, 1992; Allmon *et al.*, 1995; Lan, 1997; Kachhara *et al.*, 2011).

Since turriteliidias represent communities needing nutrients and they are the only predominant fossils in the Ghaleh-Zou section therefore we suggest a very shallow environment for this section. It is possibly related to the areas connected to the entrance of runoff from the land surface of terrestrial regimes. Lithofacies features such as loose white sandstones and fine gypsum beds confirm the depth range of this section was lesser than the Sheikh section.

### 8.2 DEPOSITIONAL CONDITIONS

#### 8.2.1 SHEIKH SECTION

The dominant supply of siliciclastic sediments retain low levels of CaCO₃, favoring the shell dissolution of several gastropods and bivalves, yielding cemented casts. Physical reworking and transportation of skeletal elements play an important role in the concentrations of the shells in the Chehelkaman Formation. This case indicates high energy levels in this phase. As well, highly encrusted and bored shells point to a long residence time on the seafloor that correspond to low rates of sedimentation. Abrasion degree, breakage and fragmentation of turritellids, their bimodal orientations and the lithofacies demonstrate the TDAs, in both Sheikh and Ghaleh-Zou sections, have deposited in the tidally dominated paleoenvironment. The same conditions were proposed for identical succession in the Caribbean province by Kirkland *et al.* (1996).

The bimodal orientations of turritellids along their axis (Figures 8 and 12a) indicate the effect of regular oscillatory flows (*e.g.*, Seilacher, 1959, 1973) in the Sheikh section (Figures 8 and 12a). In contrast, turbulent flows made irregular replacement of skeletal elements in the Ghaleh-Zou section, which has led to fractures and destroyed elements within the shell beds (Figures 12b - 12d).

These geometric and biostratonomic features represent concentrations under the influence of fair-weather waves, where abrasion and breakage should have been salient features (Fürsich and Oschmann, 1993).

The *Pycnodonte-Turkostrea* shell bed was deposited in the intertidal zone and laterally shows more turbulent conditions, around normal storm wave base. Disarticulated, shell valves - oriented convex-up valves indicate storm-flow conditions (Tomašových, 2006). Likewise a predominance of heavy and thick left valves clearly indicates sorting and influence of currents in storm flows. This high accumulation of the shells was the result of a high productivity combined with the low rates of sedimentation and occasional winnowing. The thinner right valves might have been selectively removed during winnowing, but they also could have been destroyed via bioerosion, as most shells are moderately encrusted or bored.

The degree of disarticulation, fragmentation and abrasion in bivalves suggests a low to intermediate sedimentation rate from 1 to 10 cm10⁻³y (Brett and Baird, 1986). Notwithstanding, the thick and heavy left valves of oysters, less prone
to transported concave down shells (Figures 9c - 9d, 10a, 10b and 10d), suggest rapid current flows (Kerans et al., 1994). Also, the absence of thin light right valves suggests a moderate energetic regime (Stenzel, 1971; Brett and Baird, 1986). Except for a few rare cases, almost all small attachment areas in oysters suggest they cemented in the mode of life only during the juvenile stage (Fürsich and Pandey, 1999).

There are not any traces of bioerosion and encrustation in the *Pelecyora (Cordiopsis) subathoensis* – *Cardium* shell bed and most of the shells are articulated. Tomašových (2006) have attributed the settling preference of such shell beds to a low energy environment, low rates of sedimentation and high frequency of dead bivalve shells. These conditions suggest a comparatively motionless environment after a storm, which produce a pavement arrangement of the shells (Figure 11).

Lithostratigraphically, the *Globularia* sp. shell bed is mostly marly and the rate of sandy-silty grains decreases to the topmost bed. Lack of breaking, fragmentation and a little evidence of encrustation even within the muddy matrix suggests deeper settling conditions below the influence of storm activity (Bressan and Palma, 2010).

8.2.2 GHALEH-ZOU SECTION

Laterally facies changes, from WNW to ESE in the SSA, have triggered differences in marine sequence from the Sheikh section into the Ghaleh-Zou section. Lithological similarity of the Pesteligh and Chehelkaman formations in the later section represent more consistent depositional conditions. The occurrences of two successive turritellid shell beds with intercalations of gypsum appear to directly indicate a shallow environment condition than the Sheikh section (Figure 10). Disturbed and irregular orientation of the suspension feeder turritelline gastropods indicates their development coincides with the energetic and turbulent stages of the rich nutrient environment (Figures 12b - 12c).

9. Conclusion

Clastic facies of Chehelkaman Formation in the northeast Iran at the Kopet-Dagh Basin consist of multiple molluscan assemblages in some shell beds. Paleontological investigations on these fos-
systematic paleontology and taphonomic studies of Ypresian mollusks at the Kopet-Dagh Basin, NE Iran

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