

Paleostress and Geometric Analysis of the Southern Rio Grande Rift in the Northern Chihuahua Trough: Franklin and Potrillos Mountains as examples

Dana Carciumaru, Roberto Ortega

Dana Carciumaru

danadierna@gmail.com

Orbis Consultores en Geología y Geofísica SC, Retorno Calafia 120, Col. Calafia, La Paz, BCS, 23054.

CICESE, Unidad La Paz, Miraflores 334, Fracc. Bellavista, La Paz, BCS, 23050.

Roberto Ortega

CICESE, Unidad La Paz, Miraflores 334, Fracc. Bellavista, La Paz, BCS, 23050.

ABSTRACT

We present a deformation history of Potrillo and Franklin Mountains in El Paso, TX, using paleostress and geometric analysis. Application of paleostress and geometric analysis shows that low angle normal faults exposed at Anthony Gap and regions to the north of the Franklin Mountains record differential displacement that increase from North to South along the strike. As displacement on the faults increases, the hanging walls of the faults rotated counterclockwise and transverse faults which accommodate flexure developed as a consequence of that movement. The hanging wall of the faults rotates in a counter-clockwise direction and the transverse faults accommodate the developing flexion. The transverse faults are thought as the result of an isostatic rebound of the footwalls due to the tectonic denudation of the hanging wall. The average elastic thickness of the lithosphere is 23 km. This flexure and isostatic rebound take place in a late stage which was concurrently developed to the mountain range genesis. However, transverse faults do not cut the low angle faults suggesting that the two sets of faults are part of the same system. Increase in displacement along the faults is mirrored by an overall increase in elevation in the Franklin Mountains, which suggests that fault activity is critical to the uplift of the mountain range with respect to the surrounding basins. Estimates of the stresses from small scale faults in the Franklin Mountains indicate that the greatest extension direction was ENE-WSW and the greatest compression direction was subvertical. Low angle normal faults in the Potrillo Mountains belong to two different deformation episodes. The first family of faults occurs along the eastern side of the mountain range. This is cut by younger NNW trending normal faults which bound the range at the East. Fault slip analysis, calcite vein arrays and the geometry of the faults indicate that the first set of low angle normal faults record NNE directed extension and vertical shortening rotated approximately sixty degrees in a counter-clockwise sense as inferred from the trend of the east Potrillo fault. The second generation of faults have strike parallel slickenlines which trend to the ESE. Fault slip inversion indicates that these faults were formed in response to left oblique slip, with a moderately Southeast plunging shortening direction and a shallowly plunging WNW trending extension direction. These observations are most consistent with the first low angle faults in the Potrillo Mountains tracking an early phase of extension distinct from low angle faults in the Franklin Mountains. The second set of faults may have formed in a similar stress field to that recorded in the Franklin Mountains. The combined dataset is most consistent with two phases of extension across low angle normal fault in the southern Rio Grande rift, the first was caused by NNE extension, and the second by ESE directed extension.

Keywords: geometric analysis, low angle normal faults, paleostress, Rio Grande rift.

RESUMEN

Presentamos un estudio de la deformación de las montañas Potrillo y Franklin, en El Paso Texas, usando un análisis geométrico y de paleoesfuerzos. Utilizando análisis geométrico se muestra que las fallas normales de bajo ángulo expuestas en Anthony Gap y en las Montañas Franklin registraron un desplazamiento diferencial, con aumento de los desplazamientos de norte a sur a lo largo de las fallas. Las partes superiores de las fallas giran en un sentido contrario a las manecillas del reloj y se desarrollan fallas transversales que acomodan la flexión. Se cree que las fallas transversales resultan del rebote isostático del piso de la falla debido a la denudación tectónica del techo de la falla. El espesor promedio de la litosfera es de aproximadamente 23 km. Esta flexión y el rebote isostático ocurren en una etapa tardía de la extensión la cual ocurre paralelamente a la formación de las cordilleras. Sin embargo, las fallas transversales no cortan a las fallas de bajo ángulo, sugiriendo que los dos pares de fallas se forman durante el mismo evento de deformación. El aumento de los desplazamientos en las fallas se refleja en un aumento general de la elevación de las Montañas Franklin, sugiriendo que las fallas son críticas para el levantamiento de la cordillera con respecto a las cuencas adyacentes. El análisis del estado de esfuerzos de las fallas de pequeña escala en las Montañas Franklin nos indica que la máxima dirección de extensión era este-noreste/oeste-suroeste y la máxima dirección de compresión era subvertical. Las fallas normales de bajo ángulo en Las Montañas Potrillo pertenecen a dos episodios de deformación diferentes. La primera familia de fallas se produce a lo largo del lado oriental de la cordillera. Ésta es cortada por fallas normales más jóvenes con dirección norte-noreste que limitan la cordillera de las Montañas Potrillo. El análisis de las fallas de deslizamiento y la geometría de las fallas indican que el primer conjunto de fallas normales de bajo ángulo registró una extensión norte-noreste y un acortamiento vertical girado aproximadamente sesenta grados en un sentido contrario a las manecillas del reloj, el cual fue inferido a partir del rumbo de las fallas en la parte este de la falla Potrillo (la cordillera que limita una falla de ángulo alto). La inversión de fallas transtensionales indica que estas fallas se forman en respuesta a un deslizamiento oblicuo de falla izquierda, con dirección de acortamiento moderadamente basculado al sureste, también basculando al oeste superficialmente, mientras que al noroeste el rumbo es en la dirección de extensión. Estas observaciones son más consistentes con las primeras fallas de bajo ángulo en las Montañas Potrillo siguiendo una primera fase de extensión distinta a las fallas de bajo ángulo en las Montañas Franklin. El segundo conjunto de fallas puede haberse formado en un campo de extensión similar a las Montañas Franklin. El conjunto de datos combinados es más consistente con las dos fases de extensión a través de fallas normales de bajo ángulo en el sur del Rift Rio Grande.

Palabras clave: Análisis geométrico, fallas normales de bajo ángulo, paleoesfuerzos, Rio Grande Rift.

BOL. SOC. GEOL. MEX. 2017
VOL. 69 NO. 3
P. 669 – 689

Manuscript received: November 2, 2016
Corrected manuscript received: May 8, 2017
Manuscript accepted: July 3, 2017

1. Introduction

The Rio Grande rift is an important lithospheric scale feature associated with significant thinning of the crust and mantle lithosphere, high heat flow and low upper mantle seismic velocities (Keller *et al.*, 1990). One striking feature is the presence of low angle normal faults. The fault systems in the Rio Grande rift have been affected by considerable rotations, tilting and change in the stress field (Seager, 1981). Thus, understanding the nature of the low angle normal faults is not an easy task.

The post-Laramide stress regime is widely recognized from previous works (Baldrige *et al.*, 1980; Morgan and Golombek, 1984; Aldrich *et al.*, 1986; Morgan *et al.*, 1986). They suggest that the opening of the Rio Grande rift is divided in two distinct periods associated to different stress regimes. The first phase of deformation occurred from 30 to 18 Ma, forming shallow basins bounded by low-angle normal faults. This period is associated with volcanism and is attributed to thinning of hot lithosphere with a shallow brittle-ductile transition (Morgan and Golombek, 1984; Keller *et al.*, 1990; Baldrige *et al.*, 1995). Such low angle normal faults are only known from isolated areas, and their nature is still a point of debate. A later rifting phase, beginning 10 Ma is associated with classic Basin and Range style block faulting (Keller *et al.*, 1990) and the delineation of the present interconnected rift basins and flanking uplifts. In contrast to these conclusions, Henry and Price (1984) inferred that the extension direction during the early phase of rifting was East-West and then it switched to Northwest-Southeast recently. The current stress field in the Rio Grande rift is West-Northwest, although the stress field may vary locally (Zoback and Zoback, 1980, 1989).

According to Aldrich *et al.* (1986) many faults and dikes are associated with reactivated structures and their orientations cannot be used to infer the orientations of the principal stress. Using dike orientations in plutons, in addition to detailed dating studies, they found the orientation of the least compressive stress was N-NE-S-SW and during the

latter phase was associated with West-Northwest to East-Southeast extension. However, Aldrich *et al.* (1986) and Chapin *et al.* (1994) regarded early Miocene dikes in the roofs of plutons as being the best indicators of regional stress orientations since these are only known at two locations in the rift. Therefore, they may not be representative of the regional stress regime. In this study, fault slip inversion has been used to determine the state of stress adjacent to major low angle normal faults in the southern Rio Grande rift in order to clarify the orientations of the principal stresses during rifting. One of the biggest questions in the southern Rio Grande rift is determining if normal faults were formed in two phases of deformation or they evolved in a single progressive deformation event. Extensive faulting and tilting are inferred to have occurred during two distinct extensional phases in the Rio Grande Rift on the basis of structural studies (Collins *et al.*, 1994; Henry and Price, 1984; Aldrich *et al.*, 1986; Morgan *et al.*, 1986; Seager *et al.*, 1984).

This study focused on two ranges in the southern Rio Grande rift which expose low angle normal fault systems, the first in the Franklin Mountains at Anthony Gap, the second exposed in the East Potrillo Mountains (Figure 1). These two ranges are found at the northern margin of the Chihuahua Trough, a major Mesozoic basin. Through detailed structural analysis we attempted to constrain the stress field during the formation of the normal faults, determine if the faults are rotated due to younger extensional phases and analyze the relationship between the low angle faults and range uplift.

2. Methods

The main assumption necessary to obtain the stress tensor based on fault orientation is that the direction of the tangential traction on the plane tends to be parallel to the slip direction; this tangential traction is also assumed to be parallel to the shear stress (Ramsay and Lisle, 1997). In this way,

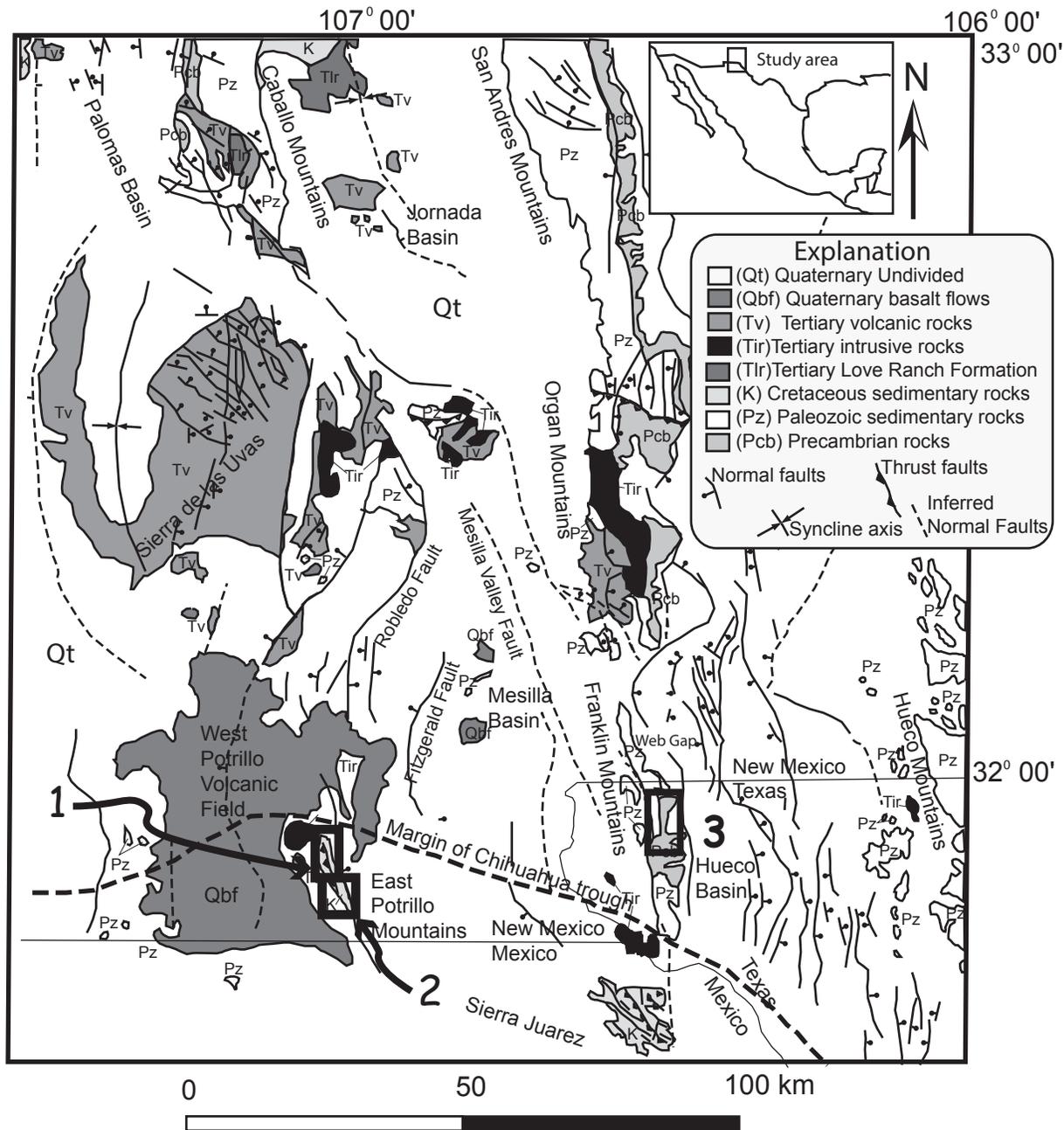


Figure 1 Geological map of the region. The two studied areas are the East Potrillo Mountains (insets 1 and 2) and the Franklin Mountains (inset 3). Both ranges are located at the northern margin of the Chihuahua trough (modified from Ruiz, 2004).

we assume that the increment of the principal strain axes coincide with the principal stress axes. This is the basic assumption to apply stress inversions. Therefore, if the tangential traction can be determined on a variety of fault planes formed at the same time in a region, it may be possible to make inferences about the regional stress field

(Angelier, 1984). The direction of the maximum resolved shear stress is constrained by experiments and observations to lie in the plane of the fault and in the σ_1 - σ_3 plane. Mohr-Coulomb behavior allows making several assumptions about the orientations of the principal stresses. These assumptions include the fact that fractures form at $\sim 30^\circ$ to σ_1 and

the least compression direction, σ_3 , generally lies at $\sim 60^\circ$ to the fractures. However, low angle normal faults are difficult to reconcile in terms of stress inversion as they cannot be easily explained by assuming Mohr-Coulomb failure criteria (Davis *et al.*, 1983).

In the Potrillo Mountains, faults with the best slickenlines were found in silica-cemented calcareous siltstones, which are common in the Hell-to-Finish formation. Slickenlines within major fault zones are rarely observed due to poor exposures and pervasive cataclasis. Therefore, our analysis focused on minor faults with small displacements, which may be more appropriate for stress inversion than large displacement faults (Erslev, 2001). At each outcrop, fault strike and dip as well as slickenline trend/plunge, and shear sense were measured. All data were inverted using the inverse methods of Michael (1984), and the inverse method of Marret and Allmendinger, (1990) to estimate the regional stress tensor. The data were collected using the regional geological map of Seager and Mack (1994) as a guide, we identified the major structures and collected the field data in general. Subsequently, it was possible to analyze the stress tensor in a regional sense by comparing our results with the data using only the geological map information.

3. Data

We used fault data from existing 1:24000 geological maps (Seager and Mack, 1994; Harbour, 1972; Kelley *et al.*, 1983) as well as field data using topographic maps as field guide. Because slickenline data are very important in stress and strain analysis, we focused on fault striations. In addition, we collected other structural data such as extensional veins, fold limb orientations, cleavage planes and bedding directions. This data set is used to analyze differential and mean stresses using the Mohr circle relationships of the associated faults. In the Anthony Gap area, we collected fault data from the high angle normal faults which connect to the

large displacement low angle normal faults for this analysis.

3.1. FRANKLIN MOUNTAINS

Low angle normal faults in the northern Franklin Mountains have been described by several researchers (Harbour, 1972; Seager, 1981; Kelley *et al.*, 1983; Figures 1987; Ruiz, 2004). Kelley *et al.*, (1983) and Figures (1987), who recognized large displacement gradients from North to South across the normal faults, with displacements close to zero at Web Gap and approximately 1000 meters at the state line (Figure 2). This differential displacement requires that the hanging walls of the normal fault system undergo rotation in the axis oriented normal to the fault as a consequence of the differential displacement. To quantify the rotations in the axis normal to the fault, three cross sections (Figure 3 and 4) were drawn across the region of the largest gradient of displacement (Box in Figure 2).

These sections show that horizontal displacement increases from ~ 160 meters at the location of section A-A' to ~ 220 meters at section B-B', and to 880 meters at section C-C' (Figure 3 and 4). This slip distribution implies that the degree of rotation of the hanging wall between section A-A' and B-B' is less than that between B-B' and C-C'. If the rotation occurs only by a rigid body, then the displacement gradient gives a direct measure of the angular rotation. However, this is not the case because rotation and displacement occur at the same time (Figure 4b). To quantify this effect, bedding in the footwall was used as a reference frame to estimate the angular rotation. The average bedding plane within the footwall in the area of the box in Figure 2 was calculated. Then the angle between the bedding planes in the hanging wall was measured with respect to the average bedding in the footwall and plotted as a function of the along strike distance. If the axis of rotation occurred entirely by rigid-body rotation the angle between bedding in the footwall and hanging wall should mimic that one estimated by the displacement gradient (Figure 4). These measurements should be

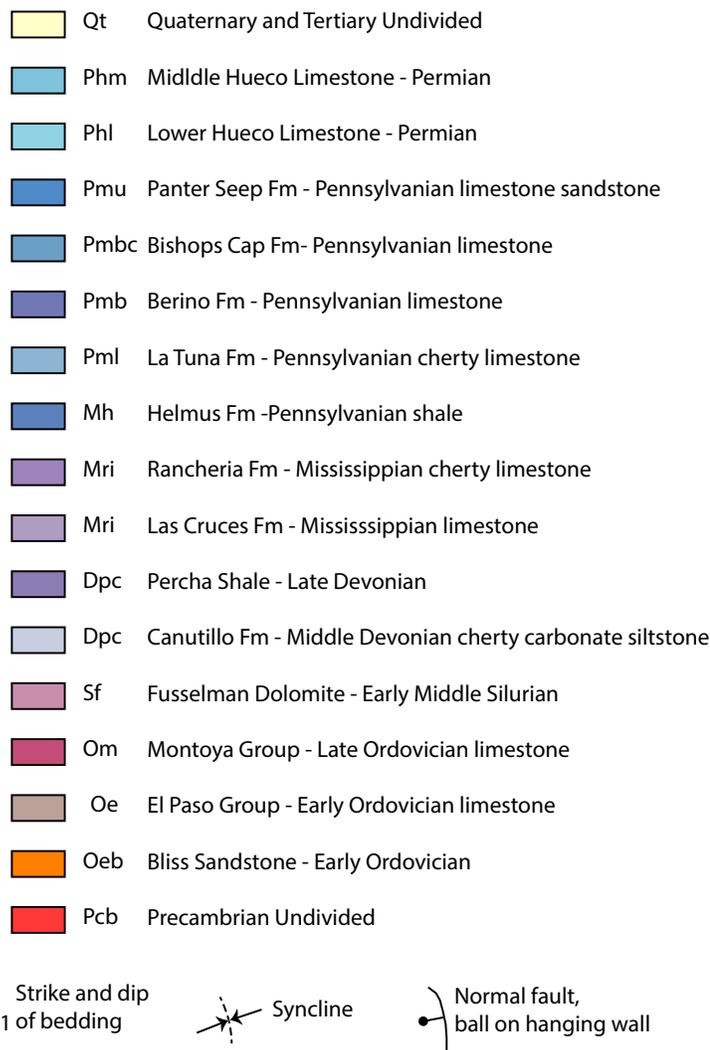


Figure 2 Continuation. Geological map of the Anthony Gap area, northern Franklin Mountains (after Kelley *et al.*, 1983; Figures, 1987, Harbour, 1972).

taken within the same lithology which in this study was in the Panther Seep and Tuna formations (Fig. 2). Inspection of Figure 4 shows that in most cases the angle between bedding planes is less than that one predicted by rigid body rotation alone, this is due to displacement that occurred previously at different piercing points which imprinted that rotation angle before (Figure 4b). It is worth noting that we only collected bedding information from La Tuna and Helms formations (Figure 2) to avoid errors from different lithology. It is possible that other events of deformation caused variations of angles of bedding planes. For example, there

are different thrust sheets in the study area which preceded normal faulting and could have caused variations in bedding. Therefore, we are afraid that this may be source of errors in our interpretation. On the other hand, there is a general increase in discordance between the hanging wall and footwall as displacement increases indicating rigid-body rotation was an important process. Other processes might have accommodated the rotation of the hanging wall, given the observed displacement gradient. In the Northeast corner of the box in Figure 2, two-left lateral strike slip faults occur, and are approximately parallel to the

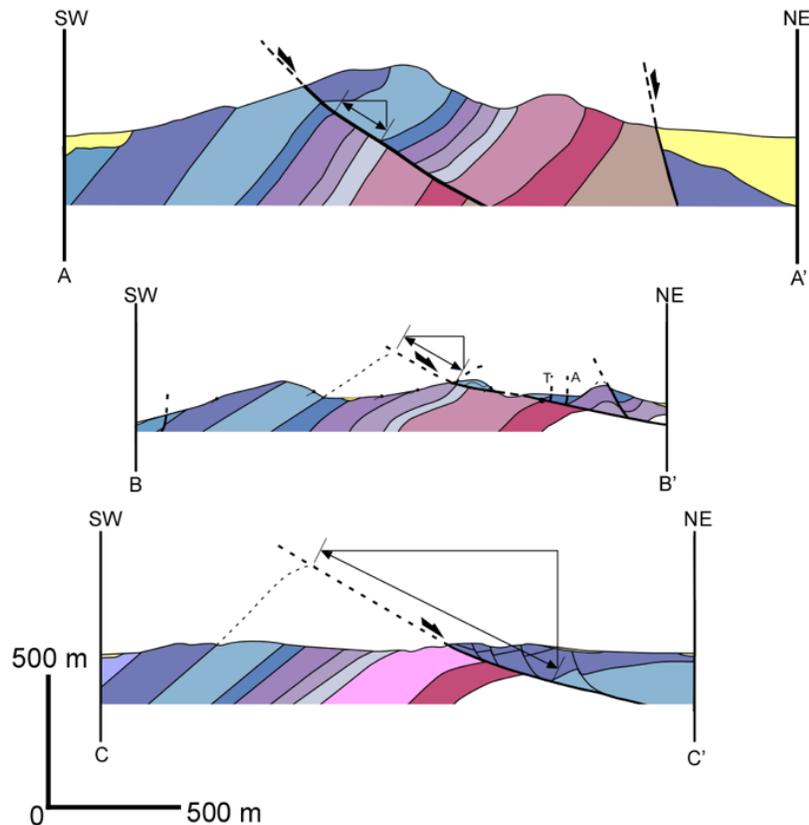


Figure 3 Cross sections A-A', B-B', and C-C'. Cross section A-A' is redrafted from Kelley and Matheny (1983). On section B-B', T indicates displacement toward the viewer, A indicates displacement away from the viewer. Right triangles above the sections show horizontal and vertical displacements inferred from stratigraphy. Color scheme for units is the same as Figure 2.

strike of the normal faults. These faults are associated with folding of the hanging wall strata and a dramatic change in bedding strike (figures 2 and 3). These faults do not penetrate the footwall of the normal fault system, have a short strike length, and occur in the region of the largest displacement gradient. These characteristics suggest that these faults were formed during normal faulting. Kinematically, these faults are consistent with accommodation of counter-clockwise rotation of the hanging wall. Therefore, it is likely that bending and rotation of the hanging wall was in part accommodated by left lateral strike slip faulting in the hanging wall of the fault.

Another geometric feature of the normal faults in the northern Franklin Mountains is the development of hanging wall synclines (Figure 2 and 3). The synclines trend to the N-NW, and are parallel to fault strike, have gently East dipping western

limbs adjacent to the faults, flat hinges and moderately East dipping eastern limbs. The fold axes are shallowly plunging and generally perpendicular to the inferred fault slip direction.

Cutting the hinges of the folds there are numerous small displacement normal faults, many with a conjugate geometry. Numerous tension veins occur and generally bisect the acute angle between the conjugate fault sets. Stereographic projections of poles to faults planes, poles to tension veins, and poles to bedding, show that all three structures are coaxial with respect to the syncline axis (Figure 5) demonstrating that all these structures probably formed during a single deformation event. The folds are geometrically similar to extensional fault propagation folds which are observed in nature (Shaw *et al.*, 1997) and numerically modeled (Suppe *et al.*, 1990; Hardy and McClay, 1999). The tightness of the synclines is clearly related to fault

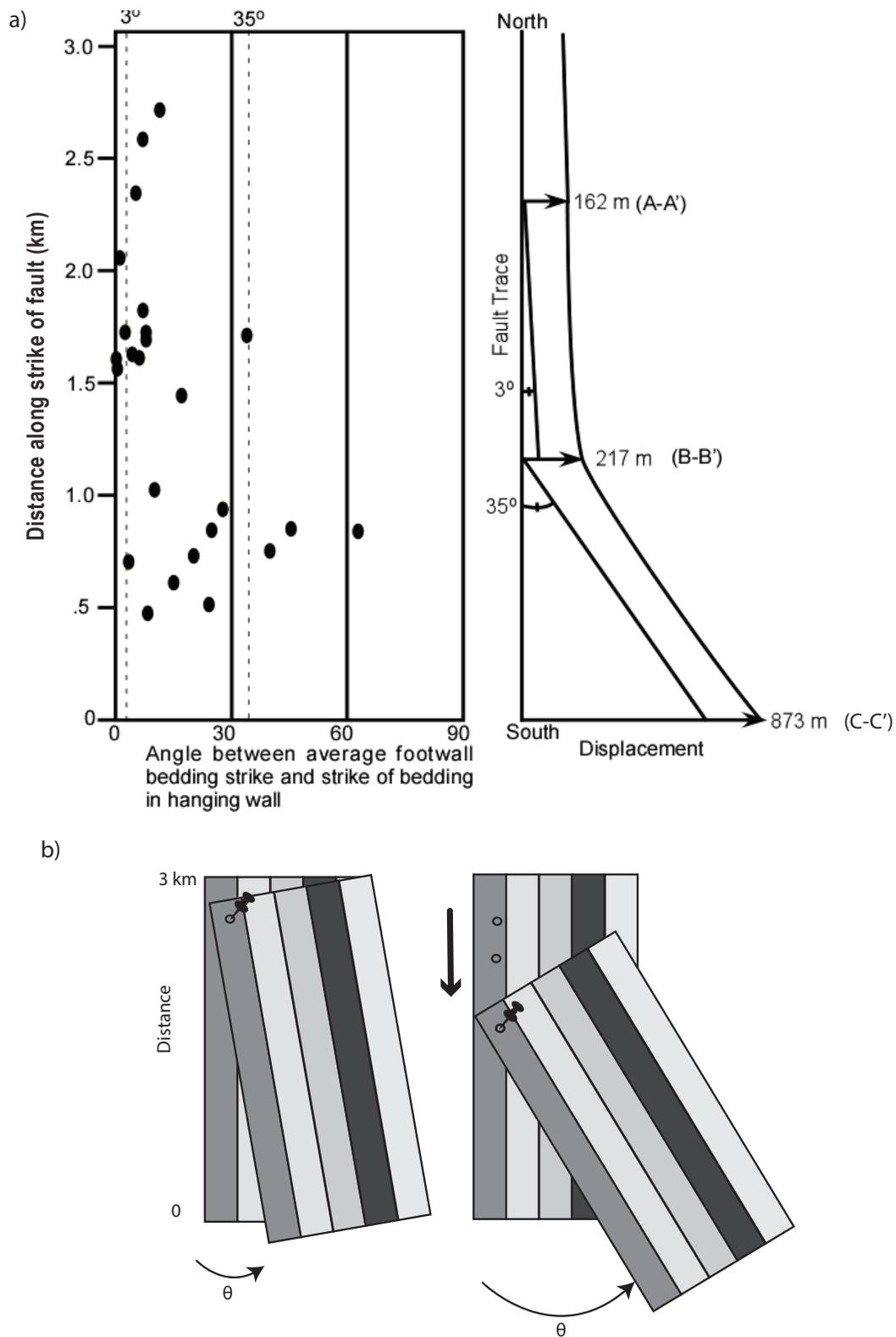


Figure 4 Plot of the angle between average footwall bedding strike and individual bedding measurements in the hanging wall in the box shown in Figure 2. Adjacent diagram shows inferred vertical axis rotations predicted from differential displacement along the normal fault.

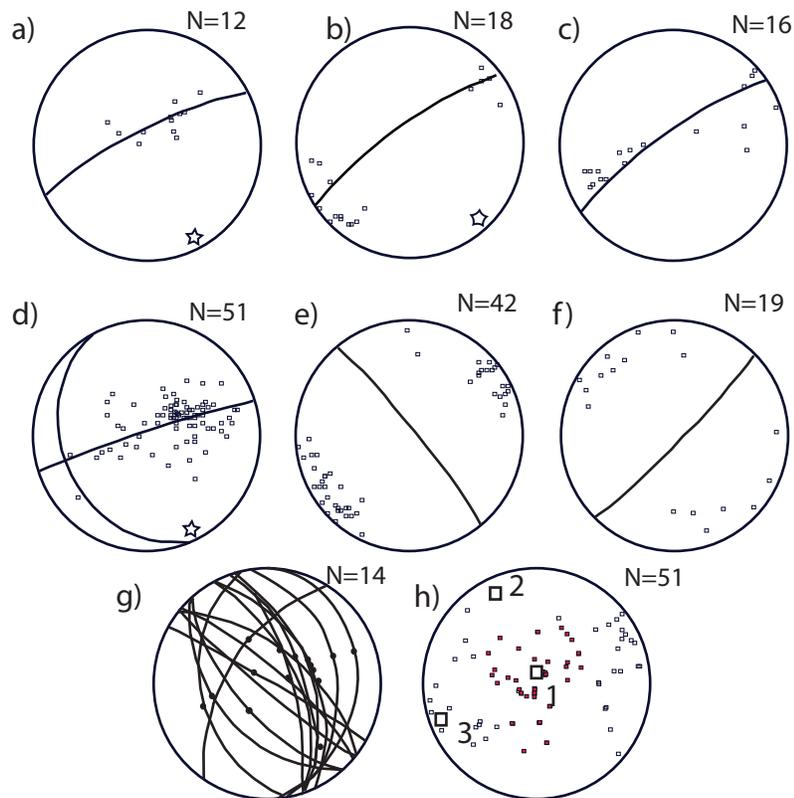


Figure 5 Stereographic projections from Franklin Mountains. a) Poles to bedding for outcrop shown in Figure 6. b) Poles to calcite veins for the same outcrop. c) Poles to fault planes for same outcrop. Note poles in a, b, and c share a common fold axis (shown by star). d) Poles to bedding for all beds shown in Figure 4. Star indicates inferred fold axis. e) Poles to first generation calcite veins. f) Poles to second generation calcite veins. g) Fault planes with slip direction indicated. h) Calculated greatest compression axis (red squares) and least compression axis (white squares) estimated from inversion of the data in h).

displacement with the folds dying out to the North and tightening to the South.

In order to analyze the stress-states during normal faulting and folding, small scale normal faults and associated calcite veins were studied, due to these structures have small displacement and are most likely to be representative of the instantaneous deformation (Ersley, 2001). Stereographic projections shown in Figure 5 summarize the data from the northern Franklin Mountains. As can be seen, most faults strike North-Northwest parallel to the axis of the range and have steep to moderate dips, with nearly down dip slickenlines. Stress inversion using the Michael code (Michael, 1984), of the data shows an extension direction towards East-Southeast-West-Northwest with a shallow plunge. The shortening direction is subvertical. Poles to calcite veins confirm a stretching direction

towards East-Northeast-West-Southwest, approximately perpendicular to the trend of the range (Figure 5).

A further constraint on paleostress state can be obtained from the study of conjugate fault sets. Most of the conjugate fault sets have dihedral angles (the acute angle between the fault sets) which are less than the 60° angle, as predicted from the Mohr-Coulomb failure criteria (Ramsay and Huber, 1983). Additionally, the faults frequently have irregular traces turn into calcite filled veins along their trace. These geometric features are consistent with the conjugate faults being transitional tensile fractures (Engelder, 1999). Transitional tensile fractures were formed where the least compressive stress was tensile. Such fractures were described according to the parabolic Griffith failure criteria, which predict a closure of the di-

hedral angle as the stress field becomes more tensile (Ramsay and Huber, 1987). Failure criteria in Mohr space is defined as:

$$\tau^2 = 4T_0^2 + 4T_0\sigma_n \quad (1) \quad \text{Equation 1}$$

where τ is the shear stress, T_0 is the tensile strength of the rock ($T_0 = C_0/2$, with C_0 = cohesive strength) and σ_n is the normal stress (Ramsay and Huber, 1987). This equation can be solved for magnitudes of the principal stresses if the values of T_0 are known and the dihedral angle can be determined. This is solved by differentiating equation 1 and solving for values where the tangents to the parabola are perpendicular to the predicted 2θ (twice the angle between the fault plane normal and the greatest compressive stress) derived from the dihedral angle between the faults. This gives values for shear and normal stress on the failure envelop that are tangent to the Mohr circle at failure, which fixes the radius of the Mohr circle, the

mean stress and magnitudes of the greatest and least compressive stress, as long as the dihedral angle is small. For large dihedral angles (greater than 45°) the parabola in Mohr space becomes effectively linear and many different Mohr circles become tangent to the envelop making a unique solution impossible. Fortunately, the maximum dihedral angle observed in this study is $\sim 45^\circ$, with a mean dihedral angle of 36° (Figure 6).

Values of T_0 were chosen to bracket the likely range of strengths of carbonate rocks, as limestones and dolomites dominate the rock types exposed in the Franklin Mountains and no rock mechanics experiments have been conducted using materials from the Franklin Mountains. For a lower bound, T_0 for Carrera Marble (Suppe, 1985) was used. This is a coarsely crystalline marble and has very low tensile strength relative to rocks with a finer grain size. For an upper bound, T_0 values for Blair dolomite (Suppe, 1985) were chosen as this rock has a very high tensile strength comparable to diabase

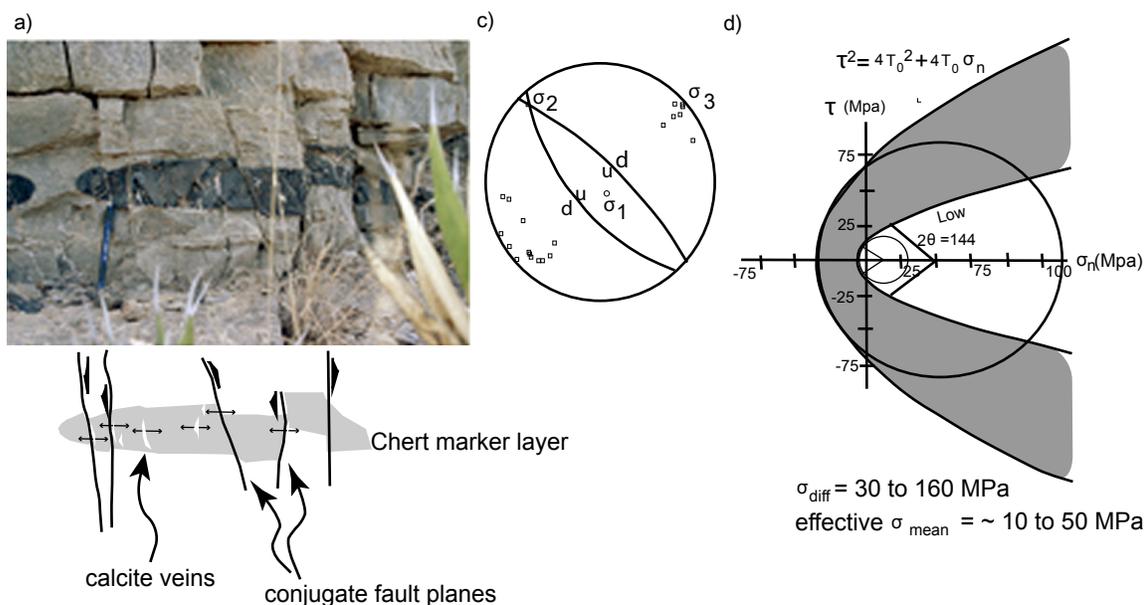


Figure 6 Diagram illustrating how stresses are estimated from geometry of conjugate fault sets. a) Photograph of outcrop with conjugate normal faults cutting the Berino Formation at Anthony Gap. b) Sketch of photograph showing relationship between calcite veins and the faults. c) Lower hemisphere equal area projection showing the orientations of the principal stresses derived from the conjugate faults. Dots represent poles to calcite veins which confirm the orientation of the least compressive stress derived from the conjugate faults. d) Mohr diagram illustrating the Griffith failure criteria and inferred Mohr circle derived from the geometry of the fault sets. Smaller envelope is calculated using T_0 values for Carrera Marble, outer envelope uses Blair Dolomite values. Gray shaded region corresponds to the range of likely stresses during faulting recorded in the Franklin Mountains.

and gabbro. Only maximum values for differential stress and effective mean stress can be made because all available experimental data have been collected on dry rocks, which are much stronger than rocks that deform in the presence of fluids (Ramsay and Chester, 2002). Additionally, the estimated principal stresses and mean stress are effective values because pore fluid pressure lowers the values of the principal stresses due to hydrostatic pressure counteracting applied compressive stress. Since only maximum values for the effective stresses and differential stress can be obtained the maximum dihedral angle of 45° was used in the calculations. This also provides a maximum estimate, since faults with smaller dihedral angles will result in stresses with smaller values.

The results suggest that the differential stress during faulting in the Franklin Mountains range between 30 and 160 MPa and the effective mean stress range between 10 and 50 MPa. To our knowledge, these are the only estimates of differential and mean stress which have ever been made based on the geometry of faults as input. The effective mean stresses are consistent with the faults forming in the upper five kilometers of the crust assuming that hydrostatic pressure was equal to or less than the lithostatic pressure.

The final feature of the Franklin Mountains is the occurrence of transverse normal faults which cut across the range at a high angle. These faults generally have small displacements and are locally associated with large fault parallel calcite veins (up to 1 meter in width). Cross cutting relationships between calcite veins demonstrate that these veins are younger than northwest trending veins associated with range parallel faults. Unfortunately, none of the planes of the transverse faults is well enough exposed to collect the data needed for fault slip inversion. However, the geometry of the veins, orientations and displacements on the faults are all consistent with the faults forming during range parallel extension. However, the faults do not appear to cross cut the northwest trending faults, and instead terminate at the low angle faults. This suggests that these faults are like-

ly linked to displacements on the low angle normal faults. Additionally, the faults are most common where the displacement along the low angle normal faults is largest, and disappear to the north as fault displacement decreases, supporting the interpretation that the faults were formed synchronously with displacement on the low angle faults.

3.2. POTRILLO MOUNTAINS

The complexity of deformation in the Potrillo Mountains makes geometric analysis of low angle normal faults a difficult and uncertain task (Carciumaru *et al.*, 2011). Therefore, we focused only on fault slip analysis, and analysis of tension vein arrays as these data can be easily correlated with the normal faults. Seager and Mack (1994) provide detailed descriptions of the structural geology of the Potrillo Mountains (Figure 7 and Figure 8). They describe the low angle normal faults which cut the central Potrillo Mountains, and concluded that the largest of the faults, exposed in the central Potrillo Mountains, is an extensional detachment surface active during Miocene rifting. This was based on the low angle of the structure and profound differences with the structures between the hanging wall and the footwall. Specifically, they noted that large scale Laramide folds and thrust faults predominate in the hanging wall, whereas the structure of the footwall is much less affected by Laramide contraction. Due to the effects of Laramide folding, no suitable stratigraphic or structural piercing points are found, thus it is difficult to estimate displacement on the fault system (Seager and Mack, 1994). However, intensity of fault gouge, stratigraphic constraints and structural discordance require a minimum of several hundreds of meters of displacement, although this value could be greater.

The low angle normal faults can be divided into two geometric groups (Figure 9). The first group has northeast strikes, shallow dips and down dip slickenlines. Based on cross cutting relationships these faults appear to be the oldest set of normal faults exposed in the Potrillo Mountains (Seager and Mack, 1994). These faults are often associated

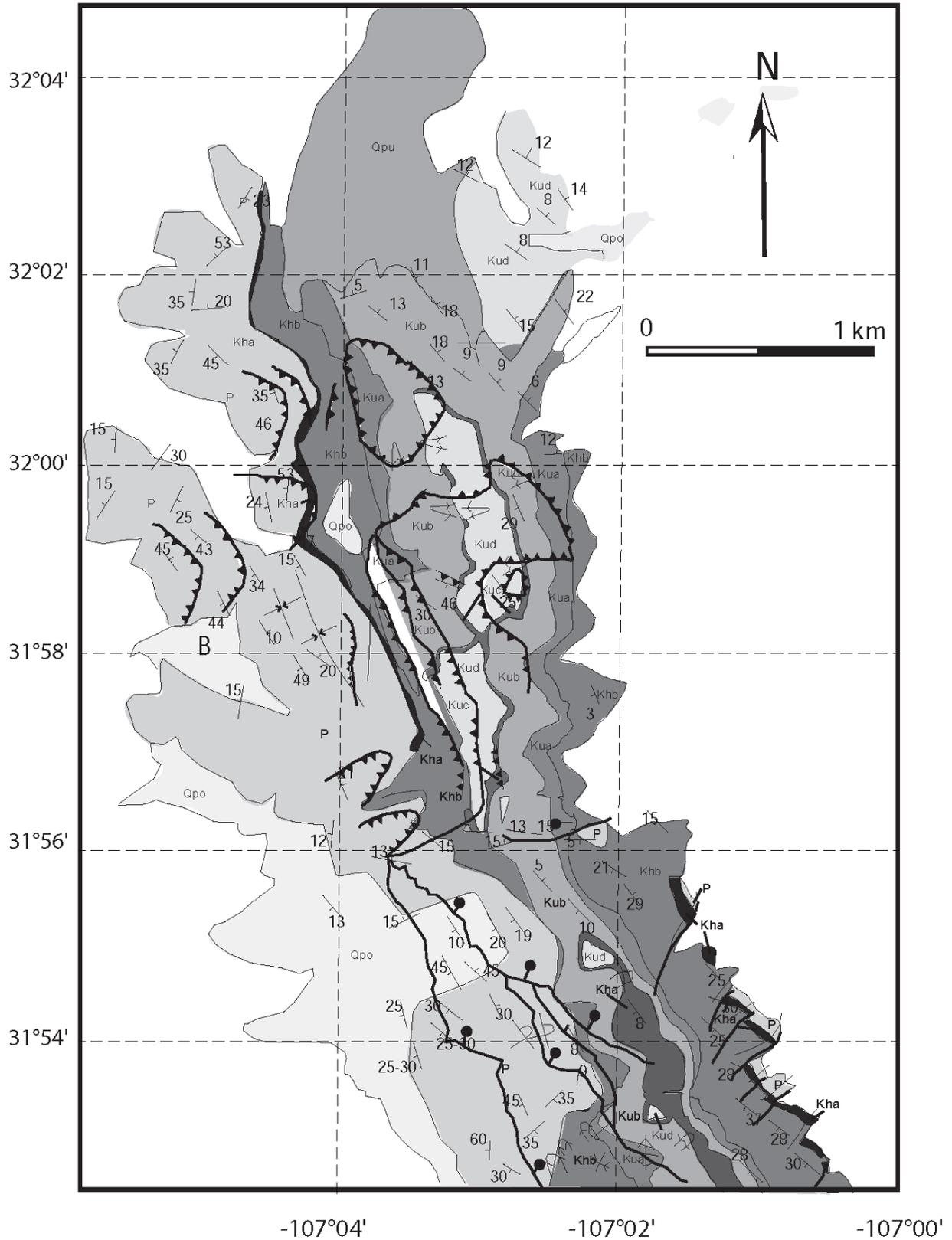


Figure 7 Geological map of the northern part of the East Potrillo Mountains. The two main formations are Hell-to-Finish and U-Bar, overlying Permian rocks. The general trend of the structures is oriented N36° W (after Seager and Mack, 1994).

- Qpu Younger piedmont slope deposits - Quaternary
- Qpo Older piedmont slope deposits - Quaternary
- Qcp Santa Fe Group. Camp Rice Fm - Quaternary
- Qpa Undifferentiated piedmont deposits - Quaternary
- Kue U-Bar Fm - Early Cretaceous massive limestone
- Kud U-Bar Fm - Early Cretaceous silstone-limestone
- Kuc U-Bar Fm - Early Cretaceous rudistic limestone
- Kub U-Bar Fm - Early Cretaceous sandstone
- Kua U-Bar Fm - Early Cretaceous lower limestone
- Khb Hell-to-Finish Fm - Early Cretaceous silstone
- Kha Hell to Finish Fm - Early Cretaceous conglomerate
- P Yeso San Andreas Fm - Permian Limestone

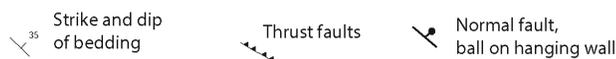


Figure 8 Continuation. Geological map of the southern part of the East Potrillo Mountains. The two main formations are Hell-to-Finish and U-Bar, overlying Permian rocks (after Seager and Mack, 1994).

with bleach zones around the faults with a strong fault parallel cleavage. The faults contain 10 to 50 cm thick gouge zones cored by brecciated cataclastite. Slickenlines on the faults are well developed, but are often obscured by a breccia. Applying fault slip analysis to this population of faults indicates that the faults record North-Northeast directed subvertical extension (figures 9a and 9b).

The second set of faults has a Northwest strike, shallow dips to the northeast and highly oblique slickenlines that plunge to the East-Northeast (figures 9c and 9d). Slickenlines are often nearly strike-parallel suggesting a significant transcurrent component to the displacement despite the shallow dips of the faults. In other ways, these faults are geometrically and texturally similar to the older set of low angle faults, which appear to sole into. Fault slip analysis suggests that these faults recorded East-Southeast subhorizontal extension with a moderately Southeast plunging shortening direction. Kinematic data are consistent with right lateral oblique normal displacement, quite distinct from the earlier set of normal faults. Geometry and kinematics of both sets of faults is quite distinct from normal faults exposed in the Franklin Mountains.

To further constrain the orientation of the least

compressive stress calcite veins associated with the normal faults were analyzed (Figure 10). The average orientation of the veins is approximately East-West, with steep to vertical dips as shown in the rose diagram of Figure 10. This geometry is consistent with North-Northeast directed horizontal extension. This is also consistent with extension direction estimated from the fault slip analysis performed on normal faults with North-Northeast plunging slickenlines. This is in agreement with the fact that the majority of the veins are found in the hanging wall of the large low angle detachment fault which cuts the central part of the East Potrillo Mountains.

4. Discussion and conclusions

The study of Carciumaru and Ortega (2011) from low angle normal fault exposures in the Franklin and Potrillo Mountains indicate a polyphase deformation history and cannot be reconciled with a single progressive deformation event. The oldest fault sets in the Potrillo Mountains appear to have formed during North-Northeast directed horizontal extension. Seager and Mack (1994) inferred, based on correlations with regional geology that

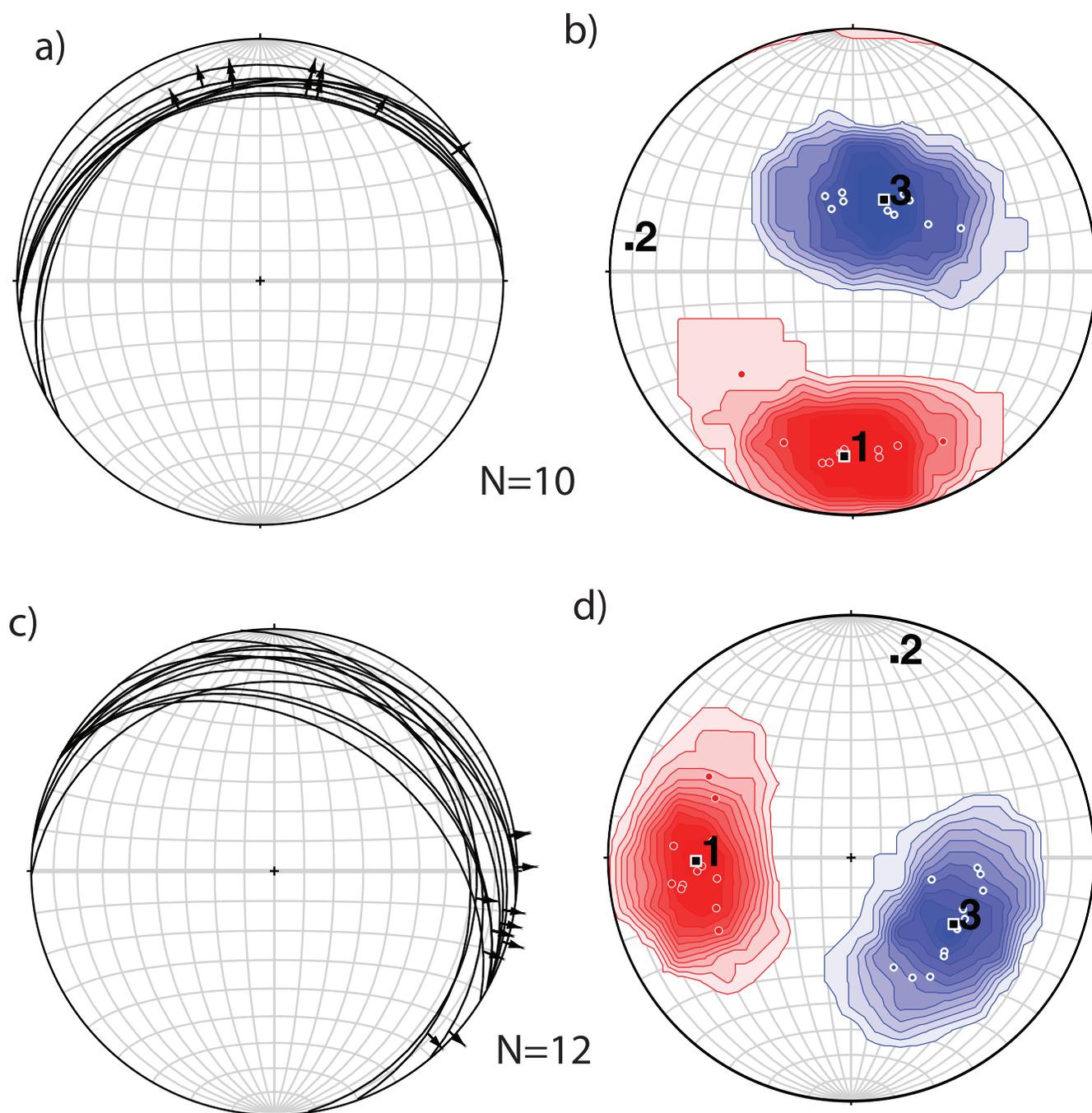


Figure 9 Stereographic projections from Potrillo Mountains using the FaultKin7 software (Allmendinger *et al.*, 2012). a) First generation fault planes, arrows show movement direction of the hanging wall. b) Stress tensor solution using the code of Michael (Michael, 1984) (black squares), in addition the contours using the Kamb method (Allmendinger *et al.*, 2012) with one sigma as significance level. c) Second generation fault planes, arrows show movement direction of the hanging wall. d) contours of the kinematics solution for the second generation.

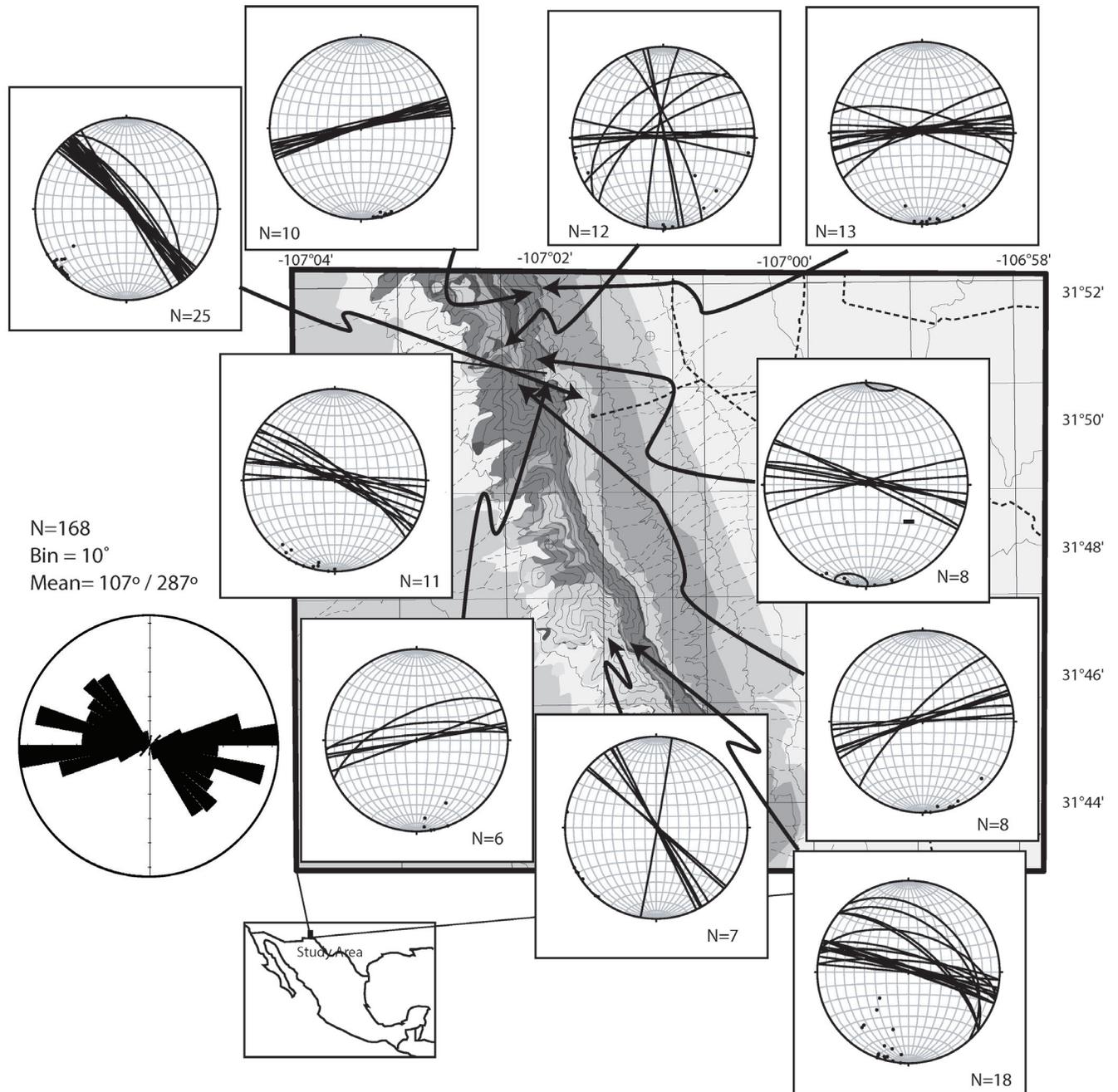


Figure 10 Map of the calcite veins in the central part of the East Potrillo Mountains, showing the planes and the poles to the planes. Also, the rose diagram shows the general tendency of the data.

these faults were formed during an early Miocene phase of extensional deformation. The newly collected data on the Rio Grande fault system support this interpretation as they demonstrate this fault system was formed in a stress field unlike from any other population of normal faults here studied.

Additionally, Miocene patterns of sedimentation within the rift are consistent with early rift basins trending to the Northwest, as well as North-Northeast extension (Seager and Mack, 1994).

The second set of low angle normal faults in the Potrillo Mountains is distinct in geometry and

kinematics to faults found in the Franklin Mountains and the older set of faults found in the Potrillo Mountains. Fault slip analysis on these faults indicates that they record East-Southeast extension with a moderate Southeast plunging extension direction. This stress field is similar to the East-Northeast horizontal extension inferred for the Franklin Mountains, although the inferred shortening direction is distinct. It is possible; perhaps even likely, that the orientation of these faults was in part controlled by the orientation of the preexisting low angle normal faults. This is indicated by the fact that the two fault sets appear to merge along strike. If true, then the geometry of the faults would not be fully representative of the stress field in which they were formed. Therefore, they may have formed during the same deformation event recorded in the Franklin Mountains. Alternatively, the faults may have formed during an intermediate stage of extension. If this interpretation is correct, this intermediate stage would be largely transtensional as is indicated by the highly oblique nature of the slickenlines. However, similar faults have not been found regionally, therefore the reactivation interpretation is preferred.

Low angle normal faults in the Franklin Mountains contrast strongly with faults in the Potrillo Mountains because they appear to record a single episode of extension. Additionally, the stress field for the faults is generally consistent with the extension direction inferred from the orientation of Quaternary normal faults found in both the Mesilla and Hueco Bolson's which flank the range. Furthermore, Zoback and Zoback (1980), based on the alignments of volcanic centers in southern New Mexico (Potrillo volcanic field), inferred an East-West oriented extension direction. Additionally, the data indicate that faults have not been significantly tilted following formation. A further observation is that fault displacement increases to the South as the elevation of the Franklin Mountains increases, suggesting that this displacement is related in a gross way to the uplift of this Mountain range. All of the features of the low angle normal faults in the Franklin Mountains are con-

sistent with the faults forming relatively recently, as these Mountains were uplifted. This interpretation is further supported by recent $^{40}\text{Ar}/^{39}\text{Ar}$ dating of jarosite from fault gouge in the Franklin Mountains which indicates a low angle fault movement in the last 6 Ma (Lueth *et al.*, 1998).

The late transverse faults in the Franklin Mountains indicate range parallel extension, distinct from that inferred from the low angle faults. The geometry of the faults is consistent with synchronous formation with motion on the low angle faults. Axen and Wernicke (1989) emphasized the role of isostatic rebound of the footwalls of the normal fault systems due to tectonic denudation. Their model emphasizes vertical displacements on a plane parallel to tectonic transport (plane strain) and did not address the effects of an along strike displacement gradient, like that seen in the Franklin Mountains. Since the normal fault footwall is denuded more to the South than to the North, the amount of footwall rebound would be expected to vary along strike. This type of gradient would produce flexure of the footwall parallel to strike. In turn, flexure leads to extension along the length of the flexure, providing a kinematic explanation for the development of the transverse faults. The magnitude of extension accommodated in this way would be small ($< 10\%$) given the flexure wavelength (~ 4 km) consistent with the small displacements on the transverse faults.

Displacements on the low angle normal faults also resulted in counterclockwise vertical axis rotation of the hanging walls of the faults. This rotation ranged from 3° in the North to as much as 35° in the South. Rotations of this magnitude have been measured paleomagnetically in the Española basin of the northern Rio Grande rift (Salyards *et al.*, 1994). This indicates that this process is wide spread within the Rio Grande rift. Notably the faults have listric curves (spoon shape), resulting in an arcuate fault trace. These fault traces are similar to those arcuate spoon shaped fault traces common in the map pattern of Quaternary normal faults in southern New Mexico (see Figure 1). This suggests that faults with similar listric geometry

underlie much of the southern Rio Grande rift. Overall the geometry and kinematics of the low angle normal faults is consistent with two phases of extension. The first phase of extension, is based on regional correlations, likely occurred in the Early Miocene, following wide spread volcanism (Chapin *et al.*, 1994). The second phase of extension has occurred over the last 10 Ma, and has had a dominantly East-West extension direction and has been accompanied by small-volume basaltic volcanism (Seager and Mack, 1994).

This study shows that low angle normal faults were formed during extensional phases in the southern Rio Grande rift. Prior studies about the rift have concluded that low angle normal faults were formed due to tilting of initially high angle faults (Seager, 1981; Kelley and Matheny, 1983; Lewis and Baldrige, 1994). These prior studies inferred that the thermal structure of the Rio Grande rift was fundamentally different from that one which is present today, leading to a shallow brittle ductile transition producing low angle faults (Morgan and Golombek, 1984; Keller *et al.*, 1990; Baldrige *et al.*, 1995). The presence of young low angle normal faults challenges this view. The results also suggest that low angle normal faults are common in the upper crust and form basin bounding normal fault systems which influence the development of surface topography. Large displacement gradients along the faults also produce vertical axis rotations independent of strike slip faulting and range parallel extension due to flexure of the footwall. The stress fields associated with low angle fault systems varied both in time and space with each of the fault systems forming in response to distinct stress fields. However, the overall pattern of deformation is consistent with two episodes of deformation, the first with North-Northwest directed extension and the second with East-West extension.

The results show that the Franklin Mountains underwent major low angle normal faulting during uplift of the range. A detailed study of the rift uplift in New Mexico using thermochronology is described by House and Kelley, 2003. In contrast, low angle normal faults in the Potrillo Mountains were

formed during an earlier phase of deformation and have likely been tilted from their initial orientation (Carciumaru and Ortega, 2011). These results support a two-phase extension model for the Rio Grande rift and demonstrate that low angle normal faulting occurred during both phases of extension and is in fact characteristic of the southern Rio Grande rift.

This article has reviewed the structural aspects of this region. We observe that the state of stress in Franklin and Potrillo Mountains does not necessarily represent two different periods. Stress inversion should be performed from data that represents a homogeneous state of stress. It is preferred to use contemporaneous fault data when reconstruction studies are performed. However, this paper is focused on determining if Rio Grande Rift was formed during two phases of extension. So, using the criterion of contemporary faults is not very adequate. For this reason, we have decided to analyze the faults arguing that the data belong to a homogeneous state of stress.

In a previous article, Carciumaru and Ortega, (2011), demonstrated that low angle normal faults can be analyzed by a back tilting technique using the cluster analysis method (Yamaji, 2000). In that study, Carciumaru and Ortega (2011) demonstrated that this data set can be represented in a homogeneous state of stress. The error range was low after performing a statistical analysis because the data converged to a single stress tensor. The method is based on a comprehensive combinatorial analysis. Data are analyzed by systematically varying the number of samples. This is done to eliminate outliers, and to reduce variability in case of epistemic errors. Once a series of input data is obtained, a new reference system results by slightly rotating the components of the faults. This simulates a back tilting.

Convincingly, it is inferred that these data belong to a homogenous state of stress. In the present article, we have used virtually the same data set, so we strongly believe that the analysis of stress is based on homogeneous data.

It is important to note that the study conducted on

the faults associated to rocks of different strength appear to have a high range of variability. Although it is preferable to know exactly the dynamic parameters of the rocks, this is not always possible because these studies are costly. We consider that an average rock range reported in the literature is sufficient for a general analysis. We observed that the origin of these faults occurs in a range of 40 to 150 MPa. We know that for Mexican earthquakes the stress drop is 4 MPa. The mean stress of these faults is of one order of magnitude greater than that observed in the crust in case of tectonic earthquakes. Probably the main reason is that the tensional stress requires higher strength at the beginning of the first crack rupture than a pure shear stress during an earthquake.

This analysis is an important achievement in geology. It is well known that stress inversion with fault data only uses angle measurements (strike, dip, rake); therefore, they do not contain absolute measurements. In general, the studies of structural geology rarely report absolute magnitudes of the stress values. In this article, we could identify at least a range of values.

As a practical example, if absolute stress values are known for a geomechanical analysis then, we can better predict a blowout during a drilling process. Rarely rotational changes are reported in structural geology studies when stress-strain inversions are performed. It is well known that the simplification of stress-strain tensor is made by assuming a symmetrical strain tensor. This symmetry appears when the rotation part is canceled. We have observed that the rotation component is very important, at least, in the case of the Franklin Mountains; therefore, our study went beyond a classical stress-strain tensor inversion adding information of the fault rotation in Franklin Mountains.

IN SUMMARY:

- Stress inversion shows that low angle faults do not always have the same origin despite being in close regions. At least this is the case of the Franklin and Portillo's Mountains.
- It is possible to add absolute values of magnitude to the stress inversion. This could be

done by adding a detailed analysis of the related structures in the field, which is a breakthrough for structural geologists.

- It is necessary to understand the limitation of our method as the case of a stress-strain analysis with significant rotations.
- The geometric characteristics of this area show that the region has a complex geology and the geological history should be shown in a broader perspective.

5. Acknowledgements

This article is part of a Ph.D. thesis by D.C at the University of Texas at El Paso under the supervision of Christopher Andronicos. D.C. thanks to Dr. Andronicos for his field assistance and his suggestions in an early version of this manuscript. The authors want to acknowledge Dr. Angel Nieto Samaniego and Dra. Elisa Fitz Díaz for their constructive revision. This article was partly supported by CICESE – La Paz, internal project number 691106.

6. References

- Allmendinger, R.W., Cardozo, N.C., Fisher, D., 2012, *Structural Geology Algorithms: Vectors & Tensors*: Cambridge, England, Cambridge University Press, 289 pp.
- Aldrich, M.J.Jr., Chapin, C.E., Laughlin, A.W., 1986, Stress history and tectonic development of the Rio Grande rift, New Mexico: *Journal of Geophysical Research*, 91, 6199–6211.
- Angelier, J., 1984, Tectonic analysis of fault slip data set: *Journal of Geophysical Research*, 89, 5835–5848.
- Axen, G.J., Wernicke, B.P., 1989, On the role of isostasy in the evolution of normal fault systems: *Geology*, 17, 775–776.
- Baldrige, W.S., Damon, P.E., Shafiqullah, R., Bridwell, R.J., 1980, *Evolution of the central Rio Grande rift, New Mexico*: New

- Potassium – Argon ages. *Earth Planetary Science Letters*, 51, 309–321.
- Baldrige, W.S., Keller, G.R., Haak, V., Wendlandt, E., Jiracek, G.R., Olsen, K.H., 1995, Rio Grande rift, *in* Olsen, K.H. (ed.), *Continental rifts: Evolution, Structure and Tectonics*, *Developments in Geotectonics*, 25, 233–276.
- Carciumaru, D., Ortega, R., 2011, On the origin of low angle normal faulting in the southern Rio Grande Rift, *Geofísica Internacional* 50-2, 9–22.
- Chapin, C.E., Cather, S.M., 1994, Tectonic setting of the axial basins of the northern and central Rio Grande rift, *in* Keller, G.R., Cather, S.M. (eds.), *Basins of the Rio Grande Rift: Structure, Stratigraphy and Tectonic setting*, *Geological society of America Special Paper*, 291, 5–25.
- Collins, E.W., Raney, J.A., 1994, Tertiary and Quaternary tectonics of the Hueco Bolson, Trans-Pecos Texas and Chihuahua, Mexico: *Geological Society of America, Special Paper* 291, 265–280.
- Davis, D., Suppe, J., Dahlen, F.A., 1983, Mechanics of Fold-and-Thrust Belts and Accretionary Wedges: *Journal of Geophysical Research*, 88, B2, 1153–1172.
- Engelder, T., 1999, Transitional-tensile fracture propagation: a status report: *Journal of Structural Geology*, 21, 1049–1055.
- Erslev, E.A., 2001, Multistage, multidirectional tertiary shortening and compression in north-central New Mexico: *Geological Society of America Bulletin*, 113, 63–74.
- Figures, S.H., 1987, *Structural geology and geophysics of the Pipeline complex, Northern Franklin Mountains: El Paso, U.S.A*, University of Texas at El Paso, PhD dissertation, 279 p.
- Harbour, R.L., 1972, *Geology of the northern Franklin Mountains, Texas and New Mexico: U.S Geological Survey Bulletin* 1298, 129 p.
- Hardy, S., McClay, K., 1999, Kinematic modelling of extensional fault-propagation folding: *Journal of Structural Geology*, 21, 695–702.
- Henry, C.D., Price, J.G., 1984, Variations in caldera development in the Tertiary volcanic field of Trans – Pecos, Texas: *Journal of Geophysical Research*, 89, 8765–8786.
- House, M.A., Kelley, S.A., 2003, Refining the footwall cooling history of a rift flank uplift, Rio Grande rift, New Mexico, *Tectonics*, 22, 5, doi:10.1029/2002TC001418.
- Keller, G.R., Morgan, P., Seager, W.R., 1990, Crustal structure, gravity anomalies and heat flow in the southern Rio Grande rift and their relationship to extensional tectonics: *Tectonophysics*, 174, 21–37.
- Kelley, S.A., Matheny, J.P., 1983, *Geologic Map 54, scale 1:24000, Geology of Anthony Gap quadrangle, Dona Ana County: New Mexico, New Mexico Bureau Mineral and Resources. Color map with text.*
- Lewis, C., Baldrige, W.S., 1994, Half – grabens, accommodation zones, and shoulder uplifts in the Ladron Peak – Sierra Lucero area, *in* Keller, G.R., Cather, S.M. (eds.), *Basins of the Rio Grande Rift: Structure, Stratigraphy and Tectonic setting: Geological Society of America Special Paper*, 291, 135–155.
- Lueth, V.W., Goodell, P.C., Heizler, M.T., Peters, L., 1998, Geochemistry, geochronology and tectonic implications of jarosite mineralization in the northern Franklin Mountains, Dona County, New Mexico: *New Mexico Geological Society Guidebook*, 49, 309–316.
- Marret, R., Allmendinger, R.W., 1990, Kinematic analysis of fault-slip data: *Journal of Structural Geology*, 12, 973–986.
- Michael, A.J., 1984, Determination of stress from slip data: faults and folds: *Journal of Geophysical Research*, 89, 11, 517–11, 526.
- Morgan, P., Golombek, M.P., 1984, Factors controlling the phases and styles of extension in the northern Rio Grande rift: *New Mexico Geological Society Guidebook*, 35th Field Conference, 13–19.
- Morgan, P., Seager, W.R., Golombek, M.P., 1986, Cenozoic thermal, mechanical and tectonic

- evolution of the Rio Grande rift: *Journal of Geophysical Research*, 91, 6263–6276.
- Ramsay, J.G., Huber, M.I., 1983, *The Techniques of Modern Structural Geology. Strain analysis*, vol. 1: San Diego, U.S.A, Academic Press.
- Ramsay, J.G., Huber, M.I., 1987, *The Techniques of Modern Structural Geology. Folds and Fracture*, vol. 2: San Diego, U.S.A, Academic Press, 329 p.
- Ramsay, J.G., Lisle, R.J., 1997, *The Techniques of Modern Structural Geology. Applications of continuum mechanics in structural geology*, vol. 3: San Diego, U.S.A, Academic Press. 1061 p.
- Ramsay, J.M., Chester, F.M., 2002, Hybrid fracture and the transition from extension fracture to shear fracture: *Nature*, 428, 63–68.
- Ruiz, A., 2004, An integrated analysis of the Laramide orogeny, and the effects of the Rio Grande rift in southern New Mexico and west Texas: El Paso, U.S.A, University of Texas at El Paso, MS dissertation.
- Salyards, S., Ni, J.F., Aldrich, M.J., 1994, Variations in paleomagnetic rotations and kinematics of the north – central Rio Grande rift, New Mexico, in Keller, G.R., Cather, S.M. (eds.), *Basins of Rio Grande rift: Structure, Stratigraphy and Tectonic setting*: Geological Society of America Special Paper, 291, 59–71.
- Seager, W.R., 1981, *Geology of the Organ Mountains and southern San Andres Mountains*, New Mexico: New Mexico Bureau of Mines and Mineral Resources Memoir, 36, 1–97.
- Seager, W.R., Shafiqullah, M., Hawly, J.W., Mandarvin, R.F., 1984, New K-Ar dates from basalts and the evolution of the southern Rio Grande rift: *Geological Society of America Bulletin*, 95, 87–99.
- Seager, W.R., Mack, G.H., 1994, *Geology of East Potrillo Mountains and vicinity, Dona Ana County, New Mexico*: New Mexico Bureau of Mines and Mineral Resources Bulletin, 113, 1–27.
- Shaw, J.H., Hook, S.C., Sitohang, E.P., 1997, Extensional fault bend folding and synrift deposition: An example from the Central Sumatra Basin, Indonesia: *American Association of Petroleum Geologists Bulletin*, 81, 367–379.
- Suppe, J., 1985, *Principles of structural geology*: New Jersey, U.S.A. Prentice Hall, 560 p.
- Suppe, J., Medwedeff, D., 1990, Geometry and kinematics of fault-propagation folding: *Eclogae Geologicae Helveticae*, 83, 409–454.
- Yamaji, A., 2000, The multiple inverse method: A new technique to separate stresses from heterogeneous fault-slip data. *Journal of Structural Geology*, 22, 441–452.
- Zoback, M.L., Zoback, M.D., 1980, State of stress in the conterminous United States: *Journal of Geophysical Research*, 85, 6113–6156.
- Zoback, M.L., Zoback, M.D., 1989, Tectonic stress field of the continental United States. *Geological Society of America Memoir*, 172, 523–539.