

The first crocodyliforms remains from La Parrita locality, Cerro del Pueblo Formation (Campanian), Coahuila, Mexico

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ABSTRACT

The record of land tetrapods of the Cerro del Pueblo Formation (Late Cretaceous, Campanian), in Coahuila, includes turtles, pterosaurs, dinosaurs, and crocodyliforms. This last group is represented only by goniopholidids, indeterminate eusuchians, and *Brachychampsia montana*. In this work we report the first crocodyliform remains from La Parrita locality, Cerro del Pueblo Formation, based on one isolated tooth, vertebrae, and osteoderms. The association of crocodyliforms, turtles, dinosaurs, and charophyte oogonia provide evidence for stagnant to fluvial environments on a delta plain with tropical climate for the Cerro del Pueblo Formation during the Late Cretaceous.

Keywords: **Crocodyliforms, Globidonta, Late Cretaceous, Coahuila, Mexico.**

RESUMEN

*El registro de tetrápodos terrestres en la Formación Cerro del Pueblo (Cretácico tardío, Campaniano) en Coahuila, incluye tortugas, pterosaurios, dinosaurios y crocodyliformes. Este último grupo está representado por goniofolididos, eusuquios indeterminados y *Brachychampsia montana*. En este trabajo se reportan los primeros restos de crocodyliformes de la localidad La Parrita, Formación Cerro del Pueblo, con base en un diente aislado, vértebras y osteodermos. La asociación de crocodyliformes, tortugas, dinosaurios y oogonias de cariofitas, proporcionan evidencia de ambientes fluviales en una llanura deltaica de tipo tropical para la Formación Cerro del Pueblo durante el Cretácico tardío.*

Palabras clave: **Crocodyliformes, Globidonta, Cretácico tardío, Coahuila, México.**

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1. Introduction

The fossil record of land vertebrates from the Cerro del Pueblo Formation in Coahuila includes turtles, squamates, crocodyliforms, pterosaurs, and dinosaurs. The turtle record includes a pleurodire with similarities to *Chedighaii*, the paracryptodires *Compsemys* and *Neurankylus*, the cheloniid *Mexichelys coahuilaensis*, the kinosternid *Hoplochelys*, the trionychid *Adocus*, and indeterminate chelydrid, kinosternoid, and trionychid remains (Brinkman, 2014). The only squamate is “*Coniophis*” sp., a taxon based on vertebrae collected near Rincon Colorado (Aguillon-Martinez, 2010). Pterosaurs are represented by one indeterminate ornithocheiroid (Frey and Stinnesbeck, 2014). The dinosaur record comprises the hadrosaurids *Velafrons coahuilensis* and *Latirhinus uistlani*, the ceratopsian *Coahuilaceratops magnacuerna*, ankylosaurids, the dromeosaurids *Trödon* and *Saurornitholestes*, the problematic coelurosaur *Richardoestesia*, tyrannosaurids, and ornithomimosaurids (Rivera-Sylva and Carpenter, 2014a, 2014b). Records of crocodyliforms in the Cerro del Pueblo Formation include indeterminate eusuchian and goniopholidid material (Rodríguez-de la Rosa and Cevallos-Ferriz, 1998), along with a skull table, two frontal bones, isolated teeth, and isolated osteoderms that Aguillon-Martinez (2010) referred to the globidontan alligatoroid *Brachychampsa montana*.

The material described in the present work comes from the La Parrita locality, located 54 km to the west of Saltillo (Figure 1), and includes fragments of a maxilla and lower jaw, vertebrae, osteoderms, and one bulbous tooth. The morphology of the tooth allows the classification as Globidonta, and thus is the second record of this clade of crocodyliforms in the Cerro del Pueblo Formation, Coahuila. The rest of the material can only be assigned as Eusuchia.

Globidonta is an alligatoroid clade characterized, at least ancestrally, by the presence of bulbous posterior teeth. Its content in recent phylogenetic analyses is redundant with that of Alligatoridae (e.g., Brochu, 2011; Skutchas *et al.*, 2014; Salas-Gis-

mondi *et al.*, 2015; Hastings *et al.*, 2016), but Globidonta is explicitly a stem-based group, and Alligatoridae is node-based (Brochu, 1999). Furthermore, some Late Cretaceous North American globidontans were excluded from Alligatoridae in earlier analyses (e.g., Brochu, 2004; Martin and Lauprasert, 2010), such as *Stangerochampsa mccabei* from the Early Maastrichtian Horseshoe Canyon Formation, southern Alberta (Wu *et al.*, 1996); *Albertochampsa langstoni* from the middle Campanian Dinosaur Park Formation of southern Alberta (Erickson, 1972), and *Brachychampsa* reported from several sites of early Campanian through late Maastrichtian age throughout western North America (Gilmore, 1911; Carpenter and Lindsay, 1980; Bryant, 1989; Norell *et al.*, 1994; Williamson, 1996; Sullivan and Lucas, 2003; Monroy Mújica, 2009; Aguillon-Martinez, 2010; Irmis *et al.*, 2013). Two species of *Brachychampsa* are recognized: *B. sealeyi* from the Campanian De-Na-Zin Member of the Kirtland Formation, San Juan Basin, New Mexico (Williamson, 1996); and the type species, *B. montana* from the Hell Creek Formation of Montana (Gilmore, 1911; Bryant, 1989; Norell *et al.*, 1994) and possibly elsewhere in western North America.

Institutional abbreviations: CPC, Colección Paleontológica de Coahuila (Paleontological Collection of Coahuila), Museo del Desierto (MUDE), Saltillo, Coahuila, Mexico.

2. Geological setting and paleoenvironment

The locality of La Parrita is part of the Cerro del Pueblo Formation, which is the oldest geologic formation from the Difunta Group (Figure 2A). The Cerro del Pueblo Formation is dated between 71 and 72.5 million years old (Obradovich, 1993; Eberth *et al.*, 2004), which indicates Campanian age. The Cerro del Pueblo Formation was deposited under marginal marine and shallow marine conditions, which were influenced by sea-level oscillations and storm events (Eberth *et al.*, 2004).

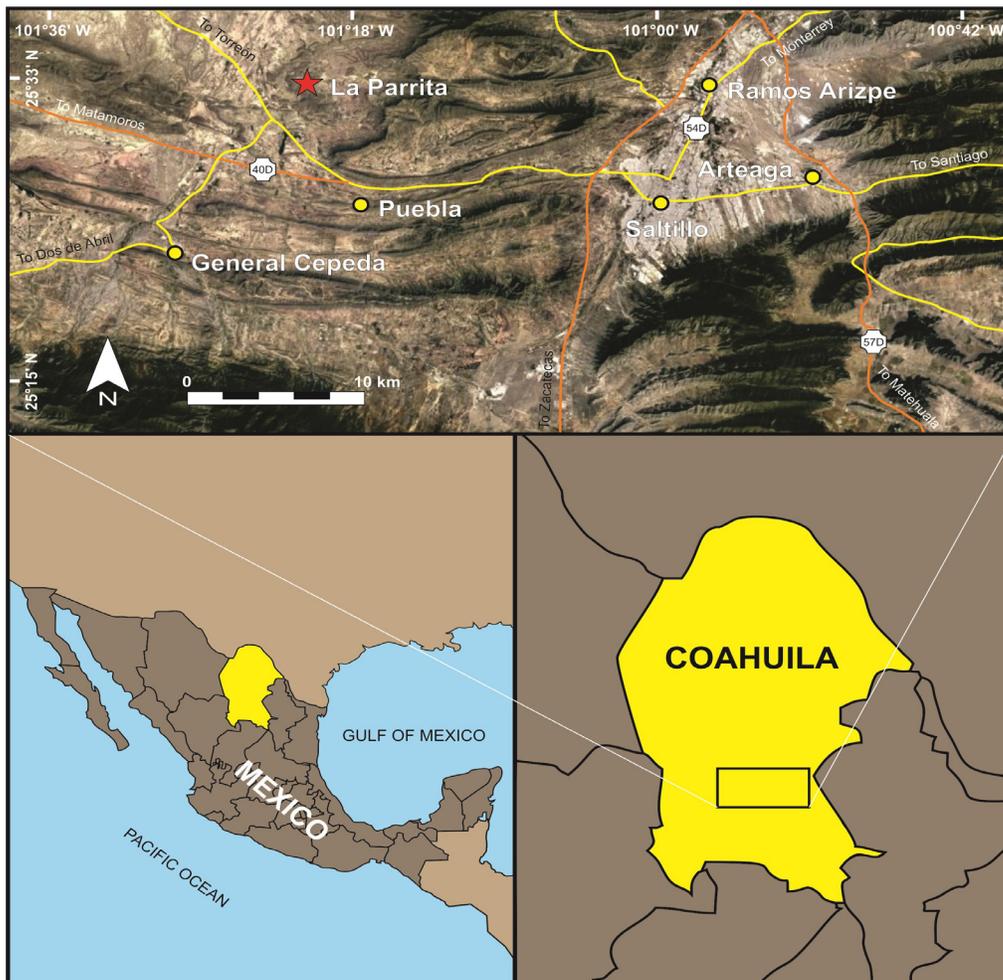


Figure 1 Location of the La Parrita locality (red star), Cerro del Pueblo Formation, Coahuila state.

The facies suggest deltaic systems with shallow flow channels, oxbow or residual lakes, bays and lagoons, coastal swamps, and wetlands with exuberant vegetation (McBride, 1974; Hill, 1988).

The stratigraphic sequence that outcrop in La Parrita locality is composed of alternating layers of sandstone, siltstone and shale (Figure 2B). At the base, it is comprised of 0.20 m of ochre-gray shales with greenish tonalities. Marine taxa such as the ammonite *Sphenodiscus* and some isolated, weathered dinosaur bones are associated with this layer. According to Eberth *et al.* (2004: 350-352, fig.5), *Sphenodiscus* is present in facies 1 corresponding to the Parras Shale, which means this first layer is the base in which Cerro del Pueblo Formation is supported (Vivas-González, 2013).

Above the base follows a 0.30 m layer consisting of exfoliated siltstones with a high concentration of oysters of diverse size, which indicate a brackish environment. According to Stinnesbeck and Frey (2014), the Cerro del Pueblo Formation sandstone and siltstone deposits are characteristic of brackish conditions, indicated by oyster banks and abundant non-ostrean bivalves and gastropods. Above the siltstone lies a 0.90 m hard, massive, coarse-grained sandstone layer with high concentrations of gastropods at the top, which indicates marine coastal environments (Vivas-González, 2013). A grit layer continues by 1 m above, with a 0.40 m thick concentration of gastropods lies above the grit. The next layer consists of 10 m of ochre-gray shale characterized

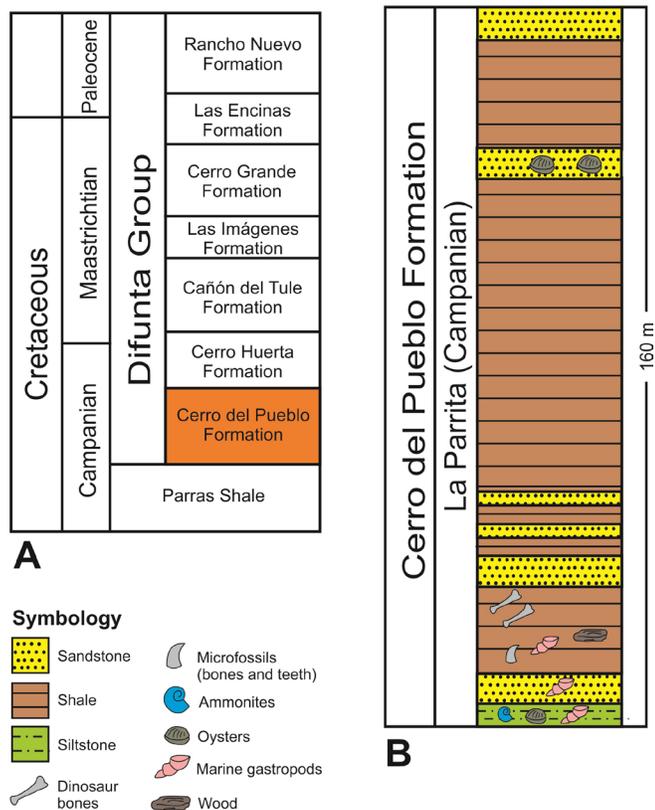


Figure 2 Stratigraphic chart showing the age and divisions of the Difunta Group with the stratigraphic position of the Cerro del Pueblo Formation (A), and local stratigraphy of La Parrita locality (B), showing the sequence of alternant strata (sandstone, siltstone, and shale). A large diversity of vertebrate and invertebrate fossils occurs in the second layer of shale (from the bottom).

by fossil diversity: the marine gastropods *Cerithium nodosa* and *Lissapiopsis* sp. are common, as well as oysters (*Flamingostrea* sp.). These mollusks represent the most common invertebrates in this shale layer (Vivas-González, 2013). Wood fragments are also present, and among vertebrate fossils, reptiles like trionychid turtles, some goniopholidids crocodyli-forms, bone and teeth of dinosaurs (hadrosaurids, tyrannosaurids and dromeosaurids) are present, as well as coprolites (Vivas-González, 2013). Eberth *et al.* (2004: 350–352) placed this kind of fossil assemblage between facies 10–12. The assemblage is suggestive of overflooded coastal plains, and the marine and limnic faunal assemblages that coexist in this layer indicate an estuarine environment (Vivas-González, 2013; Stinnesbeck and Frey, 2014).

The next layer consists of 0.40 m of massive, gray, fine-grained sandstones followed above by 0.50 m of dark-gray shale and again by 0.35 m of fine-

grained sandstones (Vivas-González, 2013). The following 0.60 m are composed of shale, then 0.20 m of fractured sandstones. Above the fractured sandstone layer lie 32 m of shale, which is devoid of fossils. Above this shale layer, there is a 3 m layer of massive medium-grained sandstones with high concentrations of bivalves, referred to *Inoceramus vamuxemi* in the top of the layer. No other fossils are reported in this layer. The next 18 m of the section are formed by shales, and the last 3 m on the top are medium-grained sandstones. No fossils were observed or reported in both of these top layers (Vivas-González, 2013).

This alternating sediment suggests the La Parrita locality underwent a series of events related to marine regressions and transgressions during deposition, which is consistent with other localities associated with the Cerro del Pueblo Formation (Eberth *et al.*, 2004).

3. Methods

All specimens were recovered using the standard vertebrate microfossil surface collection method at the La Parrita locality. All material was collected in an area of 15 m², in association with gastropods and fragmentary fish and turtle remains. Although the material comes from a small area, we do not speculate whether or not it comes from a single individual, despite the specimens' match in size.

The material was cleaned using brushes and needles. Measurements were obtained with a digital caliper with 0.01 mm accuracy. Photographs were taken with a Canon X5 digital camera, and the images were processed using Adobe Illustrator and CorelDraw X8.

4. Systematic paleontology

Crocodyliformes Benton and Clark, 1988

Eusuchia Huxley, 1875

Crocodylia Owen, 1842

Alligatoroidea Gray, 1844 *sensu* Norell *et al.*, 1994

Globidonta Brochu, 1999

Gen. and sp. indetermined

Material. CPC 1483, isolated posterior tooth.

Locality and horizon. La Parrita is located 25°32'68.4"N, 101°20'93.7"W, in the municipality of General Cepeda, Coahuila. Cerro del Pueblo Formation, Campanian, Late Cretaceous.

Description. CPC 1483 is small, labiolingually compressed and distomesially inflated tooth crown (Figures 3A to 3C). It has fine irregular striations that radiate from the apex to the base; they are deeper at the distal side of the tooth. They lack carinae. The base of the crown is oval in outline with a maximum width of 9.2 mm.

4.1 EUSUCHIANS INDETERMINATED

Material. CPC 1479, anterolateral portion of right maxilla; CPC 1480, medial fragment of right maxilla; CPC 1481, medial fragment of

left maxilla; CPC 1482, fragment of left dentary; CPC 1484, paravertebral osteoderm; CPC 1485, paravertebral osteoderm; CPC 1486, lateral dorsal osteoderm; CPC 1487, lateral dorsal osteoderm; CPC 1489 dorsal vertebra; CPC 1490, vertebra with uncertain position.

Locality and horizon. The La Parrita site is located 25°32'68.4"N, 101°20'93.7"W, in the municipality of General Cepeda, Coahuila. Cerro del Pueblo Formation, Campanian, Late Cretaceous.

Description. CPC 1479 is a fragment that corresponds to an anterolateral portion of the right maxilla (Figures 4A and 4B). The external surface is ornamented with pits and grooves. The palatal surface is smooth and slightly concave, with two vascular foramina. Two alveoli are present. One of them still contains a root of a 6.4 mm diameter circular tooth. The other alveolus is badly preserved, but is smaller. Between both alveoli is present a well-developed depression.

CPC 1480 is a quadrangular fragment that corresponds to a middle portion of the right maxilla. The external surface bears pits, grooves, and rounded edges; the ventral surface is slightly rugose, which suggests the periosteum is lost. The internal structure of the maxilla is visible at a transverse break at the anterior end of the specimen. The bone is compact and is perforated by some vascular channels.

CPC 1481 is a fragment of the posterior part of the left maxilla, and like CPC 1479 and CPC 1480, it has deep pits and grooves on its external surface. The ventral surface is slightly rugose and, on the medial margin, has transverse grooves indicating the sutural surface for the nasal (Figure 4C). CPC 1482 is a small portion of the left dentary, 10.3 mm in length (Figure 4D). In transverse view it has a semitriangular shape. The external surface is ornamented with small pits and grooves. Two oval vascular foramina are present on the laterodorsal surface. Three alveoli are preserved. Of the three preserved alveoli the middle is the only complete one. All the alveoli are semicircular in outline and are nearly confluent. The medial

surface of the dentary is flat and bears longitudinal striations indicating the contact surface for the splenial. The semitriangular shape and the evidence of the contact with the splenial suggest that it is the anterior portion of the mandible.

CPC 1489 is a procoelous vertebral centrum lacking the neural arch (Figures 5A and 5B). The anterior surface is damaged, and the anterior socket of the centrum is not visible. The size of CPC 1489 (length = 19.84 mm, anterior width = 12.30 mm, posterior width = 13.21 mm) indicates a small animal, probably 1.5 m total length.

CPC 1490 is a small procoelous vertebral centrum (length = 12.11 mm, anterior width = 13.31 mm, posterior width = 9.97 mm). The position of the vertebra is uncertain. The neural arch is not preserved. The centrum is wider than high (Figures 5C and 5D). The central part of the neurocentral suture surface is exposed and rugose.

CPC 1484 is a quadrangular paravertebral osteoderm with a surface ornamented with pits and grooves. It bears a small anteromedial longitudinally running keel. The ventral surface is planar and has fiber marks (Figures 5E and 5F).

CPC 1485 is a paravertebral osteoderm with a rectangular outline. The dorsal surface is ornamented with pits and grooves. The ventral surface has fiber marks and four foramina (Figures 5G and 5H).

CPC 1486 (Figure 5I and 5J) and CPC 1487 (Figures 5K and 5L) are lateral dorsal osteoderms with a rounded shape. They are sculpted with large pits and grooves. The internal face is planar in both osteoderms, but CPC 1486 has several foramina.

5. Discussion

5.1. TAXONOMY DETERMINATION

Among North American crocodylians, only Alligatoroidea includes taxa with bulbous teeth, and these form the clade Globidonta (Brochu, 1999). The crocodylian specimen CPC 1483 from La Parrita locality can be referred to globidontan alligatorid based upon inflated tooth shape. An accurate determination to genus or species level based upon the isolated tooth is impossible, because the teeth of North American Late Cretaceous basal globidontans *Brachychampsa*, *Stangerochampsia* and *Albertochampsia*, resemble each other (Carbot-Chanona, 2014). In all of them, the teeth are striated and inflated with oval cross sections, similar to CPC 1483. Only the posterior teeth of the globidontan *Bottosaurus*, presently known only from eastern North America, can be differentiated from those of *Stangerochampsia*, *Albertochampsia* and

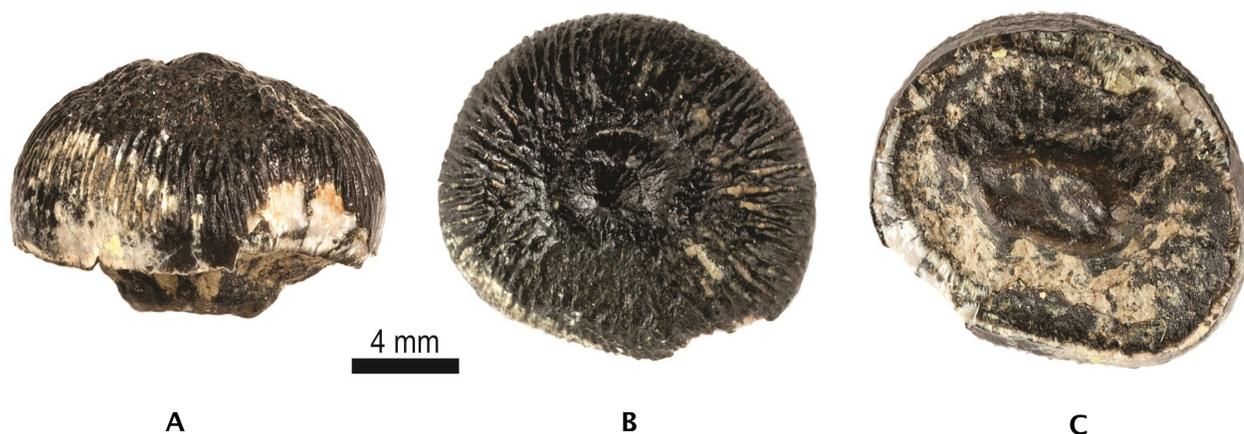


Figure 3 Globidontan alligatoroid from La Parrita site, Cerro del Pueblo Formation. CPC 1483, isolated posterior tooth in lateral (A), apical (B) and basal (C) view.

Brachychamps, because the teeth are laterally compressed and have strong striations that give a wrinkled appearance (Erickson, 1998).

Previously, crocodylian material from Cerro del Pueblo Formation was referred to as *Brachychamps montana* by Aguillon-Martinez (2010), although the identification should be treated with caution, and be referred as *Brachychamps* sp., or even better, only as Globidonta, due to the lack of diagnostic characters. In consequence, the crocodyliform specimen from La Parrita, is the second record of a globidontan

alligatoroid from the Cerro del Pueblo Formation, extending the geographical record of this taxon in North America and increasing the crocodylian diversity in the Late Cretaceous of Mexico.

On the other hand, the vertebrae CPC 1489 and CPC 1490 described in this work can only be referred to Eusuchia based on the strong procoely of the centrum, which characterizes the advanced eusuchian forms, including Crocodylia (Benton and Clark, 1988; Brochu, 1997; Molnar *et al.*, 2015). The neurocentral suture surface exposed in

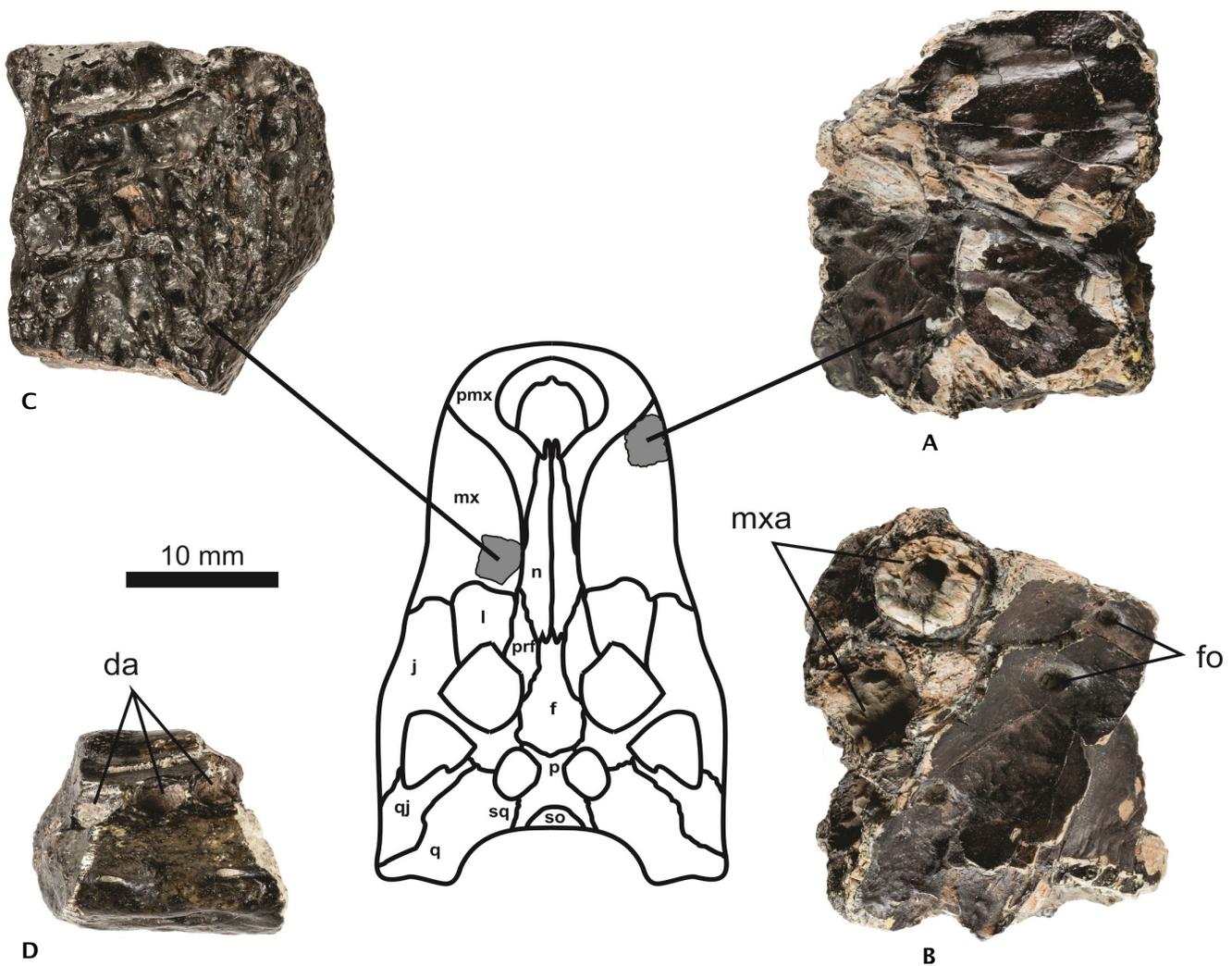


Figure 4 Eusuchia indetermined from the La Parrita site, Cerro del Pueblo Formation. Skull and mandibular elements: CPC 1479, anterolateral portion of right maxilla in dorsal (A) and ventral (B) view; CPC 1481, fragment of left maxilla in dorsal (C) view, and CPC 1482, fragment of left dentary (D). Abbreviations: da, dentary alveolus; fo, foramen; mxa, maxillary alveolus..

the vertebra CPC 1490 is indicative of a non-adult specimen, because in crocodylians closure of neurocentral sutures in cervical vertebrae happens in large individuals (Brochu, 1996). In the same way, the general shape of the osteoderms, as well as the presence of a keel in CPC 1484, allow referred the osteoderms to *Eusuchia* level.

5.2. PALEOENVIRONMENTAL SIGNIFICANCE

Temperature is the principal influence on global distribution of extant crocodylians (Marwick, 1998). Recent studies show that crocodylians are good ecosystem indicators (Mazzotti *et al.*, 2009). In consequence, it is possible to use fossil crocodylians as paleoclimatic indicators. Recent members of the Alligatoridea are intolerant to salt water because they lack salt glands (Taplin *et al.*, 1982). Consequently, the extant alligators cannot live in estuaries or coastal areas.

The association of charophyte oogonia with globidontan alligatoroid, as well as dinosaurs at La Parrita, provides evidence for fluviolacustrine environments on a delta plain (Vogt *et al.*, 2016) and a tropical climate. Mud cracks, interpreted here as desiccation cracks, repeatedly found on the surfaces of sandstone layers at Las Águilas, also suggest terrestrial or even subaerial conditions.

The vast majority of vertebrate remains at La Parrita are fragments or isolated bones that may occur isolated in the sandy layers, or in taxonomically mixed surface clusters, covering an area of a few up to about 100 m². The size of the specimens ranges between 0.5 mm up to about 150 mm. These concentrations were likely assembled in depressions or channels that were permanently or frequently flooded, and contain a remarkable amount of crocodylian material.

The vertebrate and invertebrate assemblages as well as plant material and sedimentary structures in the Cerro del Pueblo Formation suggest deposition in multiple stacked channelized sand bodies on a permanently changing delta plain (Vogt *et al.*, 2016). The depositional setting was influenced by cyclically fluctuating paleoenvironments of intermittent shal-

low marine, brackish to fresh water, or even subaerial conditions. The abundance of oysters throughout the measured section indicates frequent mixing of salt and fresh water, generating a changing zone of brackish environments (Vogt *et al.*, 2016).

6. Conclusions

We recognize the presence of the first evidence of globidont alligatorid specimen from La Parrita locality, Cerro del Pueblo Formation, based in the morphology of an isolated tooth. This extended the geographical record of Globidonta in North America. The vertebrae and osteoderms are referred to *Eusuchia*, due to the lack of morphological characters to genus or species level determination, but are important because the Mesozoic record of crocodyliforms in Mexico is scarce.

A mixture of depositional environments exists at La Parrita. The oscillating sequence of sandstones and siltstones suggests deposition under brackish conditions. The presence of oyster banks, abundant bivalves, ammonites, gastropods, and marine vertebrates (sharks and mosasauroids) indicate shallow marine environments. On the other hand, continental to brackish to fresh water conditions is indicated by characean oogonids, vascular plants, shell fragments of trionychid turtles, crocodyliforms and dinosaurs (*e.g.*, Meyer *et al.*, 2008).

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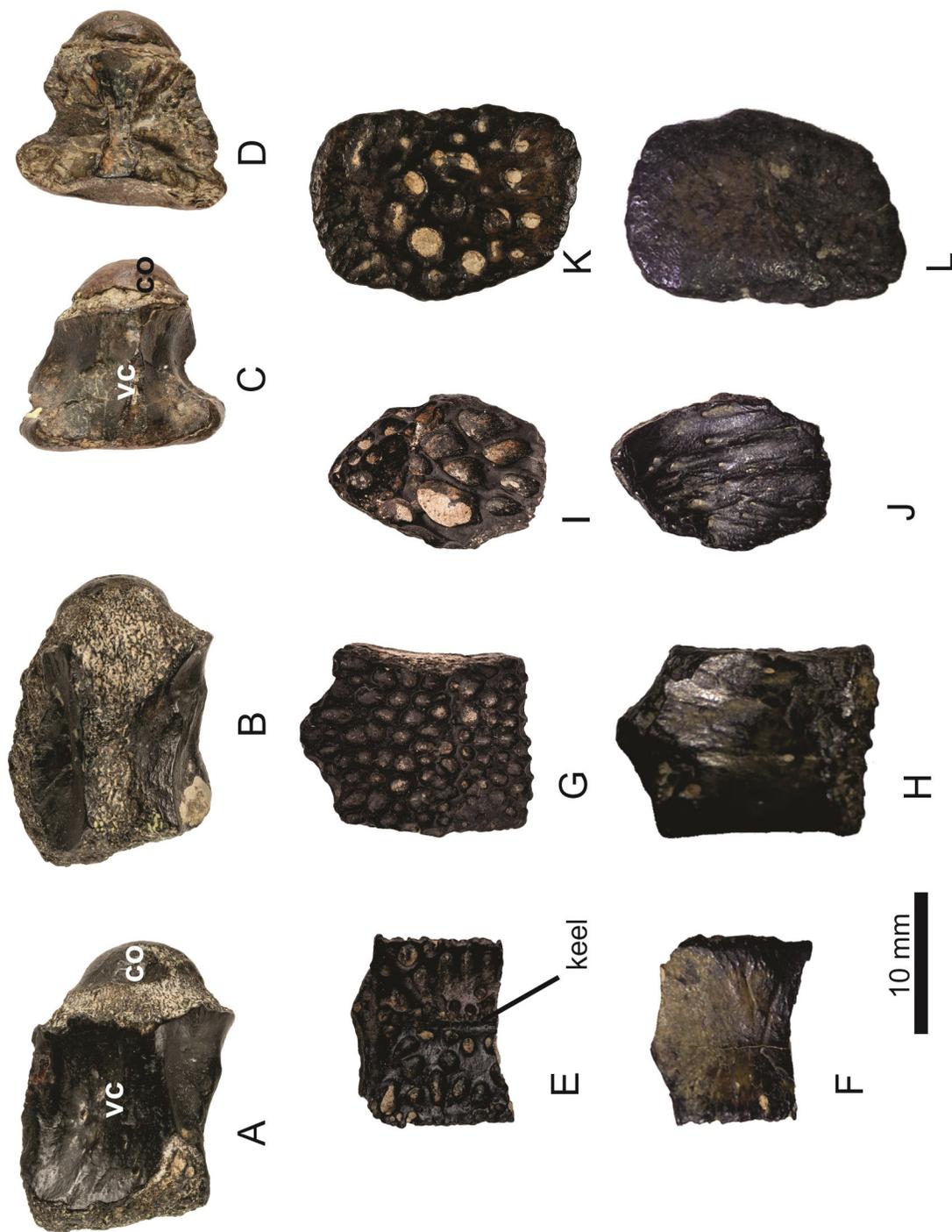


Figure 5 Eusuchia indetermined from the La Parrita site, Cerro del Pueblo Formation. Postcranial elements: CPC 1489 dorsal vertebra in lateral (A) and dorsal (B) view; CPC 1490, vertebra with uncertain position in ventral (C) and dorsal (D) view; CPC 1484, paravertebral osteoderm in dorsal (E) and ventral (F) view; CPC 1485 (I), paravertebral osteoderms in dorsal (G) and ventral (H) view; CPC 1486, lateral dorsal osteoderm in dorsal (I) and ventral (J) view; CPC 1487, lateral dorsal osteoderm in dorsal (K) and ventral (L) view. Abbreviations: co, condyle; vc, vertebral centrum. Scale bar = 10 mm.

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